THE QUATERNARY GEOLOGY OF UPPER SPENCER GULF

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Plate 2 Foraminifera

Plate 3 Foraminifera

ABSTRACT

Using seismic reflection profiles and core data a thin sequence of Holocene and Pleistocene sediment was studied in Spencer Gulf, between Whyalla and Port Pirie, and again near Port Augusta. Six seismic sequences were recognized and these were related to stratigraphic units deposited with changing Pleistocene sea levels in the Gulf. recognized within the Pleistocene sediments include aeolian, supratidal, beach and subtidal sediments. The marine Holocene sediments within the cores consist of basically one facies: the subtidal facies generally coarsening upwards. Detailed sediment particle analyses were carried out on the top 2 cm. of the Holocene. It was noted that a reciprocal relationship exists between the proportion of quartz and bivalves, quartz and foraminifera, and between quartz and total calcium carbonate. A positive relationship is noted between bivalves, foraminifera and total carbonate. It was also observed that the sedimentary assemblage, both organic and non-organic, differed within the two regions studied, with the northern Gulf being richer in quartz and poorer in biota than the southern Gulf. Studies of the foraminifera indicated that different assemblages lived in different depths of water. It would seem, therefore, that water depth, and the factors associated with water depth, have a great influence on the distribution of fauna within Spencer Gulf.

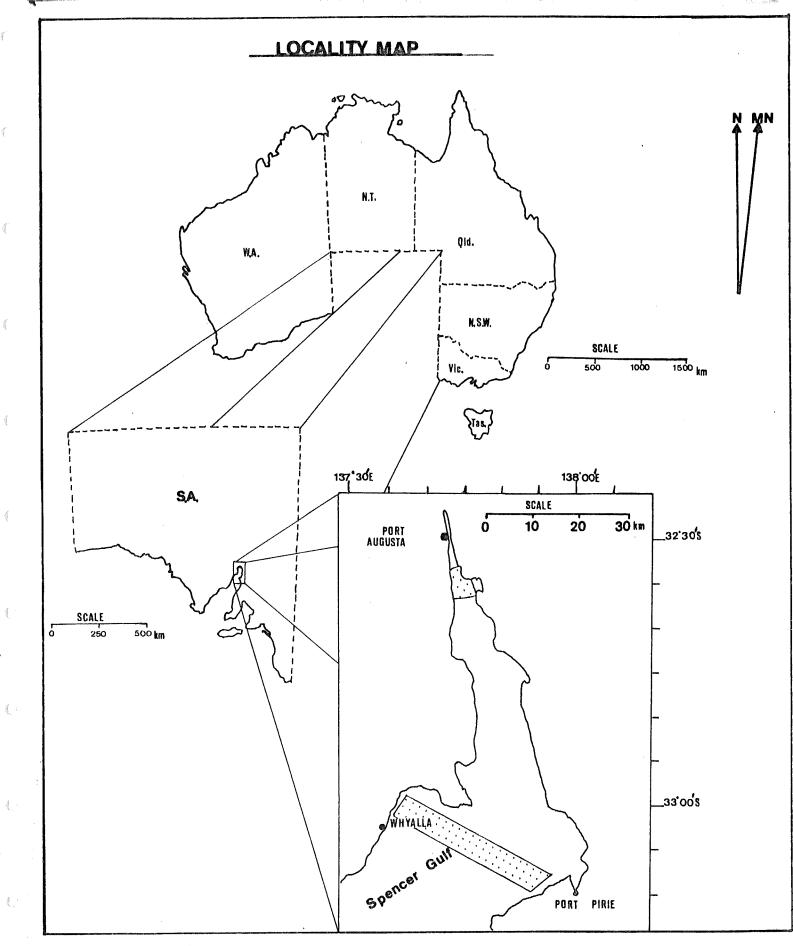


FIGURE 1: LOCALITY MAP.

1. INTRODUCTION

This study has been prompted by work done by the Environmental Studies Department and Sedimentologists at the Geology Department of Adelaide University, within Spencer Gulf over the last four years.

Two areas were chosen for comparison. One within the waters of the Gulf near Whyalla and the other within the waters of the Gulf near Port Augusta. Fig.1.

The purposes of this study were to interpret seismic reflection profiles and core data taken from the Gulf. A study of the data enabled a stratigraphy of the Holocene and Pleistocene sediments to be constituted. Six sequences The oldest, sequence V, is only noted in were observed. the seismic record, not in the cores, and is an unknown quantity, but from its seismic characteristics is probably similar to sequence IV. Sequence IV is red-coloured sediment of intertidal, supratidal, subtidal, and estuarine Sequences II and III are identical in being basically of sabkha and intertidal facies with minor subtidal deposits. Sequence I is divided in two. Sequence IA is grey and of subtidal factes. Sequence IB is reworked sequence II, III or IV. Sequence VI is alluvial material of various ages. Ages are suggested for these transgressive sequences.

Further research was carried out on the Holocene sediment. The two regions observed showed obvious differences in sediment type. The Port Augusta region shows a greater terrigenous and lower marine biota content. The Whyalla region showed the opposite. The foraminiferal distribution was also studied and the assemblages changed from one region to the other and from shoal to trough within the regions. Variables associated with water depth are thought to control this distribution.

2. REGIONAL GEOLOGY AND ENVIRONMENT

2.1 REGIONAL GEOLOGY

The study area lies within the Torrens Sunkland, a major downfaulted block, lying between two horsts, the Flinders Ranges and the Stuart Shelf.

The horst blocks on either side of the Gulf have an equivalent lithology and strati-graphic sequence, i.e. formations mapped can be related in time and rock type. The lithology is a Proterozoic sequence of sandstones, shales and quartzites and minor dolomite and limestone members.

Structurally, however, the blocks are quite different. The Flinders Ranges, including Mt. Grainger, are strongly folded whereas the Stuart Shelf plateau is flat bedded. There is evidence from bore log data and previously published maps that a minor horst block exists within the graben. The associated faults probably run in a north-south direction along the edges of the coastal monadnocks, Mt. Grainger, Mt. Gullet, and Mt. Mambray.

Within the sunkland of Spencer Gulf a thick sequence of Tertiary sediment is overlain by Quaternary piedmont sandy clays and gravels. Superimposed on this is a relatively thin veneer of sediment which is the subject of this study.

2.2 ENVIRONMENT

2.21 Climate

Records are available at Port Pirie (95 years), Port Augusta (100 years) and Whyalla (68 years) for rainfall and temperature.

Rainfall in these areas is quite low. Mean rainfall at Fort Augusta is 242 mm., at Port Firie 342 mm. and at Whyalla 269 mm. The influence of the Flinders Ranges on the rainfall of these coastal areas is very significant. Records show that most of the rain falls during the months April to October, coinciding roughly to the winter months. Winter rains are more noticeable at

Port Pirie than at Port Augusta and Whyalla where the spread of rainfall is more even over the year. Summers are generally dry.

Mean maximum and minimum temperature trends at Port Augusta and Port Pirie are similar and within 1°C throughout the year. The summer months are very hot; Port Pirie having a mean of 31.8°C and Port Augusta a mean of 32.1°C for the month of January. Winters are generally cool; Port Pirie having a mean of 16.3°C and Port Augusta a mean of 17.1°C for the month of July.

at the Port Augusta Power Station from the period 1961 to 1965 inclusive, at 9.00 a.m. and 3.00 p.m., show the following trends. During the months of June, July, and August the winds are dominantly north-north-west to north, whereas for the rest of the year, the winds are dominantly southsouth-east to south. The strength of these winds varies considerably, but is on the average between 3.5-8 km./hr. and always stronger in the afternoon than in the morning.

In conclusion, it can be seen that combined with high summer temperatures and low rainfall these areas have high evaporation rates (b/w 150 and 175 cm. per year). In summer there would be a consequent lowering of the water table in the sediments and an upward movement of water.

2.22 Tides

Tides are strongly affected by winds. At Port Augusta additional sea-level rises of up to 1.3 m. accompany south-west to south winds, while north winds lower the sea level up to 0.5 m. Mean tidal range is in the order of 1.5 m., while spring tides (range 3.6 m.) can raise the sea level more than 0.5 m. above mean hightide level.

2.23 Marine Flora

The flora of the sublittoral zone is dominated by marine angiosperms, which exist as extensive banks. These seagrasses grow in shallow water in

fine sand and silt sediments. Two dominant species of seagrass grow in the sublittoral zone and their distribution is related to water movement and light penetration.

Zostera muelleri grows at depths slightly above mean low tide to 1 m. below. The plant is very thin-leaved, between 1-2 mm. broad and 6-15 cm. long, and grows in very dense patches. It has a large rhizome-root system. Hormosira banksii is occasionally found growing with Zostera muelleri.

generally in a monospecific community but sometimes it is seen associated with Amphibolis antartica and Cymodocea griffithsi. The leaves are from 3-8 mm. wide and up to 1 m. long. P.australis grows only in areas which are permanently submerged and at depths of 0.3-4 m. below mean low water. At greater depths the broad leaf P.australis is replaced by a narrow-leaved species showing distinct morphological differences. Mixed communities of the two forms occur in transitional zones. P.australis is supported by a very large rhizome-root system.

2.24 Fauna

In the sublittoral zone the seagrasses <u>Posidonia sp.</u> and <u>Cymodocea sp.</u> support many epiphytes which grow on the broad leaves. The epibiota include foraminifera, gastropods, coralline algae and calcareous worm tubes. When the epiphytes die they accumulate in the sediments and are redistributed and broken by burrowing organisms.

The fauna within the Recent sediment is diverse. Comprehensive lists of gastropods and bivalves occurring in South Australian waters are given by CHILD (1968), COTTON (1959, 1961 1964), VERCO (1904-1918) and WIISON AND GILLET (1971).

Gazameda iredalei, Iemintina sipho, Floraconus compressus,
Zeacumanbus sp., Truncatella vincentiano, Argaliste fugitiva,
Phasionella australis, Vexillum plicarium, Philine beachportensis and Triphora sp.. The turreted gastropods are much

more numerous than non-turreted and the species <u>Zeacumantus</u>
is the most prolific of the gastropods.

Bivalves identified include Anadara trapezia, Glycymeris radeans, Chama ruderalis, Katelysia peroni, Austromytilus erosus, Pinna dolobrata, Notochlamys tasmanicus, Macla australis, Mytilus sp., and Limopsis sp..

The foraminifera for statistical purposes have been divided up into groups determined by Genera. The species identified are included.

Group	1		SPIROLOCULIN	1A	angusteoralis,
					antillarum,
					communis,
		and	QUINQUELOCUI	LINA	tropicalis.
Group	2		TRILOCULINA		oblonga,
					rotunda,
					striatotrigonula,
				and	trigonula.
Group	3		ELPHIDIUM		macellum,
•				and	pseudonodosum.
Group	4		PENEROPLIS		planatus,
		and	SPIROLINA		aciculoris.
Group	5		DISCORBIS		dimiadatus,
					<u>vesicularis,</u>
		and	ROSALINA		kennedyi.
Group	6		NUBECULARIA		lucifuga.

Other foraminifera not included in the groupings but which are found in minor numbers in the recent sediment include <u>Cribrobulima polystoma</u>, <u>Miliolinella subrotunda</u>, <u>Bolivina sp.</u>, <u>Guttulina pacifica</u>, <u>Amphicoryna scalaris</u>, <u>Iagena sp.</u>, and <u>Glabratella sp.</u>. No planktonic nor agglutinated foraminifera were observed within the study area, only benthic forms. Thus the assemblage within the study area is totally different to that found in Investigator Strait where both agglutinated, planktonic and benthic foraminifera are found. (pers.comm. J.Cann, 1981)

Aragonite needles are prolific within the study area. Lowenstam and Epstein (1957) observed that the organic tissue of living calcareous algae contains abundant needles of aragonite that are similar to those found in the bottom sediments of the Bahama Platform. The needles found on the Bahama Platform are similar to those found within the Gulf. Therefore it is highly possible that the aragonite needles found within the sediments studied are from the organic tissue of calcareous algae found within the Gulf.

Also found within the recent sediment are fragments of echinoids, both plates and spines, coralline algae, sponge spicules and bryozoans. All of these faunas mentioned here appear in minor numbers within the sediment.

No whole brachiopods were found within the recent sediments. Even within the fragments of shell no brachiopod detritus was observed. It would seem therefore that brachiopods are either small in numbers or non-existent within the study area.

The above list of organisms identified are neither comprehensive nor definitive and much further work needs to be done to complete the lists of all species of biota occurring in the area.

3. SEISMIC REFLECTION AND CORE DATA --

INTERPRETATION

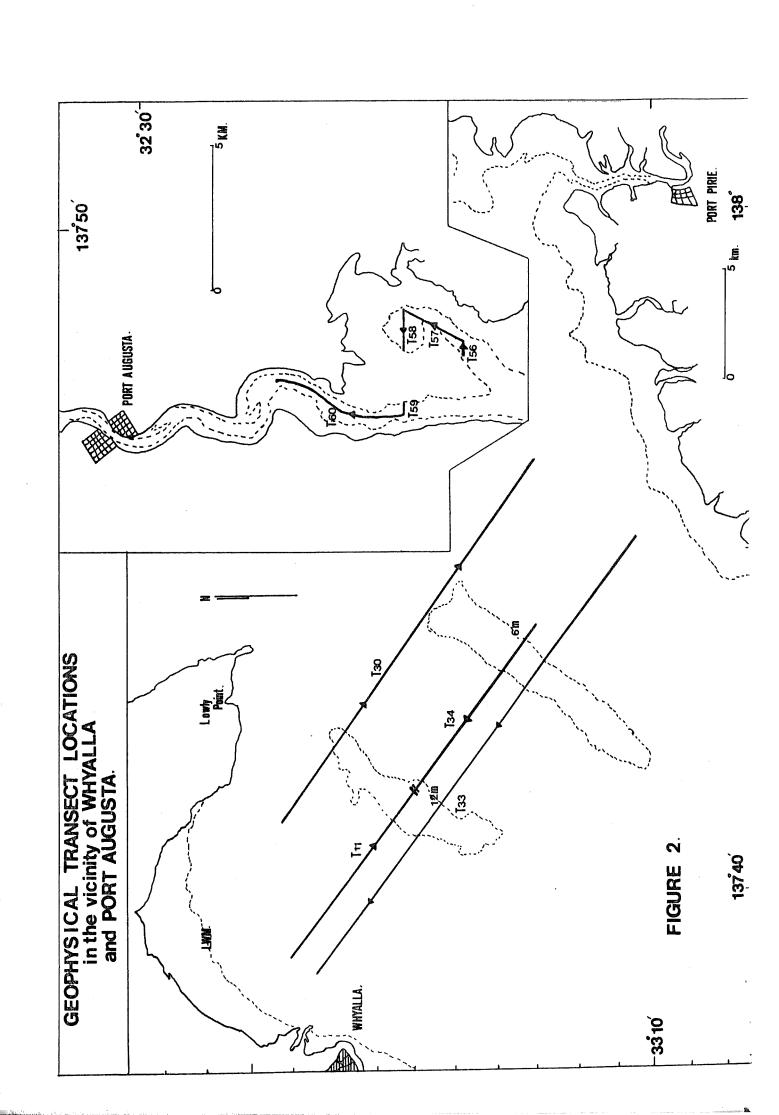
Marine seismic reflection methods were used in the study area. Using seismic reflection methods the structure of subsurface formations may be mapped by measuring the times required for a seismic wave (or pulse), generated in the water by a boomer, to return to the surface after reflection from interfacies between formations having different physical properties. In effect the seismic reflection method delineates subsurface geological structures. DOBRIN (1976).

