The Structural and Metamorphic Evolution of the Neoproterozoic Basement in Jebel Ja'alan, East Oman

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THE STRUCTURAL AND METAMORPHIC EVOLUTION OF THE NEOPROTEROZOIC BASEMENT IN JEBEL JA'ALAN, EAST OMAN

RUNNING TITLE: NEOPROTEROZOIC EVOLUTION OF JEBEL JA'ALAN

ABSTRACT

Jebel Ja'alan (east Oman) displays some of the best exposed and easternmost basement rock in the country. It comprises metasedimentary and intrusive igneous rocks, interpreted to have been generated within the Mozambique Ocean at the margin of Neoproterozoic India. The metamorphic conditions experienced by the basement and implications these conditions have for tectonic models of the region were, until now, poorly understood. The aim of this paper is to constrain these conditions in order to test the hypothesis that the basement of Jebel Ja'alan formed in a Neoproterozoic volcanic arc and unravel the relationship between the structural and metamorphic evolution of the region.

Phase equilibria modelling constrains peak metamorphic conditions to c. 670–700 °C and 4.5–6 kbar, following a clockwise P-T path. These conditions are not exclusive to an arc environment but are suggested to represent one due to current and previous interpretations of basement formation based on its geochemistry. U–Pb monazite age data of Hassan Schist samples yields a weighted average age of 833 ± 15 Ma, interpreted to be the age of near peak metamorphism, and is supported by 40 Ar– 39 Ar muscovite age data, which yields a plateau age of 830 ± 6 Ma. The age data collected is shown to be older than that previously gathered for basement in the country's south and is interpreted to represent the Tonian accretion of arc terranes. Mapping of structures in Jebel Ja'alan reveals two phases of deformation, the first involving north-south directed compression, interpreted on the basis of field and petrographic observations to have occurred contemporaneous to or slightly after peak metamorphism. The second phase of deformation involved east-west directed compression, timing of this is difficult to constrain though the reported presence of similar structures within the overlying sedimentary rock suggests the deformation occurred after the Maastrichtian.

KEYWORDS

Structure; Metamorphism; Geochronology; Pseudosection; Oman; Arabian-Nubian Shield; Jebel Ja'alan; Neoproterozoic

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Hassan Schist (JA15-40) with weighted mean 206 Pb/ 238 U age of 838 ± 11 Ma, obtained by excluding two young analyses interpreted to represent either radiogenic Pb loss or a limited phase of monazite growth; d) Hassan schist (JA15-43) with no weighted mean Figure 8. Probability density plot for monazite grains found within Hassan Schist. Ages are plotted as: a) ${}^{206}Pb/{}^{207}Pb$, weighted average of 845 ± 19 Ma; b) ${}^{206}Pb/{}^{238}U$, weighted average of 830.1 \pm 7.3; c) ²⁰⁶Pb/²³⁸U weighted average of 838 \pm 11; d) ²⁰⁶Pb/²³⁸U, with Figure 9. ⁴⁰Ar-³⁹Ar muscovite age data plot for Ja'alan Granite sample JA15-03. The sample displays a weighted average plateau age of 831 ± 15 Ma (MSWD 0.30), which Figure 10. ⁴⁰Ar-³⁹Ar muscovite age data plot for Ja'alan Granite sample JA15-04. The Figure 11. Photomicrographs of petrological relationships within the Hassan Schist and Ali Gneiss units. a) JA15-04 (Hassan Schist) showing fabric defined by biotite and muscovite, biotite and muscovite in contact with quartz, and inclusions of biotite within quartz. b) JA15-06 (Hassan Schist) showing large muscovite grains, interpreted to form part of the peak assemblage, sillimanite rosettes, separation of muscovite and sillimanite by quartz, and the replacement of large muscovite grains with very fine muscovite grains, interpreted to signify the breakdown of muscovite. c) JA15-06 (Hassan Schist) showing staurolite within the muscovite–biotite–quartz–sillimanite matrix, sillimanite rosettes, and fabric defined by biotite and muscovite. d) JA15-36 (Ali Gneiss) showing inclusions of biotite and muscovite within quartz, replacement of large muscovite grains with finer muscovite grains, crosshatch twinning in plagioclase, and separation of biotite and muscovite by quartz. e) JA15-36 (Ali Gneiss) showing fabric defined by muscovite, biotite and sillimanite, inclusions of biotite within quartz, inclusions of quartz within biotite and muscovite, and replacement of large muscovite grains with finer muscovite grains. f) JA15-36 (Ali Gneiss) showing separation of biotite and plagioclase by quartz, and replacement of large muscovite grains with finer muscovite grains. g) JA15-41 (Hassan Schist) showing replacement of large muscovite grains with finer muscovite grains, sillimanite rosettes, and separation of muscovite and sillimanite by quartz. h) JA15-43 (Hassan Schist) showing fabric defined by biotite, muscovite and sillimanite wrapping around quartz and larger muscovite grains, plagioclase featuring simple twinning and quartz inclusions, and randomly oriented biotite separated by quartz. i) JA15-43 (Hassan Schist) showing a strong fabric defined by biotite, muscovite and sillimanite which wraps around quartz grains, replacement of large muscovite grains with finer muscovite grains, separation of muscovite and biotite by quartz, and Figure 12. Microprobe image showing the compositional transect undertaken on a garnet grain in Hassan Schist sample JA15-41. Accompanying end-member cation Figure 13. Microprobe image showing the compositional transect undertaken on a garnet grain in Hassan Schist sample JA15-43. Accompanying end member cation Figure 14. $T-M_0$ pseudosection for sample JA15-36. Compositions given are in mol%. The red line represents the Fe_2O_3 value used for P-T modelling of sample JA15-36... 42 Figure 15. $T-M_0$ pseudosection for sample JA15-04 (and by proxy JA15-43). Compositions given are in mol%. The red line for JA15-04 represents the Fe_2O_3 value

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Figure 23. Time-space plot for Terranes in the Arabian-Nubian Shield and locations within Oman, with a gap between locations signifying a >800 km gap between the closest basement outcrops. A location map of terranes displayed in the plot is presented in Figure 24. It can be observed that both protolith ages and timing of suturing of terranes decreases to the modern day east of the ANS, reflecting amalgamation of younger arc terranes. The opposite trend can be seen within the Omani basement, with ages decreasing towards the modern day west. Furthermore, it can also be seen that accretion in Oman finished earlier than in the ANS, reflected in the earlier deposition of sediments. Ages in the ANS are from Johnson et al. (2011), Mirbat ages from Figure 24. Location map of the Arabian-Nubian Shield (ANS) and Oman (adapted from Blades et al. (2015)). The map shows terranes and locations used for the time-space plot Table 1. Summary of previous geochronological data for crystalline basement lithologies within Oman, adapted from (Rantakokko et al. 2014). Ages in Ma $\pm 2 \sigma$. WR: Whole rock, Mineral abbreviations from (Whitney & Evans 2010). References: a) (Bowring et al. 2007); b) (Mercolli et al. 2006); c) (Rantakokko et al. 2014); d) (Gass et
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