The geophysical transects focused upon in this study are shown in Fig.2. For equipment used during the geophysical programme see APPENDIX 1.A.

VAIL et al. (1977) look in particular at seismic stratigraphy and much of the following general comments are from their work.

The depositional sequence is an objectively defined stratigraphic unit composed of genetically related strata and bounded by unconformities and their correlative conformities. Using seismic reflection methods the depositional sequences in the subsurface can be delineated. This method can be employed because seismic reflections are composites of the individual reflections generated by surfaces separating strata of differing acoustical properties. Hence reflectors usually correspond with horizons marking the boundary between rocks of markedly different lithology. However, such a boundary does not always occur exactly at a geological horizon of major chronostratigraphic importance, such as the base or top of a system or series, but may be simply a seismic marker horizon which occurs close to the boundary. This problem is overcome by the correlation of seismic and core data.

But, on the whole, the reflectors tend to parallel stratal surfaces and have the same chronostratigraphic significance as stratal surfaces. Therefore it is possible to make chronostratigraphic correlations using seismic reflection patterns.



When a depositional sequence is identified within a seismic section it is labelled a seismic sequence. It is a relatively conformable succession of reflectors within a seismic section interpreted as genetically-related strata. This succession is bounded at its top and base by surfaces of discontinuity marked by reflection terminations and interpreted as unconformities or their correlative conformities.

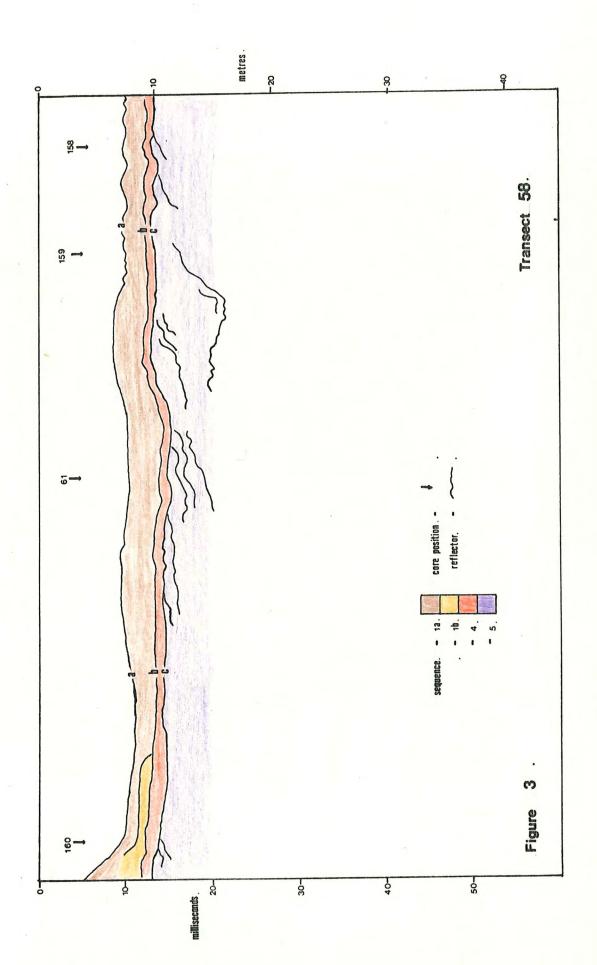
Reflection terminations are the principal factors for recognition of seismic sequence boundaries. Reflection terminations interpreted as stratal terminations include erosional truncation, toplap, onlap and downlap.

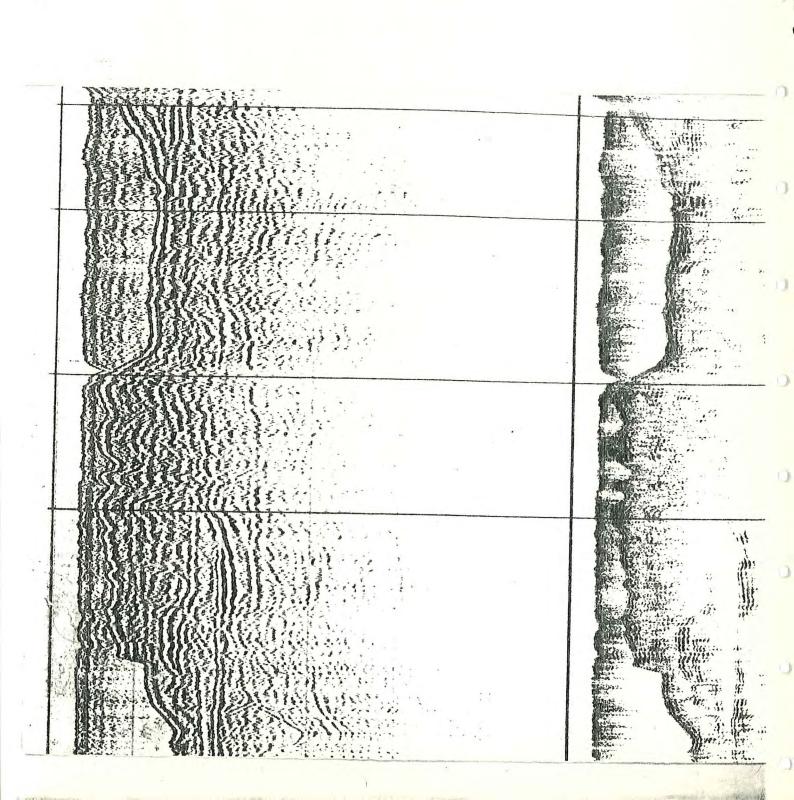
Seismic transects 58 and 56 (Figs. 3 and 4 respectively) are examples of the seismic reflection records, and transparencies overlain show the interpretation.

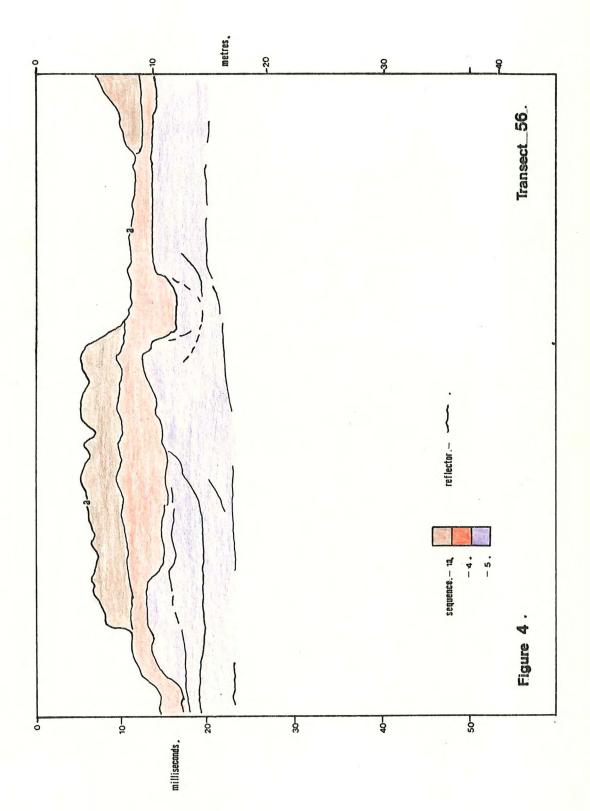
Transect 58 shows 3 sequences. The surface of the Gulf floor is marked by the first strong reflector (a). The next strong reflector (b) marks the base of the first sequence and the initiation of the second sequence. The second sequence is extremely narrow for a third strong reflector (c) marks the base of this sequence within a short period. Below reflector (c) a series of reflectors of a concordant nature and of identical strength are observed lying at angles to reflector (c). Here an actual unconformity can be recognized. Three sequences are delineated each marked at top and base by strong reflectors. The reflective pattern within sequences I and II is basically reflection-free. Within sequence III the reflective pattern is strong and concordant.

Knowing that strong reflectors mark the top and base of sequences and that different sequences usually have different reflective patterns the depositional sequences can be identified. With erosional truncation between the sequences, as seen between sequences II and III, the task of interpretation becomes easier and with cores through these sequences correlations of data can be made and a definite stratigraphy indicated provided there are

LINE 58 (REV)







enough cores.

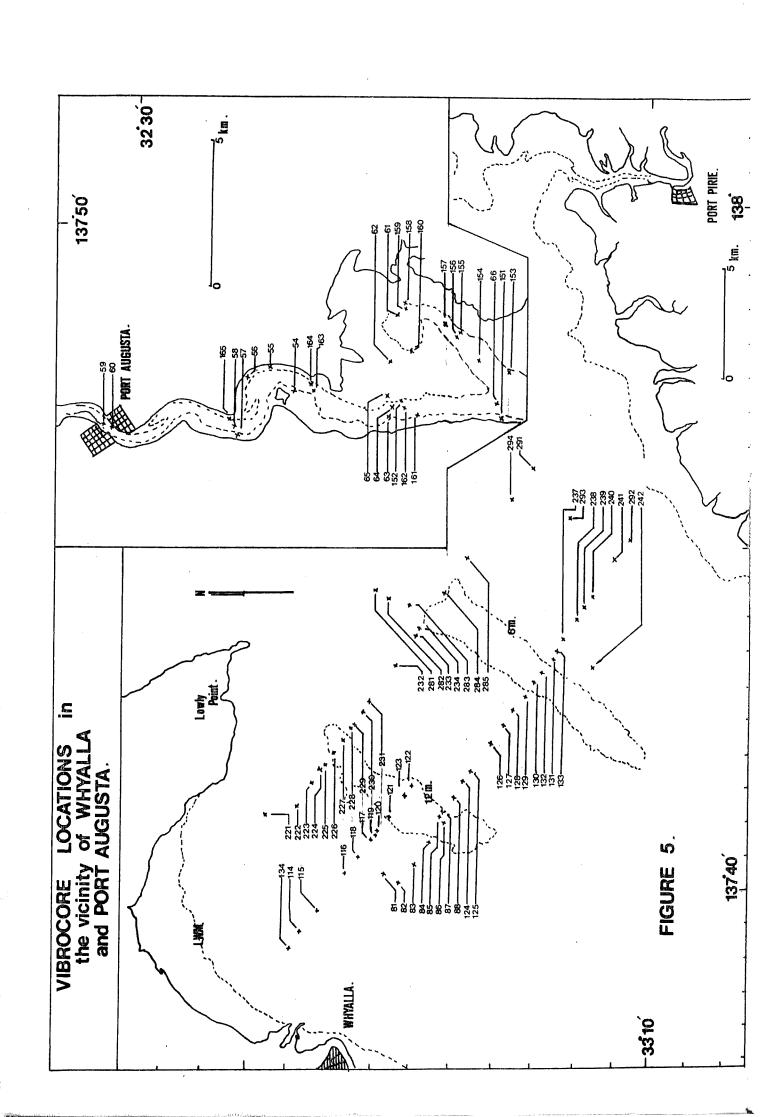
The seismic reflection pro-The upper profile is a low files available give two modes. resolution/high penetration profile while the lower is a high resolution/low penetration profile. The lower profile may look too reflection-free to be of any use but in Transect 56 and others it is invaluable. The low resolution/ high penetration profile 56 is extremely hard to interpret as there are many strong reflectors in close proximity and no differential patterns are observed. Using the high resolution/low penetration profile, boundaries between sequences are identified. Reflector (a) is again strong but so are all the reflectors beneath it. Using the lower profile, sequence I is marked by a darker granular haze. Where this reflective pattern ends, sequence I ends. Using the lower profile again, strong thick reflectors are identified. Using this interpretation and overlaying this on the upper profile a series of sequences can be identified.

Using both high resolution/
low penetration and low resolution/high penetration methods
sequences can be interpreted within the section and with a
number of cores pushed through into these sequences the
stratigraphy and facies of the sediments can be identified.

Figure 5 illustrates the position of the cores taken within the study area. The vibrocoring transects are as near as possible to the geophysical transects, in position. The cores were taken using a Vibrocorer (see APPENDIX 1.B.). Cores of up to 4.5 m. in length were taken. Many cores are of shorter lengths. The cores were cut in half and one side used for sampling, the other described and interpreted. For description and interpretation of selected cores see APPENDIX 2.A.

The facies identified within the study area are of a very flat, shallow water, depositional, coastal origin. The facies can be divided into 3 groups, Supratidal, Intertidal and Subtidal.

The facies classified as Supratidal include dune, sabkha, and supratidal flat. Dunes



are recognized usually because of the high proportion of well-rounded, medium- to well-sorted, frosted quartz. The quartz is usually stained an iron-oxide orange to red. Clay dunes are also identified, consisting of a very fine silt. Often the silt is in pellets. Supratidal flat deposits consist of mainly silt with small percentages of quartz and shell. Sabkhas form in an area of unconsolidated sediment at the margin of a large body of water where the ground water lies at or very close to, the land surface. These conditions permit loss of ground water to the atmosphere by upward transpiration and evaporation, causing ground-water solutes to be deposited at or near the surface (RENFRO, 1974). Both supratidal flats and sabkhas are a grey, cream to green colour and consist of calcareous mud. Within the sabkha sediment gypsum crystals are found.

The facies within the Intertidal division include intertidal flats and beaches. A beach deposit is identified by a well-sorted sand either shell or quartz or both. An intertidal flat deposit has a higher percentage of silt and less quartz and shell than a beach deposit. Usually these facies are oxidized a yellow-to-cream colour.

Within the Subtidal division basically 2 facies are identified, the subtidal and the estuarine. The subtidal deposit is very similar to the intertidal deposit except material, shell and quartz, is less well-sorted and whole organisms such as bivalves should be found. The estuarine deposit consists of mud with small percentages of shell and quartz. Again whole organisms should be found.

Besides these facies alluvial fan material is identified. This consists of red-to-brown muds with quartz. The discerning factor is that it is non-calcareous.

Indicators of breaks, diastems, erosional truncation within the cores are many.
Often between depositional sequences, nodules, often large nodules, taking up the whole core-width, are found. The

nodules are formed of quartz, shell and carbonate cement and are extremely hard. The nodulous sequence may be up to 20 cm. thick. The nodules are formed by the movement of calcium carbonate, within water, up through the sediment through capillary action. The carbonate is left at the near surface as the water is evaporated off. Because of the predominance of quartz within these nodules it could be that, of the nodules formed within dunes and then with transgressive seas, only the nodules were left. Soils quite often form beneath these nodules. BILLINGS (1981) interprets the nodules as pedogenic types of deposits formed in coastal dune situations where they cap weakly-developed truncated soils.

Repetition of a sequence is another indicator of the development of a second sequence. For instance, core 225 (APPENDIX 2.A). At the base of the core a sabkha deposit is identified covered by dune, quartz sand. On top of this another sabkha deposit is identified covered by a dune deposit again. For this sequence to form there has had to have been a regression from sabkha to dune and then a transgression to sabkha facies and then a regression to dune facies again. A transgression has occurred but the only factor identifying this is the repetition of sequences.

Also colour change can be used. The recent material, for instance, is always grey within the Spencer Gulf. Deposits beneath are always of a different colour hence the break is always indicated. Also if one bed is oxidized then it can be definitely said that the oxidized deposit has experienced very shallow water cover to non-existent water cover. If the deposit above is of intertidal-to-subtidal facies, for example, and of a different colour then it obviously has had greater water cover, hence the deposit can be labelled a transgressive deposit, while the red-to-brown oxidized deposit has felt the effects of regression. With regression of the sea the marine sediments of the upper Gulf were left stranded above water level. These deposits were then sub-aerially weathered and

with the addition of terrigenous sand and clay by both fluvial and aeolian agencies became soils. All of the soils found within the cores today are of a light-brown to dark-brown colour. Some of the soils have developed only on the marine sediment as depicted by shells within the soils but other soils show a bimodal quartz distribution whereby on fraction is obviously aeolian and the other of beach or fluviatile origin.

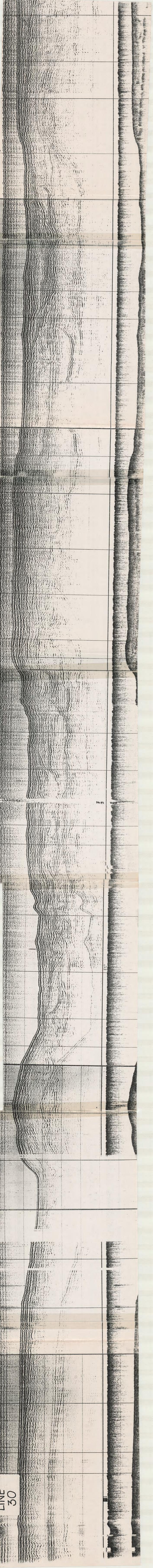
Another factor is the identification of alluvial material between facies. This means that the sea had regressed and alluvial material from the slopes had been washed down on to the marine deposits.

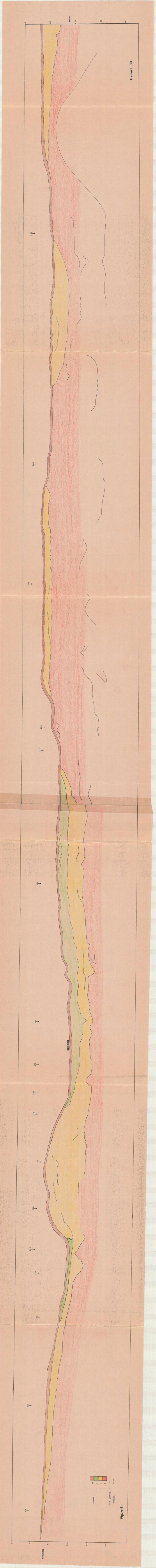
One last factor is the identification of reworking of the sediment indicating, again, a transgression. Regression of the sea leaves the sediments open to weathering and erosion and transgression of the sea also causes erosion hence the breaks between the transgressive sequences are marked by weathering and erosional effects.

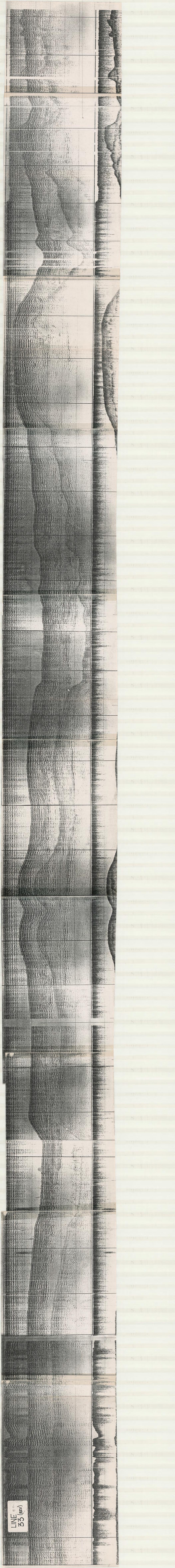
Through the use of seismic reflection profiles the stratigraphy of the Gulf can be approximated. With the use of the core data descriptions and interpretations the stratigraphy is not approximated but definite where there are enough cores.

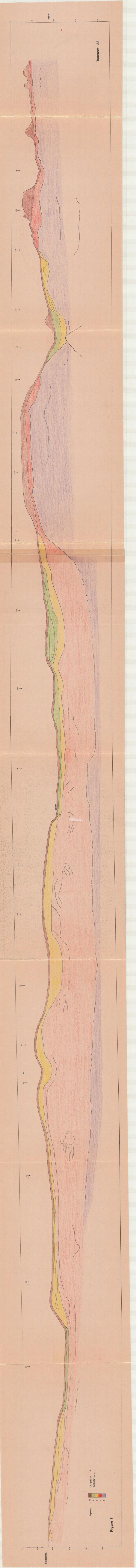
Seismic transects 30, 33, 34, 11, 56, 57, 58, 59 and 60, FIGURES 6, 7, 8, 9, 10, 11, 12, 13, 14 respectively, display the seismic interpretation of the transects in the Port Augusta region, the Whyalla North and Whyalla Central regions. This breakup of the study area can be seen in FIGS. 2 and 5. Along with the seismic interpretation information from the cores has been added and the various depositional sequences indicated by a different colour. Over all, 6 depositional sequences have been recognized using seismic stratigraphy and core information.

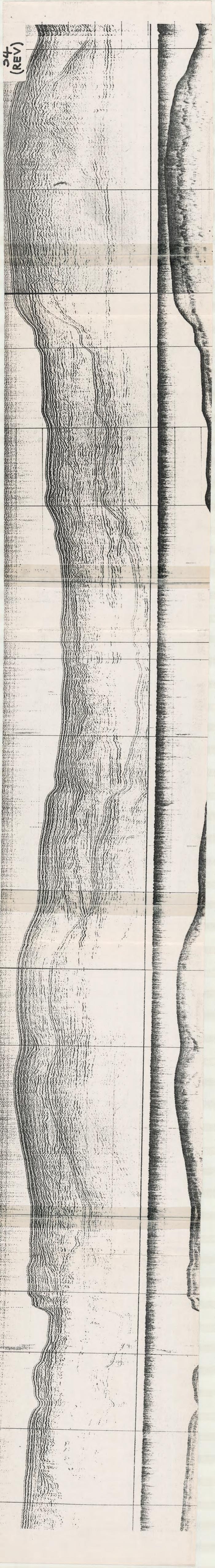
Descriptions of the sequences follow. More information can be obtained from selected core descriptions in APPENDIX 2A.

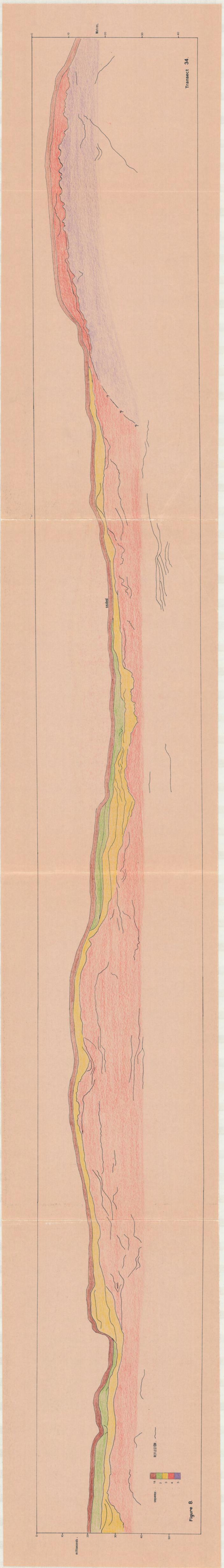


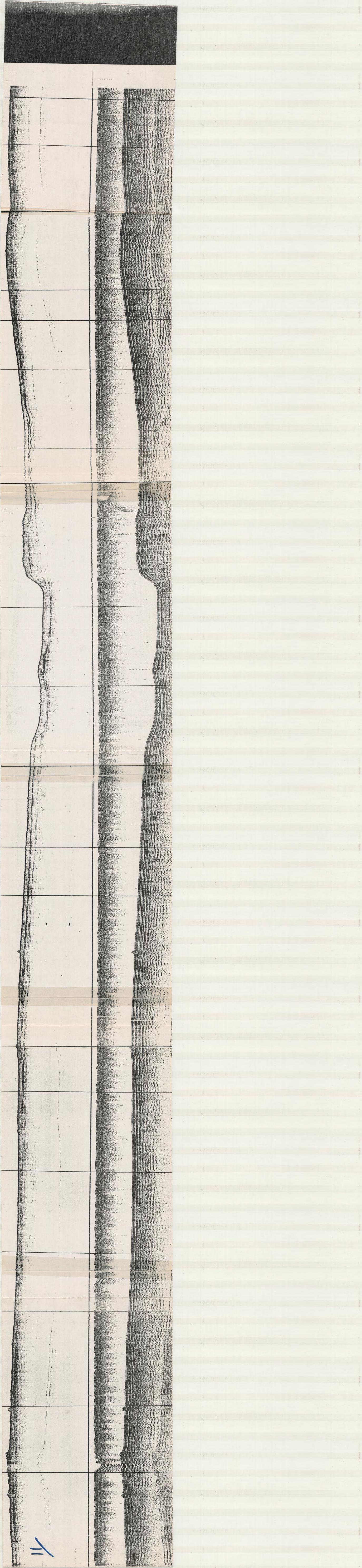


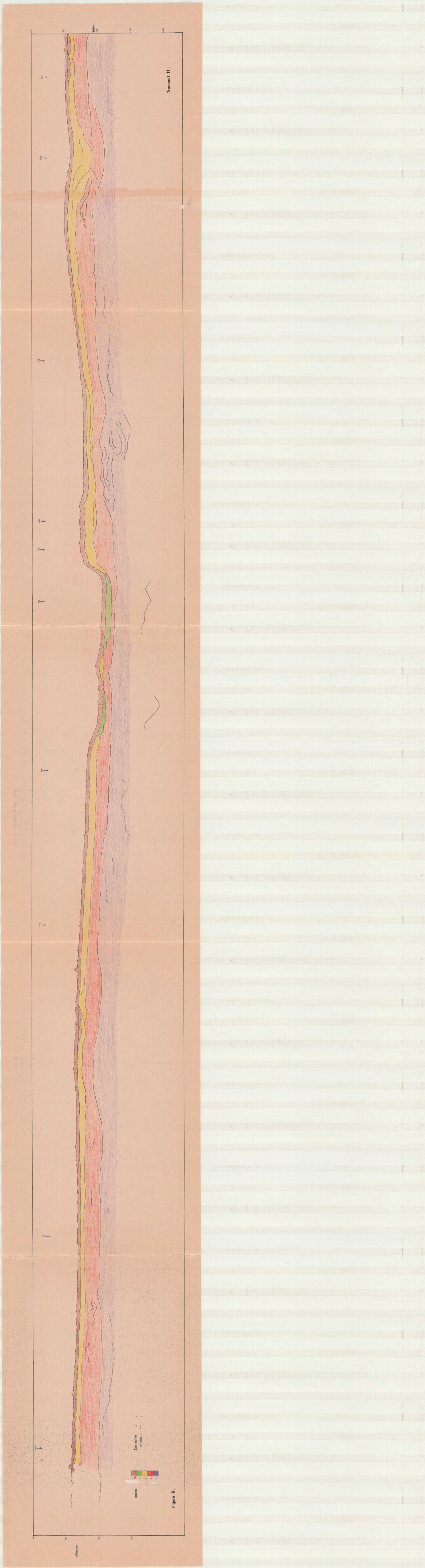


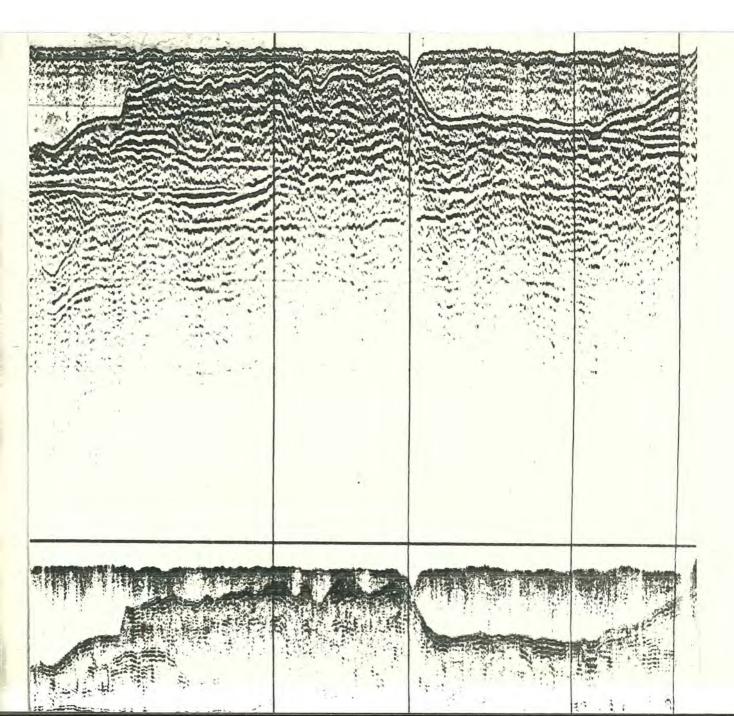


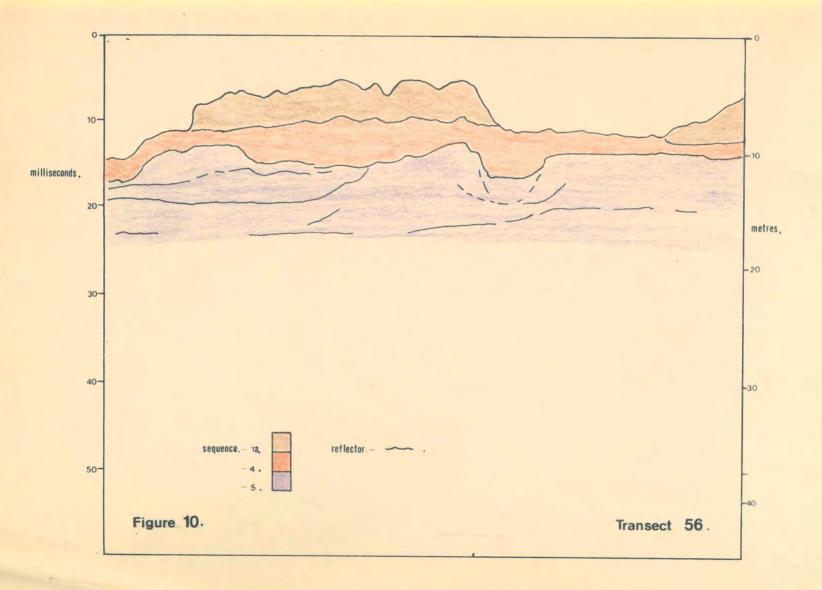




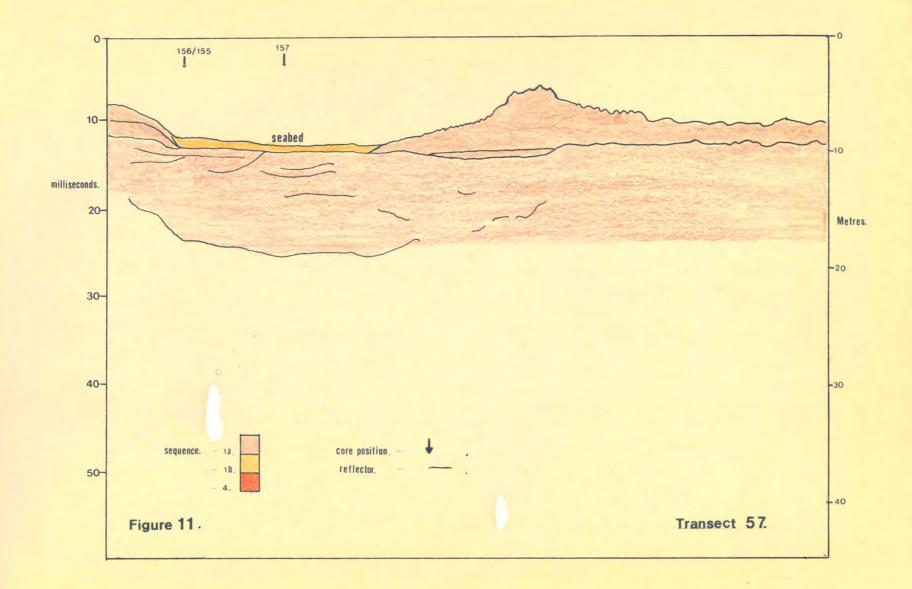




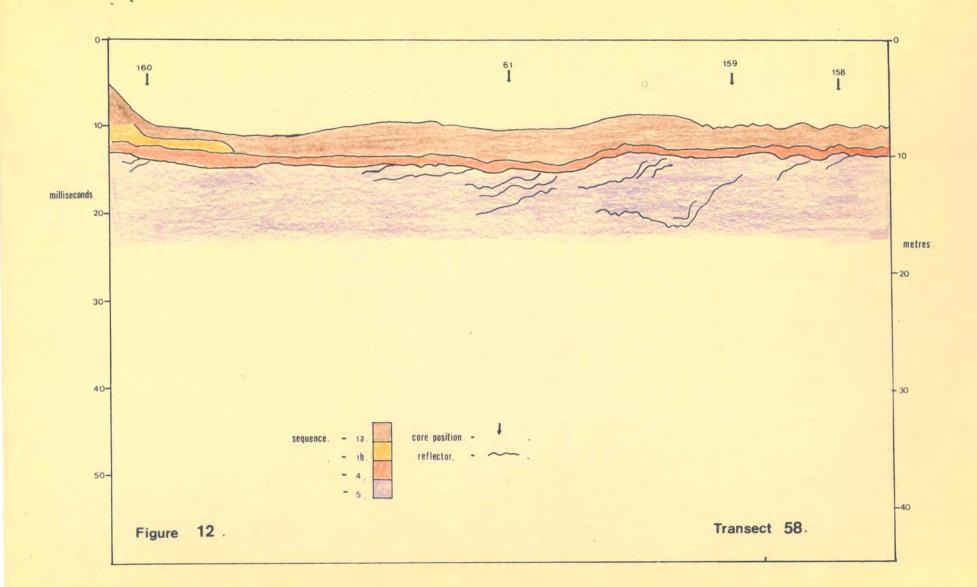


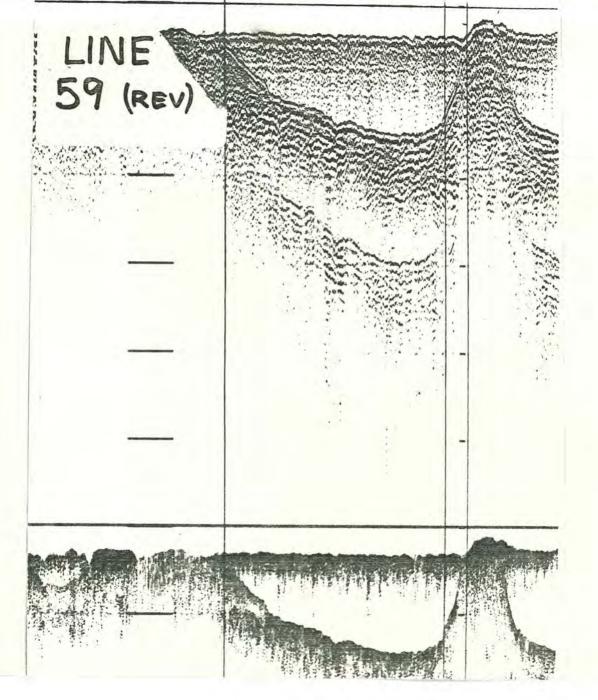


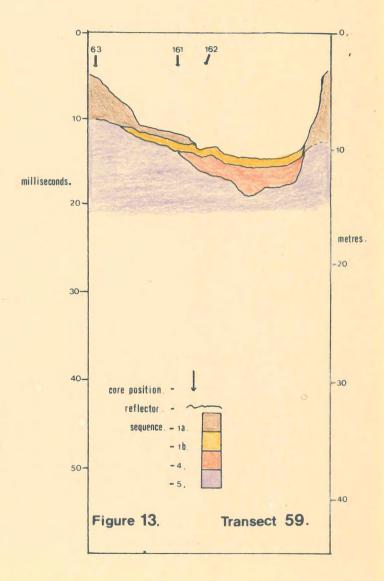
LINE 57		

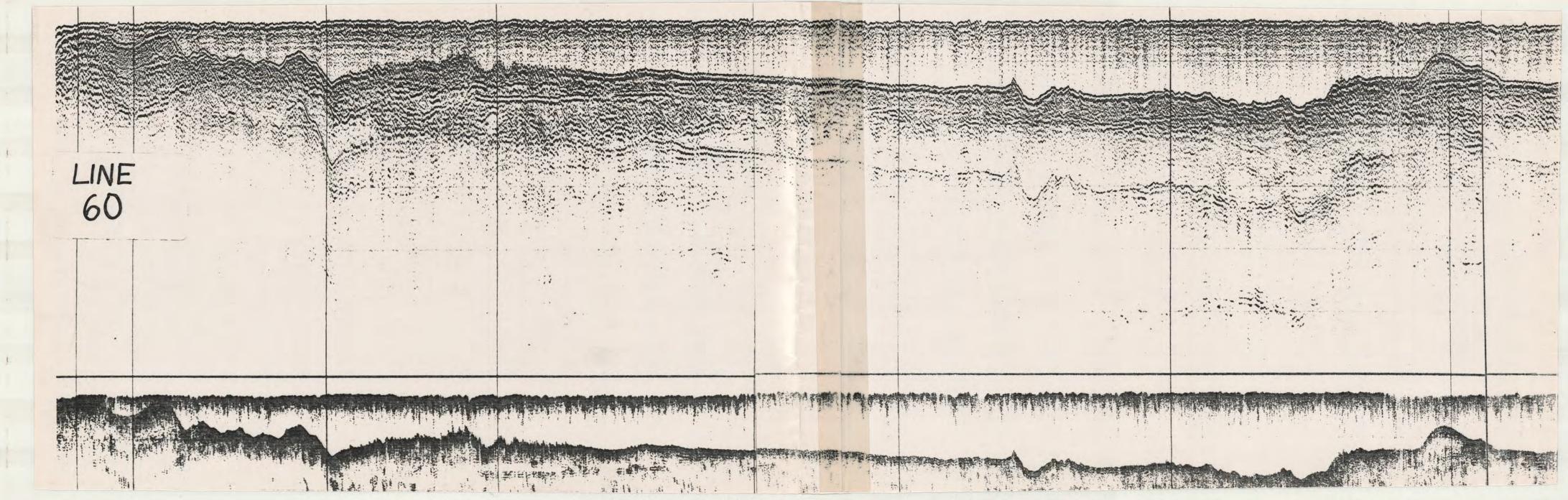


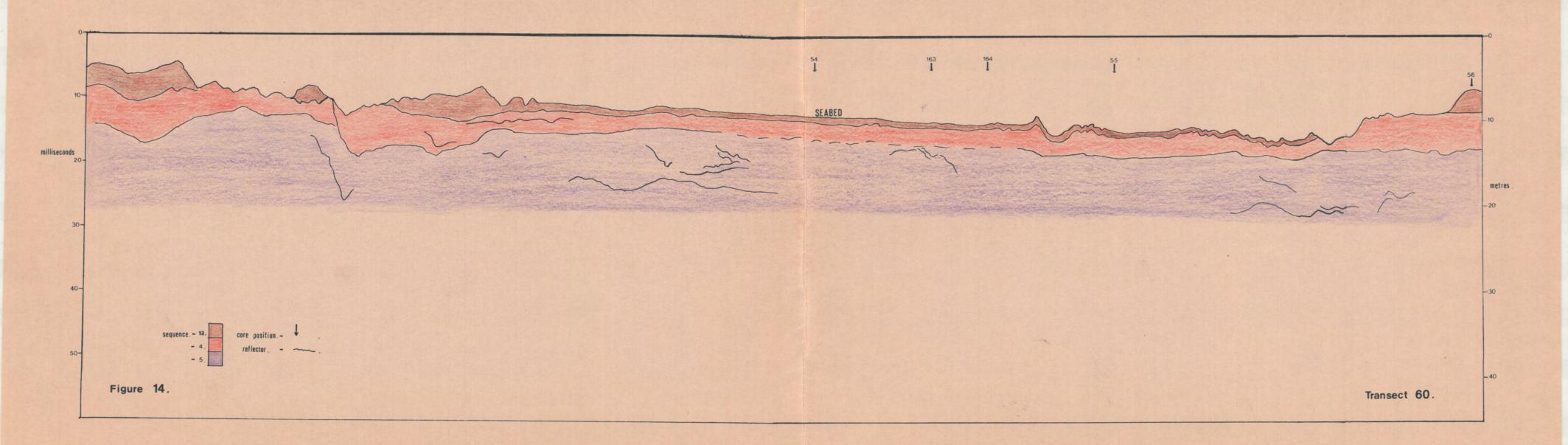
LINE 58 (REV)











LEGEND FOR THE SAND AND ORGANISM DISTRIBUTION

FIGURES: 17, 18 and 19

Quartz

Bivalves

Gastropods

Foraminifera

A Aragonite needles

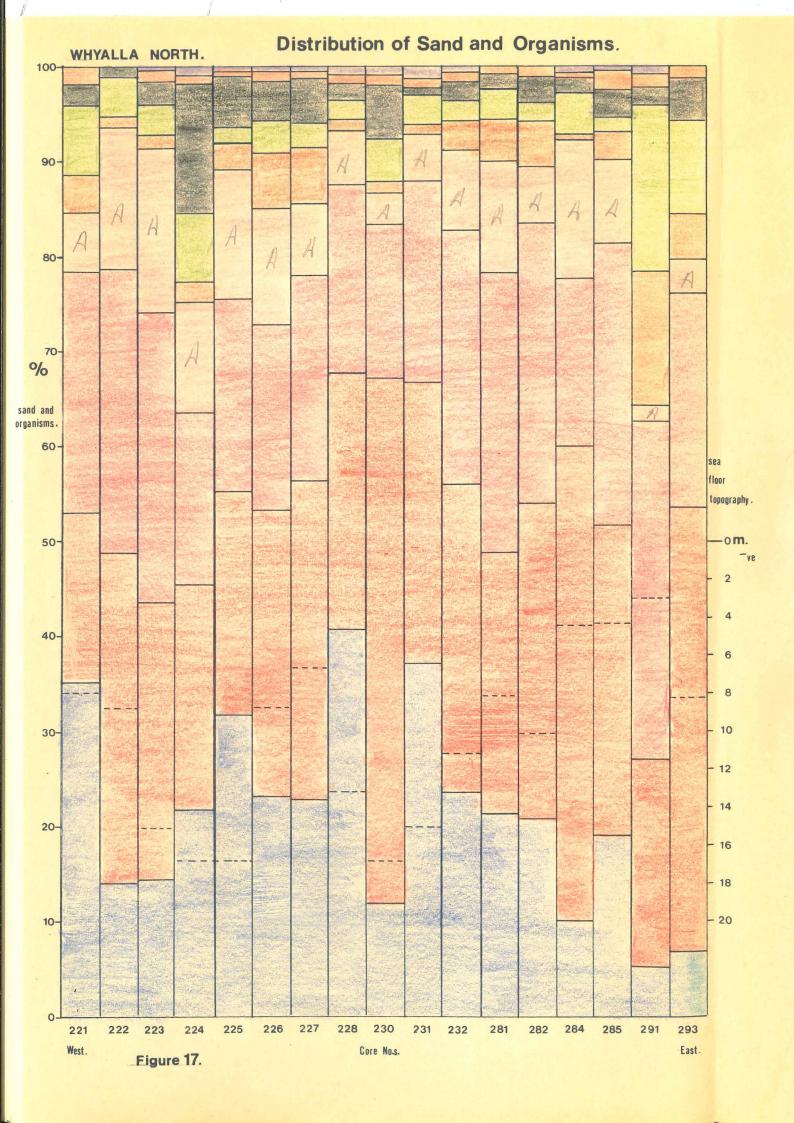
Coralline algae

Bryozoans

Miscellaneous

Echinoid spines and plates

- Line representing sea-floor topography



SEQUENCE V (OLDEST)

Within every region sequence V is known only by its seismic record because few cores pierce into this sediment. In the Port Augusta region sequences IV and V are often separated by alluvial material. Beneath this alluvial material various facies of sediment are observed. On one occasion an extremely wellsorted, polished and frosted, rounded, quartz deposit is noted. This is probably a dune. On another occasion a deposit of 80% quartz and 20% fine silt is identified. could be dune material again but a few fragments of bivalves are found and it is thought to be a beach deposit. An intertidal deposit is identified elsewhere within the record. So definitely aeolian and intertidal deposits exist but possibly more facies would be identified with more and deeper cores. Dipping beds, channels and series of contorted beds are observed within the sequence. The sediment is always dark red to orange in colour. Only in sequence V. particularly transect 58, FIG.12, of the Port Augusta region, are beds seen to lie at a dip, an obvious dip, to younger sediments. It would seem that before sequence IV was deposited there was movement within the study area causing the dip on the beds in sequence V. No other strongly dipping beds are seen in the geophysical sections implying tectonic activity stopped before the deposition of sequence IV.

SEQUENCE IV WHYALLA NORTH

Sequence IV within this region is mostly brown to red shelly silts. Shell fragments including bivalves, gastropods and foraminifera, with the bivalves dominant, is at its greatest 25% of the deposit as is quartz which is polished, subrounded and of mediocre sorting. The silt content ranges from 50% to 100% so once again a sequence is predominantly shelly silts and silts. The facies are subtidal, intertidal, supratidal and estuarine. On occasions the top of sequence IV is reduced. A marine transgression after a regression would initiate this reduction The thickness of this sequence is unknown.

Sequence IV is the oldest sequence defined in Whyalla Nth. Within the sequence relict

channels and monadnocks are observed hence differential erosion has occurred in the past. The boundaries between the various sequences also show differential erosion with a small-scale ridge and valley outline apparent. More channels appear within sequence IV than in today's topography. Actually, since only the top of sequence IV can be identified and not the base, it is quite possible that sequence V is also being studied in this case. Nodules and burrows are common.

WHYALLA CENTRAL

Sequence IV consists of intertidal, subtidal, supratidal and estuarine deposits. Again shelly silts and silts predominate. The supratidal deposits within this sequence have up to 50% subrounded to sub-angular, poorly-sorted, polished, iron-oxide-stained quartz. Little shelly material is present while silt again makes up at least 50% of the deposit.

Fauna within the subtidal, intertidal deposits include both turreted and non-turreted gastropods, coralline algae, echinoid spines, bivalves, foraminifera, aragonite needles, worm tubes, etc., with Marginopora present in the intertidal deposits.

Again silt within this sequence is a dark-red to brown colour. Nodules are found throughout and silt attaches itself to the quartz and shell on many occasions. Everything is iron-oxide stained.

As for the Whyalla North region, channels and monadnocks are apparent within the fourth sequence. Within Whyalla Central, the base of Sequence IV has been identified using seismic stratigraphy in several areas. The boundary between sequences IV and V is erosional as are the boundaries between other sequences.

PORT AUGUSTA

Within this region sequence IV is again either subtidal, intertidal, supratidal or estuarine. Intertidal deposits are prominent with a predominance of quartz, rock fragments and silt with a very small percentage of shell. The quartz and rock fragments are very

coarse. Quartz is sub-angular to mainly subrounded, polished and poorly sorted to well sorted where possibly a dune exists rather than an intertidal deposit. The thickness of sequence IV is variant. Channels and monadnocks are apparent. Beneath the intertidal deposits usually the supratidal facies is found. The estuarine deposits are mainly muds with some coarse shell.

Sequence IV is extremely extensive and is the major factor in the topography of the Gulf today for it makes up most of the sub-marine ridges and the channels of today, although not as prolific as those of old, are usually in close proximity to sub-surface relict channels. Within the Port Augusta region, sequence V is extensive and is a major factor in the production of today's topography. However, it is not just sequence V for sequence IV changes the topography of sequence V to a topography reminiscent of today's. Certainly sequence V has had its effects but it is the topography of sequence IV, or better, what is left of sequence IV that is the producer of today's topography.

SEQUENCE III WHYALLA NORTH

Sequence III within this region is very similar to sequence II in that it is predominantly intertidal and sabkha deposits with minor aeolian and subtidal deposits. Simply, shelly, sandy silts and silts with dune quartz. Descriptions are identical to that of sequence II but greater quartz exists than in sequence II. Sequence III within this region can be up to 15 m. thick but the mean thickness is 3 to 4 m.

WHYALLA CENTRAL

Sequence III in this region is identical to the above region and hence to sequence II except for the fact that often at the top of the sequence oxidation has taken place to varying degrees giving a yellow to brown colour to the sediment. Thickness is on average 3 to 4 m.

Again, as for sequence II,

sequence III is basically only found in the troughs but also it makes up some of the lower terraces of the Gulf floor.

PORT AUGUSTA

No core was emplaced into sequence III sediment within this region but again it is found at the base of troughs so it is probably subtidal, intertidal and/or a sabkha deposit. Sequence III in fact is only observed in the deepest of the trenches and only up to a thickness of 2 to 3 m. within restricted areas, because there are only a few deep trenches within the region. Sequence III, hence, is not as prolific here as it is in other regions.

SEQUENCE II WHYALLA NORTH

Sequence II consists of basically intertidal and supratidal deposits, in particular sabkhas. Again, simply, shelly, sandy silts and silts. The intertidal deposits are usually oxidized to a faint yellow colour. Quartz and the various faunas are present. Quartz proportions average out at 20% and bivalves predominate mainly as fragments within the sediment. Benthic foraminifera, turreted gastropods, aragonite needles, coralline algae, echinoid spines and other faunas are also observed. The quartz is predominantly sub-angular to subrounded, polished with some frosted, and poorly sorted. When the quartz percentage increases above 20% it usually steps up to 80% and greater. Here it becomes difficult to decide whether a dune or a beach is being observed.

The sabkha deposits of sequence II are white to green in colour. Usually less than 5% quartz and less than 5% shell fragments are noted within the sabkha. Where pockets of greater shell or quartz are noted it is thought that these pockets are but storm-water deposits on the sabkha surface. Often gypsum crystals are found.

WHYALLA CENTRAL

Sequence II in this region is basically a shelly silt, the question being whether it is

of intertidal and/or subtidal origin. A white to yellow calcareous silt predominates, usually averaging greater than 50% of the deposit. Shelly hash on occasions is more predominant. Quartz proportions at their greatest approximate 10%. Whole turreted gastropods and rarely whole bivalves are found. Benthic foraminifera, echinoid spines, aragonite needles, etc., are observed but again bivalve fragments predominate the shelly fraction of the sediment. The subtidal and/or intertidal deposits are up to 4 m. thick. Sabkha deposits are described as above. Clay-dune material of very fine silt is also observed.

Within the Whyalla region, overall, intertidal and supratidal (sabkha) deposits are the most prominent, with some subtidal deposits within what were then deeps.

PORT AUGUSTA

Sequence II within this region is non-existent for reasons that will be discussed later.

HOLOCENE SEQUENCE I (A and B) WHYALLA NORTH AND CENTRAL

Core descriptions indicate that sequence I consists of a shelly sand, basically of subtidal origin. Usually the percentage of silt increases downwards. The distinguishing factor is that sequence I is always grey. Molluscs are the major fauna with smaller proportions of benthic foraminifera, aragonite needles, echinoid spines and plates, coralline algae, etc.. Seagrass fibre is quite common within the sediment. Grey staining on shell is present quite often and is best seen on the foraminifera, echinoid spines and bivalve fragments. Quartz is always present. Silt is either brown or grey. The thickness of sequence I is variable but the mean thickness is 1 m.

PORT AUGUSTA

Sequence I in the Port Augusta region is of two types. Sequence IA is identical to sequence I found in both Whyalla North and Central regions.

Sequence IB consists of approximately 70 to 80% sub-rounded, polished and poorly-sorted quartz with little silt or shell. The shell is highly weathered and an orange iron-oxide colour, as is the small percentage of silt. The quartz is also iron-oxide-stained. It would appear that this material, sequence IB, is purely reworked sediment of an older age, Pleistocene age, as will be explained.

SEQUENCE VI

Sequence VI is of various ages being located between sequence I and IV and IV and V. It consists of alluvial material, a deep red/brown silt with polished, rounded to sub-rounded, poorly-sorted quartz. Alluvial material when observed between sequences I and IV is only seen as a small proportion of dark-brown clay with minor quartz. It is thought that a major proportion of the alluvial deposit has been reworked into sequence I. How-ever, the tell-tale factor in recognizing alluvial material is the non-calcareous nature of the sediment. Care must be taken, however, as it is possible for decalcification of the sediment to take place making an intertidal, or any, deposit non-calcareous.

In summary, the major differences between the sequences are that sequence I is grey and subtidal. Sequence II and III are identical basically consisting of intertidal and supratidal (sabkha) deposits with minor aeolian and subtidal deposits. Sequence IV is red and has coarse intertidal deposits and has subtidal and estuarine deposits. Sequence V is little known and also is red but is probably identical to sequence IV. Sequence VI is non-calcareous alluvial material.

4. QUATERNARY SEDIMENT DESCRIPTION HOLOCENE

The Recent sediment within the Gulf is, on average, 1 m. thick. Typically it is grey in colour. The flora and fauna within this sediment are described in sections 2.23 and 2.24, respectively. Often the Holocene sediments may be divided into two parts.

Firstly, the top 10 to 30 cm. often is coarser than the underlying sediment. Usually there is a greater proportion of coarse shell and a minor percentage of quartz. The coarse shell consists of whole or fragmentary bivalves and gastropods. Mostly the shell is fragmented and weathered. On occasions coarse coralline algae is prominent and is so prolific that a speckled effect results with white dots (coralline algae) on a grey background. Foraminifera, echinoid spines, bryozoans, aragonite needles and coralline algae amount to only minor fractions of the sediment. Sea-grass fibre may be observed. The proportion of shell usually ranges from 60 to 90%. Quartz is only found in minor proportions, 10 to 20%, in the Recent sediment, except in the Port Augusta Region as mentioned before. The quartz is poorly sorted, angular to rounded, polished and frosted, clear or green to red in colour and of a very fine sand grain-size. Brown or grey silt makes up the rest of the deposit.

The second, basal, section of the Holocene is finer in character with much greater grey silt and organic matter than the top section. Shell is less prolific and is comminuted. Certainly, bivalve and gastropod fragments are still observed but there is an increase in proportions of aragonite needles and organic matter. Foraminifera may occur in greater numbers as well. The silt proportion may build up to 70%. Iess quartz is observed in this section.

This sequence of sediments described above is observed in most cores, but in some areas only one of the two may be found. Where the Holocene sequence is thin it is noted that only the coarser deposit is

present. In other parts of the Gulf only the finer sequence can be found. However, often a coarsening upwards in grainsize is seen within these sequences, whether there be only one or both types.

Another noticeable factor is that the faunal assemblage within the Holocene often changes from the basal section through to the top. Core 83 (see APPENDIX 2A) is a good example. The basic faunas are bivalves, turreted gastropods and foraminifera of the genera Quinqueloculina and Spiroloculina. Aragonite needles are also common. These biota are found in the top 10 cm. of the Holocene. In the basal 10 cm. a greater prominence of foraminifera is noted. Foraminifera of the genera Elphidium, Peneroplis, Spirolina, Quinqueloculina and Spiroloculina are observed. The basal section within this core also has more bivalves, turreted and non-turreted gastropods. This change in faunal assemblage must mean there has been a change in conditions in the Gulf over the Holocene period. out what these changes were, further work is required.

Often, also, the shell, in particular bivalves, foraminifera and echinoid spines, have a grey to black stain on their surfaces. It is possible that this staining is caused by in situ reduction within the sediment.

PIEISTOCENE

Within the Pleistocene several facies are present. These facies are aeolian, supratidal flat, sabkha, intertidal flat, beach, subtidal and estuarine. The factors used to discriminate between these facies have been mentioned before. Typical examples of these facies are listed below.

Aeolian: Greater than 80% quartz. The quartz is well sorted, sub-rounded, polished and frosted, and stained an iron-oxide red. No shell is apparent. Red silt makes up the rest. Core 225.

Supratidal flat: 2% fragmentary shell consisting of bivalves, turreted gastropods and foraminifera. 30% moderately-sorted, sub-angular to sub-rounded,

polished, fine-grained quartz. 68% pale-green calcareous clay Core 233.

Sabkha: Calcareous white clay Core 225. Gypsum is often found.

Intertidal flat: 40% shell including aragonite need-les, bivalves, and turreted gastropods, both whole and fragmentary. 20% moderately-sorted, sub-rounded to rounded, polished quartz. 40% nodular white calcareous silt Core 227.

Beach: Quartz and shell. The shell consists of foraminifera, echinoid spines, turreted gastropods and bivalves. The quartz is moderately sorted, subangular to sub-rounded, polished and iron-oxide stained. The deposit is pale yellow in colour to pink at the base Core 231. A beach deposit may be all shell or all quartz or both but there is little silt. Subtidal: 70% shell consisting of foraminifera, bivalves, non-turreted gastropods and echinoid spines. Whole bivalves and gastropods are observed. Only a small proportion of quartz is noted. The rest is creamy-white silt Core 83.

Estuarine: Grey, calcareous, fine quartz sand-rich clay. One Anadara trapezia noted Core 162.

The above are only examples and the percentages of quartz, shell and silt vary for each facies. The Pleistocene sediments are of various facies but the discriminating factor between them and the Holocene sediment is colour and organic matter. The Holocene is always grey and always has some organic matter while the Pleistocene has no organic matter and is never grey except where estuarine deposits are noted, but there is a great difference in the physical character of the Holocene and the Pleistocene estuarine deposit, the Pleistocene being harder. The Pleistocene ranges from white, cream, buff, khaki, green, brown to red in colour. Pedogenic nodules and burrows often are observed in the Pleistocene in the study area.

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The fauna of both the Holocene and Pleistocene are very similar and there are few species seen in the Pleistocene and not in the Holocene.

Reduction is common in the Holocene but oxidation is more common in the Pleistocene, and this oxidation is apparent in the yellow-to-red marine deposits that now are noted as paleosols.

A major factor that has contributed to the Pleistocene sediment is the level of the sea during the formation of these specific facies, whereas the Holocene sediment seen in the submarine cores is entirely subtidal estuarine sea-grass bank facies. Further information on the sediments can be obtained from APPENDIX 2A.

5. DATING

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Amino acid racemization

and thermoluminesence dating of the pre-Holocene sequences is in the initial stages and so far the dates have proved inconclusive, only indicating ages of approximately 100,000 years plus or minus 20,000 years. Therefore other methods must be found.

Over the years, more and more work has been focused on sea levels throughout time, especially the relatively recent changes in sea level. Within the Southern hemisphere work by CHAPPELL, J. (1974) and BLOOM et al. (1974) and CHAPPELL, J. and VEEH (1978) on sea levels is extensive. It is known that sequence I is of Holocene age but sequences II to V are of unknown ages. Each of these sequences represents a transgressive unit of sediment, hence if the transgressive sea level periods are known and dated then surely approximate dates can be associated with each sequence.

CHAPPELL and BLOOM et al.

(1974) have focused their attention on the Huon Peninsula in New Guinea. Emerged coral reefs form terraces along the coastline. Interpretation of sea level oscillations superimposed on a rising land rests on a principle of reef morphology identified by DARWIN (1842, pp.76-78) viz., a barrier and a lagoon association implies a sea-level rise relative to the land. Contrarywise, a fringing reef reflects relatively stable juxtaposition of land and sea, and descending flights of reef terraces, their fronts cut by low cliffs and notches, indicate emergence. Using this information along with C^{14} and Th^{230}/U^{234} dating methods, the terraces on the Huon Peninsula were studied. ments were subtracted using methods explained by CHAPPELL, J. (1974). Results to date (1974) indicate sea level maxima at about the following times: 30,000, 40,000 to 50,000, 60,000, 80,000, 120,000, 140,000, 185,000, 220,000 in addition to the recent maxima. Further dating in 1974 confirmed these ages and found an additional high sea level at 105,000 years. These dates correspond with dated upper Quaternary terraces elsewhere, especially Barbados, and correspond to EMILIANI'S (1966) paleotemperature curve.

Along with these dates the sea levels were calculated and the two plotted against each other. FIG.15 illustrates the final product.

and VEEH (1978) has changed the heights of the sea level highs but not the ages. The 80,000, 100,000, and 120,000 transgressive sea levels reached maximums of -20 m., -14 m., and 6.5 m., respectively in comparison to today's sea levels. These sea levels were identified from work on Atauro Island and Timor and correlate well with work from New Guinea and Barbados. However, within the Gulf it is thought that the maximum sea levels for these periods are slightly greater, -15 m. and -8 m. for the 80,000- and 100,000-year transgressions, respectively (pers.comm., T. BELPERIO, 1981). No definite levels have been calculated, but certainly within the Gulf the maximum sea levels for the transgressive periods are greater than those levels calculated by Chappell and Veeh.

Within the Gulf, sequences II to V are basically transgressive sequences. The sequences, as mentioned, are identified by using both seismic reflection profiles and core interpretation. The breaks between the sequences are recognizable. These breaks, representing the maximum heights of the sequences, are at levels that correspond to the maximum sea levels of the various transgressive seas. For example, sequence II is only found where the water depth is now greater than 15 m.. approximately. Sequence III is only found in areas where the water depth is now greater than 8 m., approximately. Sequences IV and V are prolific and are observed in any depth of water. Hence it is suggested that sequence II is the equivalent of the 80,000-year-old transgressive event, and sequence III, IV and V also are equivalent transgressive deposits to the coral reefs of New Guinea. Therefore, an age of 80,000 years is suggested for sequence II, 100,000 years for sequence III, 120,000 years for sequence IV and an older age, possibly 220,000 years, for sequence V.

The distribution of sequences I to V within the Gulf, then, is controlled by transgressive

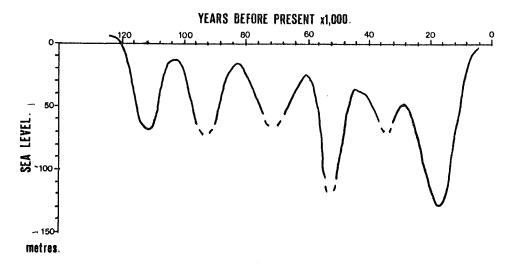


FIGURE 15.

EUSTATIC CHANGES.
CHAPPELL and BLOOM et.al. (1974),

seas and because of the shallow nature of the Gulf it is only the pronounced sea-level rises that had a depositional effect. It would be true to say that the two regions studied have been more often exposed to the atmosphere than covered by water.

6. ANALYSES OF THE HOLOCENE SEDIMENT

A sample of 5 to 10 grams was randomly chosen from the top 2 cm. of Holocene sediment within the core, for each core. The sand fraction was then separated from the whole fraction by a wet-sieving procedure. (See APPENDICES I.C. and 2.B.) The sand fraction was then studied under a binocular microscope.

To minimize any bias the sand fraction was thoroughly mixed and then divided into quarters. One quarter was randomly selected and then a count of 500 particles was carried out. Counts of 1000 were taken and the figures were identical to that of a count of 500. Counts less than 500 disclosed a marked divergence in proportions to those calculated using counts of 500 or 1000. Hence, a count of 500 was used, being the smallest count possible which gave reproducible figures.

The series of non-organic and organic particles were divided into the following types:

Coralline algae	Foraminifera -	genus	Spiroloculina and
Bryozoa			<u>Quinqueloculina</u>
Gastropods		genus	<u>Triloculina</u>
Bivalves		genus	Elphidium
Quartz		genus	Peneroplis and
Aragonite needles	•		<u>Spirolina</u>
Echinoid spines		genus	<u>Discorbis</u> and
and plates			<u>Rosalina</u>
Miscellaneous		genus	<u>Nubecularia</u>

The miscellaneous division included material that was of a very fine sand grain-size and its origin could not be recognized. It is believed that most of the material within the miscellaneous division is of bivalve, gastropod and foraminiferal origin. Also organic or non-organic matter not in the above divisions, for example, rock fragments, were put into this division.

Results from the counts

are found in APPENDIX 2.D.

tages for each division were calculated and plotted for each sample. Transparencies depicting the percentages of each division were used to perceive relationships, if any, existing between the various organisms and non-organic matter. The calcium carbonate percentage per sample was also calculated. See APPENDICES I.D and 2.C.

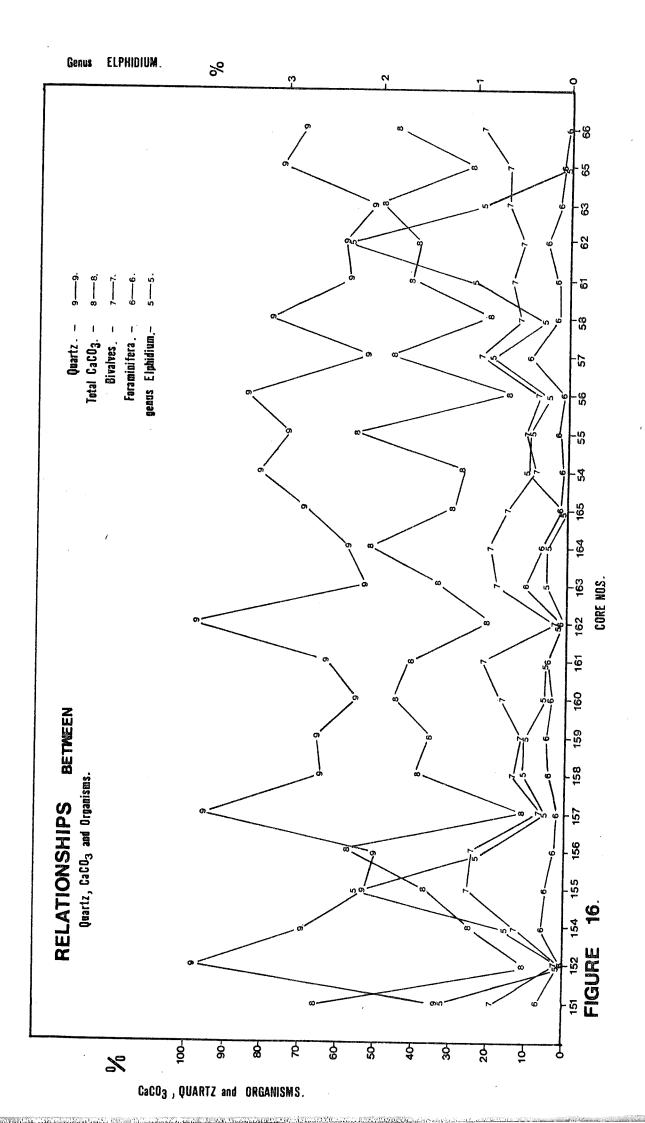
Distinct reciprocal relationships exist between the proportion of quartz and the proportion of bivalves and also quartz and CaCO3. FIG.16. That is, the greater the proportion of quartz, the less the proportion of bivalves and CaCO3 and vice versa. Between bivalves and CaCO3 a positive relationship exists. This reciprocal relationship is expected, for the greater the terrestrial influence, the greater the quartz; the less the marine influence, the less the marine fauna, e.g., bivalves. Bivalves are so prominent within the Gulf, within the faunal assemblage, that the CaCO3 curve and the bivalve proportion of the whole fauna, when graphed, are very close to identical. Hence the relationship being examined here is one depicting the proportion of terrestrial and/or marine influence.

However, a reciprocal relationship exists between the proportion of foraminifera, in particular the genus Elphidium, and quartz and a mutual relationship exists between bivalves, CaCO₃ and foraminifera. The proportions are not identical but the pattern of distribution between bivalves and foraminifera is identical. Where the proportion of bivalves increases so does the proportion of foraminifera, etc.. See FIG.16.

This would indicate that the degree of marine influence is marked by not only bivalves but by foraminifera. It would seem that the conditions conducive to bivalve numbers increasing are the conditions that benthic foraminifera also prefer, especially the genus Elphidium.

No other relationships

can be observed.

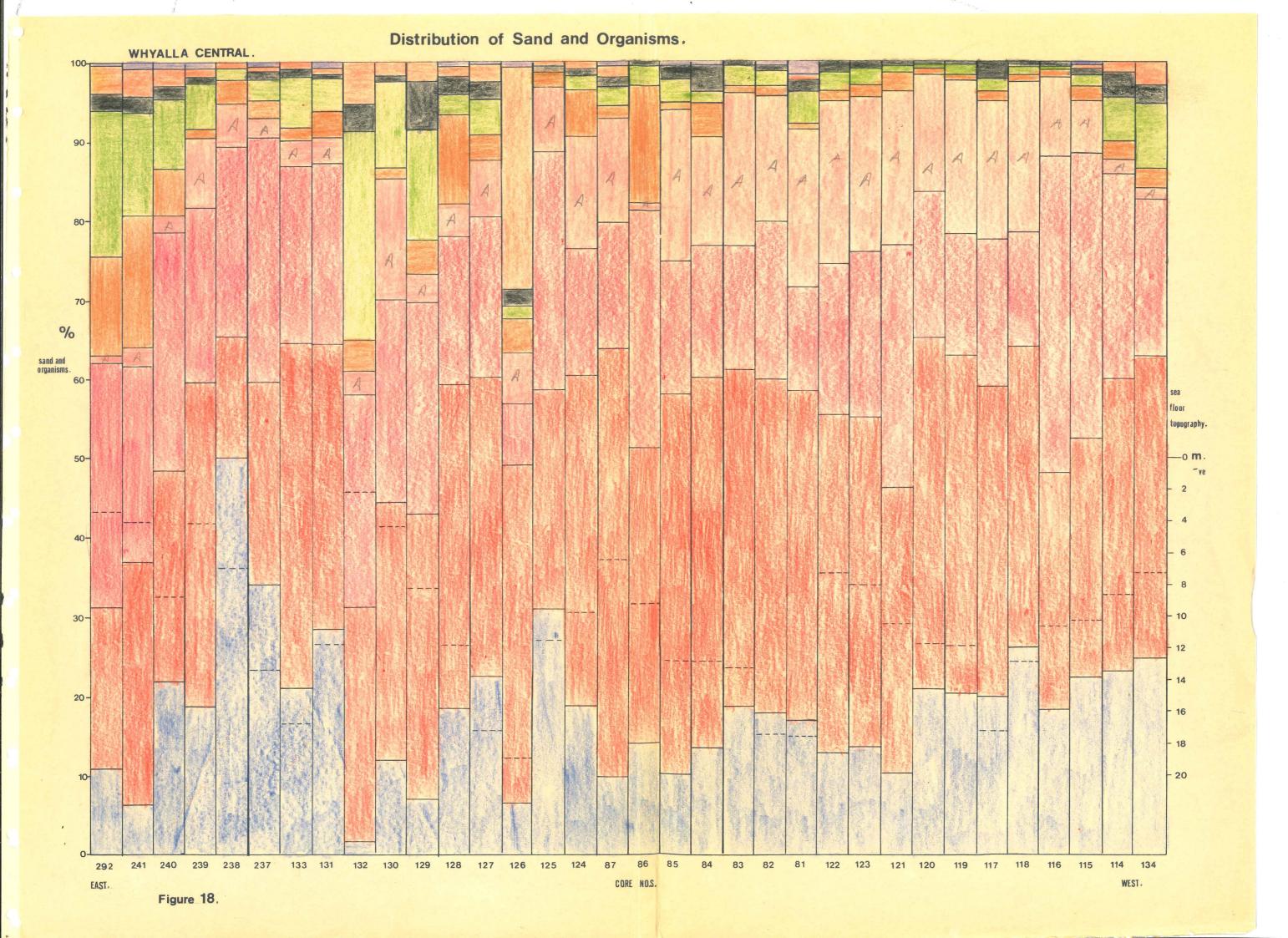


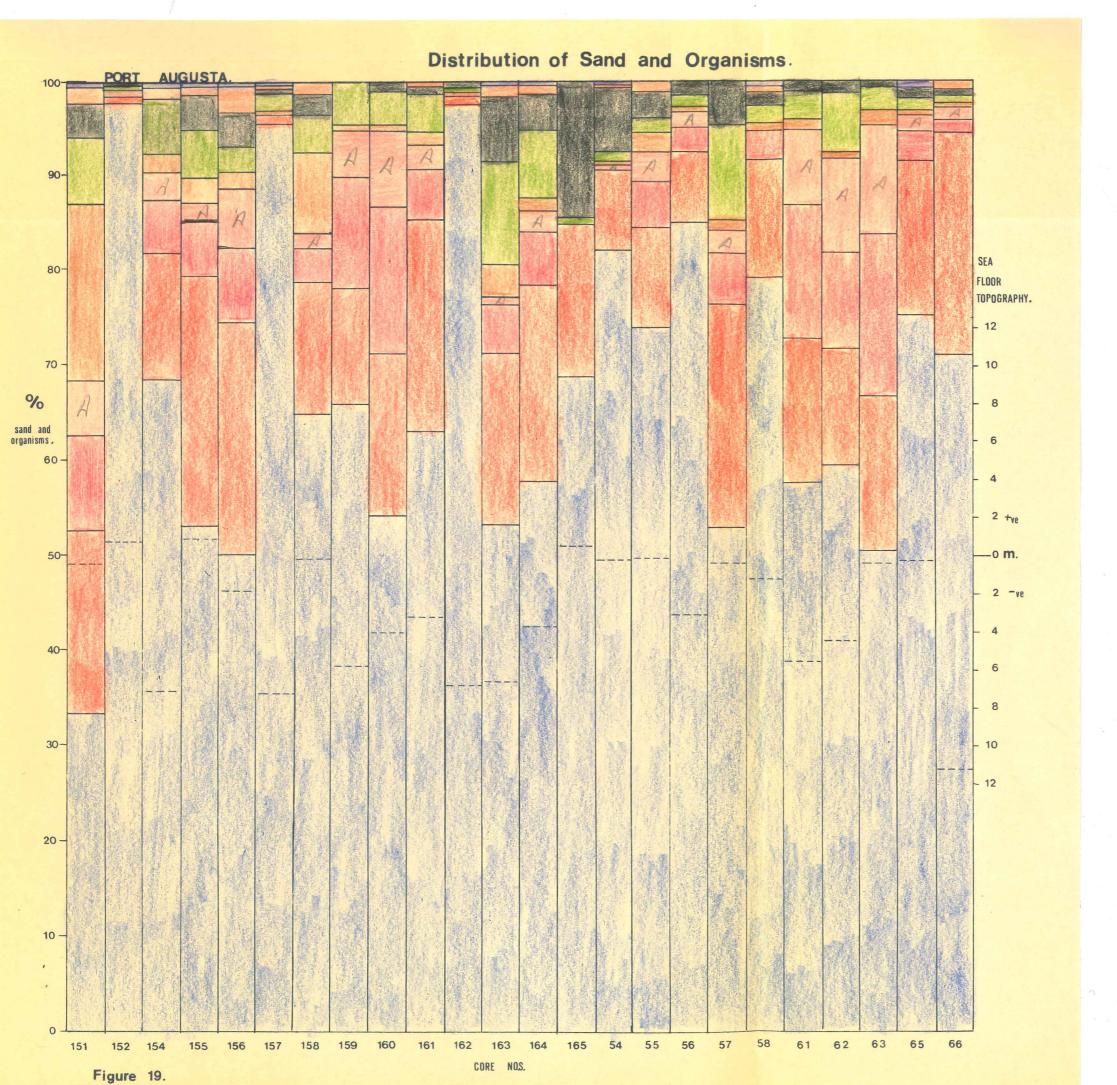
Bar graphs were used to show the proportion of organisms and non-organic matter within the sample for each core. Each column represents a core. The various foraminiferal genera were summated and the total figure used. A separate graph was used to show the proportions of various foraminiferal genera within the samples. The variation from one core to another and across all the cores discloses the distribution of the species or genera through the Gulf. FIGURES 17, 18, and 19 show this distribution for Whyalla North, Whyalla Central and Port Augusta regions, respectively.

There are definite variations between the northern part of the Gulf and the southern section. From the diagrams it is quite clear that in the Port Augusta region (FIG.19), a greater terrigenous component, quartz, exists, while in the Whyalla region, FIGS.17 and 18, bivalves are the major component within the sand fraction. In fact, in the Port Augusta region, quartz makes up to 95% of the sediment, averaging 65% of the sediment. In the Whyalla region, quartz averages approximately 20% of the sediment and with bivalves summates to an average of 65%. Hence it can be stated that quartz is extremely dominant in the northern part of the Gulf, having a greater dominance than bivalves do in the south.

In the northern region, FIG.19, there are fewer bivalves, echinoid spines, coralline algae, gastropods, bryozoans, and in particular aragonite needles in comparison with the southern region. FIGS.17 and 18.

Also, along the Whyalla Central transect, FIG.18, a definite variation in Holocene sediment can be seen. Looking at the diagram, it is quite obvious that from core 85 to 134 a different sediment-type exists to that of the east. The major change is in the aragonite-needle proportions which have at least doubled and which show a consistent proportion within the western part of the Gulf. Also, a decrease in the proportions of coralline algae, bryozoans, gastropods, echinoid spines and foraminifera exists in this regions compared to the eastern





region.

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The reason for the significant difference between one side of the Gulf and the other is unknown. Salinity and temperature are the vital factors which usually govern the distribution of organisms, but as no detailed information is available, such conclusions await further research.

However, it is noticeable that the Port Augusta region is shallower than the Whyalla region. FIGS.17, 18, and 19. Also, a large proportion of the sediment designated Holocene in the Port Augusta region is definitely reworked from Pleistocene sediment. From microscope studies it was observed that some foraminifera, bivalves and echinoid spines are an orange, iron-oxide stained, colour. Also, most of the quartz in each sample has an orange-to-red stain. On occasions, the fauna mentioned are half orange, half black, demonstrating a history of oxidation and reduction.

Now, if a transgression is rapid, then hardly any reworking will take place; but if the sea level rises very slowly, then there will be greater reworking. Hence, one can safely say that, with the reworking present in the northern part of the Gulf, the last transgression must have been a slow one.

The reason for this difference between the two regions studied, Port Augusta and Whyalla, may be greater runoff in the northern region, hence greater quartz in the north. With greater terrestrial influence in the north, less fauna will be observed and greater quartz and silt will be noted. The greater reworking in the north also influences the differences in the overall sediment. It would seem that at first the transgression must have been rapid up to a depth that included the Whyalla Central and North region, and then it must have slowed down so that sea level only gradually increased in the northern part of the Gulf. Hence the greater reworking in the northern Gulf and less in the southern Gulf.

The distribution of the foraminifera is displayed using bar graphs. From the counts, foraminifera were split up into divisions based on genera. On this occasion, the sum of foraminifera is 100%. The percentage of each division of genera is graphed. The cores again are represented by a column on the graph. FIGURES 20, 21, and 22 show the foraminiferal distribution for Whyalla North, Whyalla Central and Port Augusta, respectively.

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Again it was observed that the foraminifera differed from one region to another. In the Port Augusta region, FIG.22, there is a greater preponderance of <u>Peneroplis planatus</u> and <u>Spirolina aciculoris</u> than in the Whyalla region, FIGS.20, 21, where they are almost non-existent.

Also, it was noticed that the foraminiferal assemblages varied with the topography. FIG.21 exemplifies this very well. When the topography of the sea-bed is transformed on to the foraminifera distribution, it is obvious that the banks have a different foraminiferal assemblage to the troughs. Foraminifera of the genera Elphidium, Discorbis and Nubeculoria are the prominent foraminifera on the banks while foraminifera of the genera Quinqueloculina, Spiroloculina and Triloculina are the prominent genera within the troughs. It has been suggested that the distribution of the foraminifera is associated with the distribution of sea-grass. However, while the distribution, which does vary, does not correlate with the foraminiferal change.

From the analyses it would seem plausible that water depth is an important factor in the distribution of foraminifera for there is a distinct variation in foraminiferal assemblage from troughs to banks and from the deeper waters of the Whyalla region to the shallow waters of the Port Augusta region.

The question must be raised as to whether the foraminiferal assemblage being

LEGEND FOR THE FORAMINIFERAL DISTRIBUTION

FIGURES: 20, 21 and 22

Foraminifera of the genera:

Quinqueloculina and Spiroloculina

Triloculina

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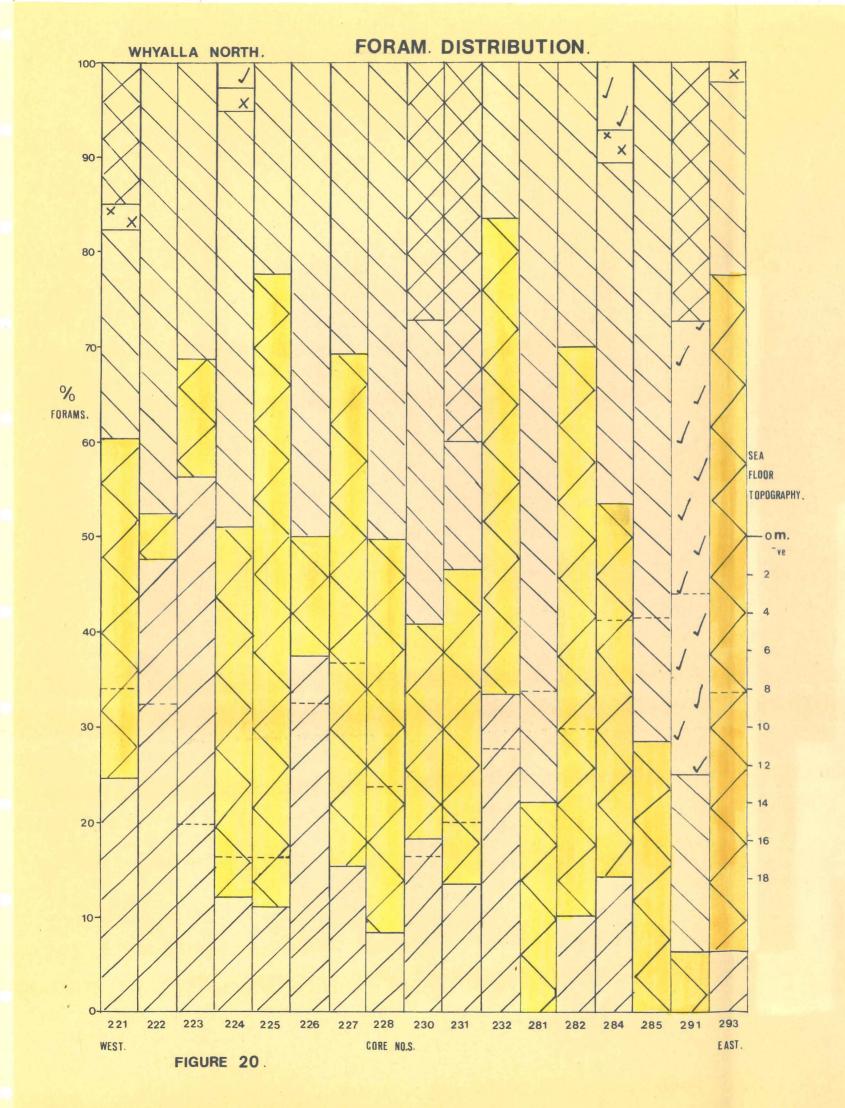
Elphidium

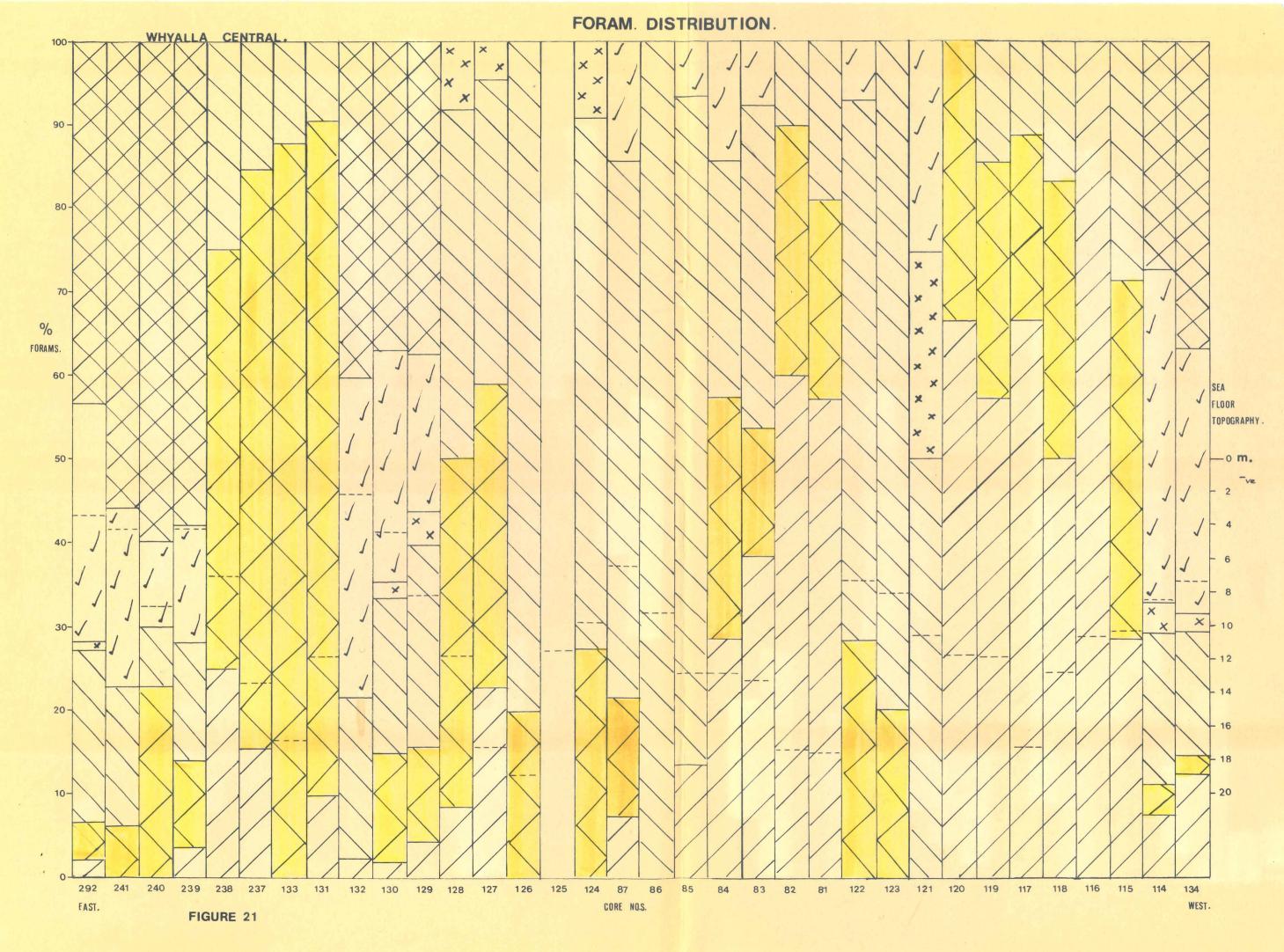
× Peneroplis and Spirolina

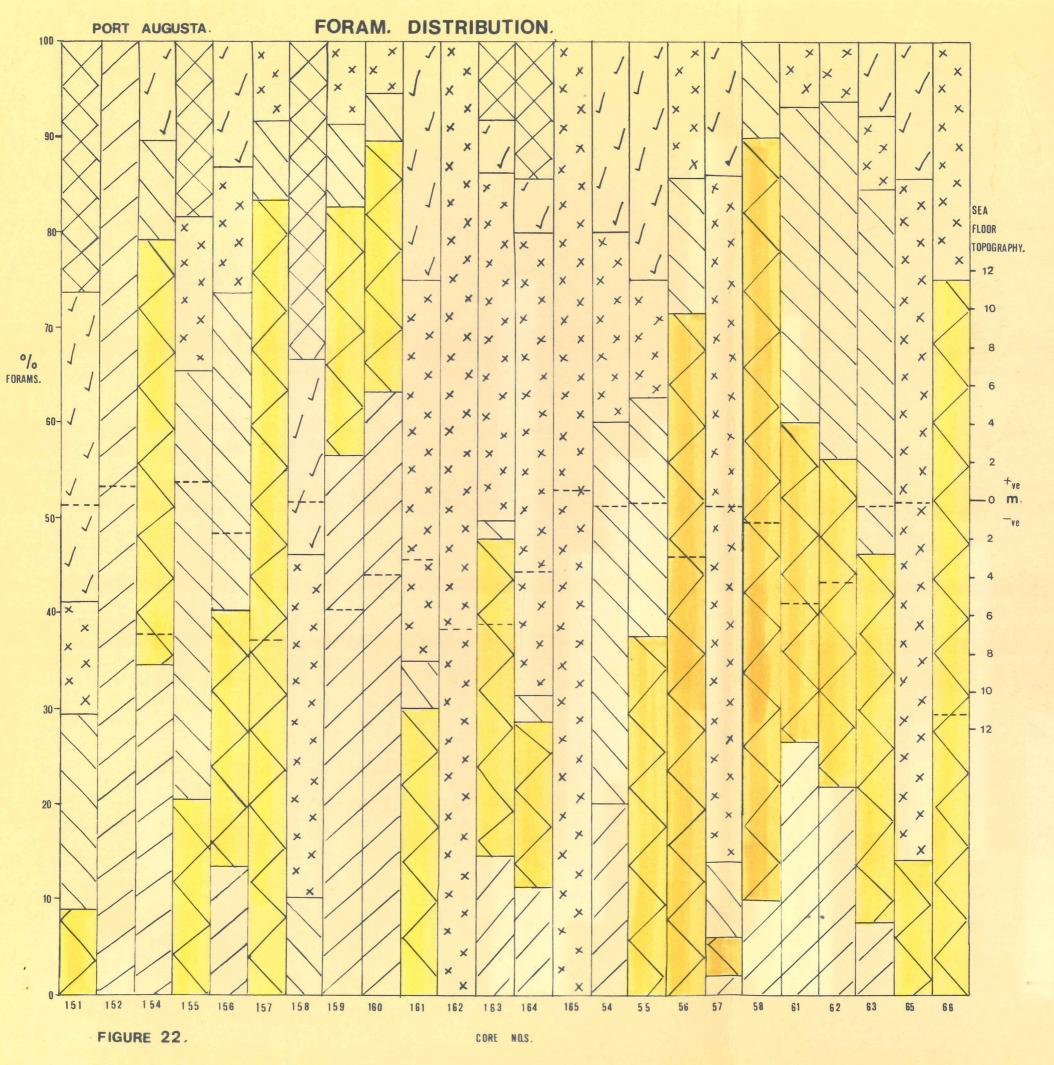
Discorbis and Rosalina

Nubecularia

Line representing sea-floor topography







studied is the life assemblage or an assemblage of foraminifera thrown together through differential sorting upon transport. To some degree it must be the life assemblage as most of the forms are fresh with no staining, pitting or scouring. Certainly forms with staining, pitting, etc., are present but the fresh forms are more numerous than the fragmentary or weathered forms, especially in the Whyalla region. It is possible that the differential distribution of foraminifera with water depth is actually caused by differential shape-sorting upon transport of the foraminifera within the Gulf. However, as many of the forms are fresh it would seem likely that little transport has taken place and that the foraminifera are in life-assemblage position. It is only in the Port Augusta region where the distribution of foraminifera is biased by the much greater degree of reworking.

It has already been cited, within this paper, that the whole faunal assemblage, to a degree, changes from one region to another. Looking at the depths of water within the two regions, it is obvious that the Port Augusta region is very much shallower, and possibly it is this difference in water depths that is behind at least part of the difference in fauna from one region to another, the difference being in species and genera proportions of fauna, in particular foraminifera.

However, in the study, depth is the only physical parameter known. According to MURRAY, J.W. (1973), it is most probable that depth is not the limiting factor, but perhaps one of the many factors related to depth: hydrostatic pressure, density, light penetration, temperature, pH, oxygen content and carbon di-oxide content. Although throughout the literature there are correlations between depth and the distribution of fauna in restricted geographical areas, on a world-wide scale individual species do not appear to be controlled by depth alone.

Within the Gulf, therefore, there are variables associated with depth which govern the distribution of fauna, mainly foraminifera. Hence these variables associated with depth vary from Port Augusta to Whyalla.

7. CONCLUSION

Using the seismic profiles and core data from the Gulf, six sequences have been iden-The oldest, sequence V, is only noted in the seismic record, and only to a minor extent in the core data, but is probably similar to sequence IV. Sequence IV is red in colour and consists of supratidal, intertidal, subtidal and estuarine deposits. The intertidal deposits are coarser than those found in the other sequences. Sequences II and III are identical in that both are basically of sabkha and intertidal facies, with minor subtidal deposits. quences V and IV are prolific with III, II and I less so, and II is certainly the least abundant. Sequence IA is grey and of subtidal origin. Sequence IB, only found in the Port Augusta region, is reworked Pleistocene sediment usually consisting of mainly quartz and iron-oxide-stained shell and silt. Sequence VI is non-calcareous, brown, alluvial sediment. Using CHAPPELL (1974) and BLOOM'S et al. (1974), and CHAPPELL and VEEH'S (1978) work, dates of 80,000, 100,000, and 120,000 are suggested for sequences II to IV respectively. It is suspected that sequence V is representative of the 220,000-year transgression. I is Recent in age, and sequence VI, alluvium of various ages. All of the sequences, except for VI, are transgressive sequences and are dated using this fact.

A study of the Holocene sediment reveals that statistical relationships existing within the fauna and non-organic matter are few. A reciprocal relationship exists between the quartz proportion and the bivalve or foraminifera or CaCO₃ proportion. This relationship probably depicts the degree of marine CaCO₃ or terrestrial Quartz influence. Both foraminifera, especially the genus Elphidium, and bivalves have a distribution which correlates well with the overall CaCO₃ distribution within the Gulf. Hence either of these forms can be used to give an indication of CaCO₃ content or marine influence.

Further work carried out on the Holocene sediment showed that the Port Augusta region sediment is dominantly quartz with minor fauna, while the

Whyalla region has greater fauna, in particular bivalves. The reason for the greater quartz content in the north is thought to be greater runoff into the area. Also, greater reworking of the sediment within this region could increase the quartz content with the fines being transported away. It would seem also that, with the smaller percentage of fauna, that the Port Augusta region is not as favourable for faunal life as is the Whyalla region. The greater reworking in this area was probably caused by a slow rise in sea level, yet this reworking is not seen in the Whyalla region. Hence, sea level must have risen at a rapid rate to a point above Whyalla and then slowed down to cause the greater reworking in the Port Augusta region.

Finally, the foraminiferal distribution was studied. The Port Augusta region, with shallower water, has a higher proportion of Peneroplis planatus, while the Whyalla region has few. Within the Whyalla region two assemblages of foraminifera are noted. One assemblage consisting of the genera Elphidium, Discorbis and Nubecularia is prominent on the shoals, while an assemblage of the genera Quinqueloculina, Spiroloculina and Triloculina is prominent in the deeps. This distribution can be related to water depth, but it is thought that one or more of the variables associated with water depth controls the foraminiferal distribution.

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Finally, thanks to Mrs. Sangster for the typing of this thesis.

	PLATE 1		Magnification
1.	Gastropods	Zeacumantus sp.	in triplate?
2.		Vexillum plicarium	5000 10 x
3.		Truncatella vincent	lano 40 x
4.		Argaliste fugitiva	50 x
5.	Aragonite needle		70 x
6.	Aragonite needle	s	10 x
7.	Aragonite needle	S	20 x
8.	Bryozoan		10 x
9.	Bryozoan		10 x
10.	Echinoid spine		60 x
11.	Echinoid spine		30 x
12.	Echinoid plate		70 x
13.	Coralline algae		lO x
14.	Coralline algae		40 х
15.	Coralline algae		20 x

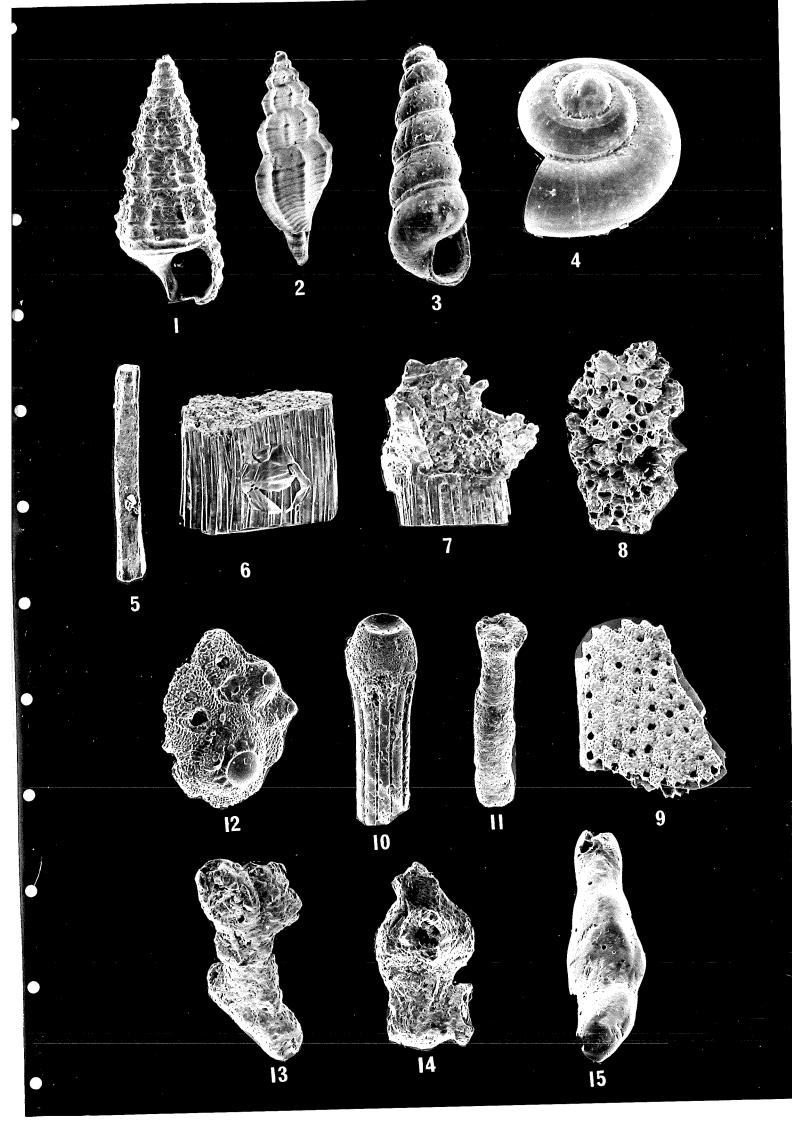
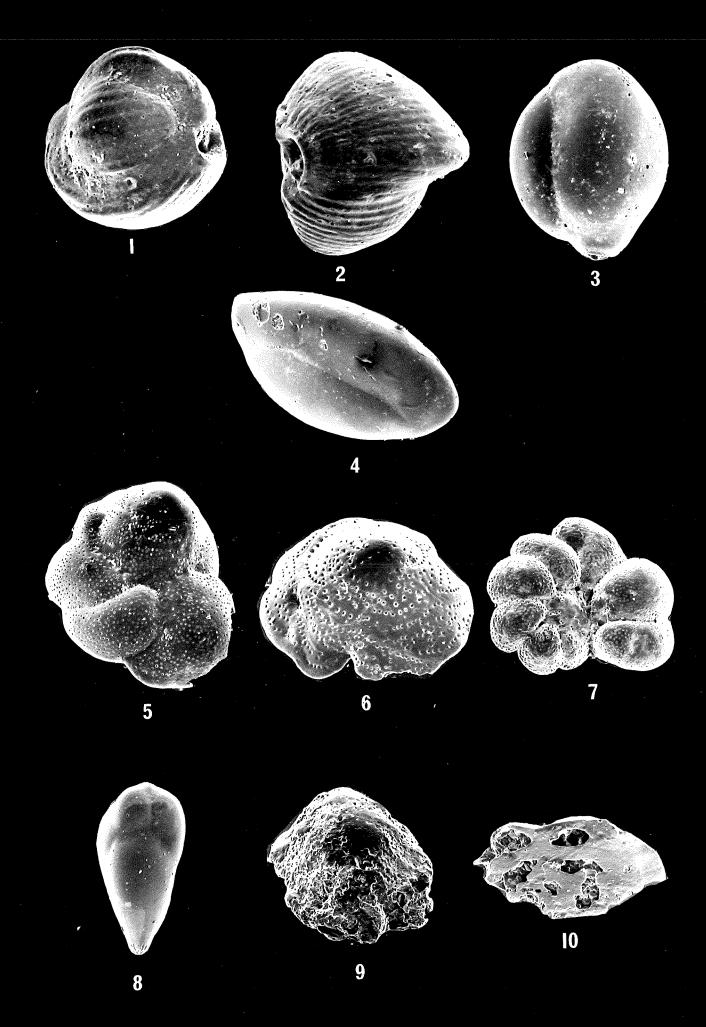
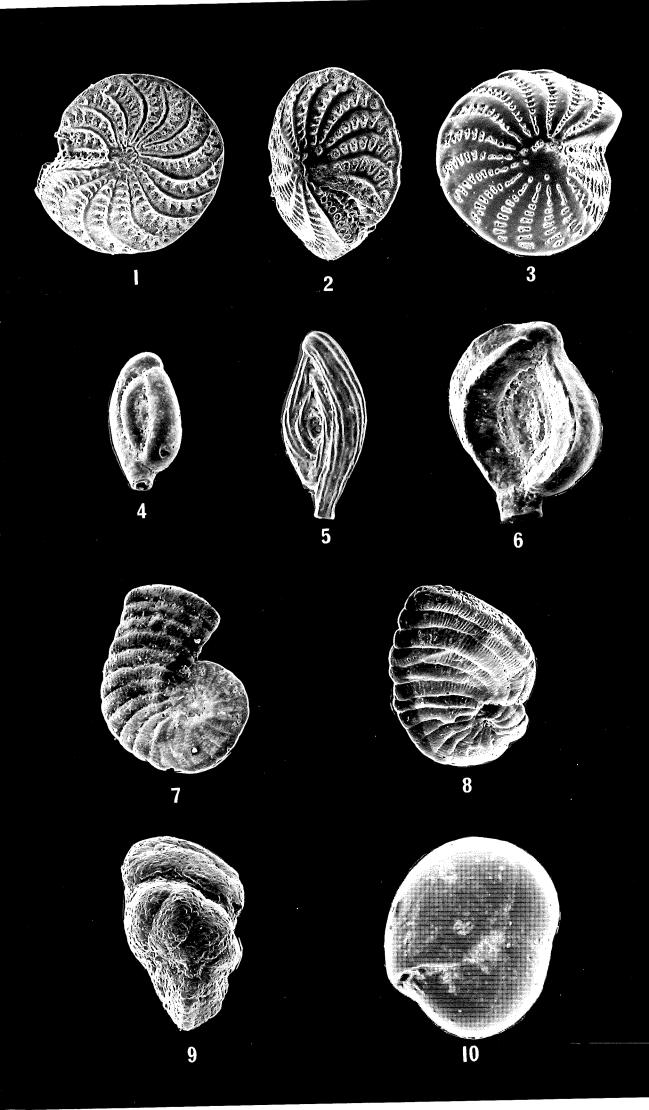


PLATE 2 Magnification Foraminifera 1. Triloculina striatotrigonula 100 x Triloculina striatotrigonula 2. x 08 3. Triloculina trigonula 50 x Triloculina oblonga 70 x 5. Discorbis vesicularis 50 x 6. Rosalina kennedyi 60 x 7. Discorbis dimiadatus 40 x 8. Guttulina pacifica 60 x 9. Nubecularia lucifuga 40 x 10. Nubecularia lucifuga 30 x



		<u>P</u> L	<u>ATE 3</u>	<u>Magnification</u>
	Foraminifera	1.	Elphidium macellum	80 x
		2.	Elphidium macellum	70 x
1		3.	Elphidium pseudonodosum	80 x
:		4.	Quinqueloculina tropicalis	. 50 x
		5.	Spiroloculina angusteorali	<u>s</u> 50 x
		6.	Spiroloculina communis	50 x
		7.	Spirolina aciculoris	40 x
		8.	Peneroplis planatus	30 x
		9.	Cribrobulima polystoma	40 x
		10.	Miliolinella subrotunda	110 x

(



APPENDIX 1:

METHODS OF STUDY

- 1.A. Geophysical equipment and technique
- 1.B. Vibrocoring and core preparation
- 1.C. Size analysis
- 1.D. CaCO₃ analysis

APPENDIX 1.A.

GEOPHYSICAL EQUIPMENT AND TECHNIQUE

- 1) An E.P.C. 4100 seismic reflection profiling recorder.
- 2) A MAG 30V. transducer with baffle.
- 3) A SR555 transmitter on 185 joules setting, firing 2 per second.
- 4) A recorder set positive phase, 1 of 4 on 125 ms. sweep speed; gain 100.
- A channel splitter SRCS Mk2 Channel A positive phase 132 gain and Channel B rectified signal 55 gain (filters Ch.A 240 H_z; Ch.B 430 H_z).
- 6) A hydrophone 20 element 45 OB pre-amplifier; 300 $\rm H_{Z}$ high pass filter.

The energy source used in this study was the Boomer. Energy is produced when two plates are driven apart suddenly by a heavy surge of electrical current through a coil on one of the plates which generates eddy currents in the other plate, resulting in it being suddenly repelled. The Boomer is a Scheidegg Monopulse Acoustic Generator, type MAG 30V as mentioned in (2) above.

The hydrophones are mounted in a long streamer towed behind the seismic ship at a depth between 1 and 2 $\,$ m.

Shooting takes place at a speed of the order of 6 knots. The transects where shooting took place are on FIG.2.

APPENDIX 1.B.

VIBROCORING AND CORE PREPARATION

The vibrocorer used for obtaining undisturbed samples consists of a submersible coring rig with power unit which, in operation, is linked to the surface by a power cable from a shipboard generator and lifting cable. A T-shaped frame maintains the core barrel in an upright position during coring.

The vibrocorer used is capable of procuring 4.5 m. cores. The plastic liner tubes containing vibrocores were cut into 1 m. lengths in the field and transported in a vertical position to the laboratory in order to minimize disturbance.

In the laboratory, they were opened by making longitudinal cuts through the plastic liner with a machine especially designed just to cut the plastic liner and not the sediment. The sediment itself is then cut with a pallette knife. The two halves of the core are treated differently according to the information required from sedimentological analysis. Usually, one half is used for grain-size and CaCO₃ analysis, whilst the other half is used for descriptive and photographic purposes. 75 cores were analyzed and described.

APPENDIX 1.C.

SIZE ANALYSIS

- 1. PREWEIGHED AND NUMBERED 250 ML. BEAKERS.
- 2. DRIED AND PREWEIGHED 15.0 CM. AND 12.5 CM. NO.54, HARDENED, FILTER PAPER.
- 3. PLACE IN BEAKERS APPROX. 5 TO 10 GRAMS OF SAMPLE.
- 4. ADD DISTILLED WATER, ALLOW TO STAND 2 HOURS, SIPHON OFF. AND DRY IN OVEN SET AT APPROX. 100°C.
- 5. COOL AND WEIGH (SAMPLE WEIGHT).
- 6. ADD 110 TO 150 ML. OF Na.HEXA.META.PHOSPHATE, STIR UNTIL SAMPLE WELL DISPERSED AND THEN WARM IN OVEN OVERNIGHT. OVEN AT APPROX.80°C.
- 7. SIEVE THROUGH SAND 63. AND COARSE SILT 38. SIEVES. COLLECT FINE SILT AND CIAY IN BEAKER, STIR AGAIN AND ALLOW TO SETTLE HALF AN HOUR.
- 8. THEN FILTER FINER FRACTION THROUGH 15 CM. NO.54. FILTER PAPER (FINE SILT).
- 9. DRY FILTER PAPER WITH SEDIMENT, IN OVEN AT APPROX. 100°C. WEIGH.
- 10. ADD WEIGHTS OF SAND, COARSE SILT AND FINE SILT AND THEN SUBTRACT THIS FIGURE FROM SAMPLE WEIGHT. THIS LEAVES WEIGHT OF CLAY.

NOTE: Samples were taken from the top and bottom of cores and also from lithological units in which the author was interested. From the sample, 5 to 10 grams was used for grain-size analysis, and 2 to 3 grams was used for carbonate analysis.

APPENDIX 1.D.

Caco 3 ANALYSIS

- 1. PREWEIGHED AND NUMBERED 50 ML. BEAKERS.
- 2. PLACE 2 TO 3 GRAMS IN BEAKERS.
- 3. ADD DISTILLED WATER, ALLOW TO STAND FOR 2 HOURS, SIPHON OFF AND LET DRY IN OVEN AT APPROX. 100°C.
- 4. COOL AND WEIGH TO 2 DECIMAL PLACES.

 (WEIGHT OF SAMPLE).
- 5. ADD H.CL. UNTIL DIGESTED, AND THEN SIPHON OFF.
- 6. WASH AND ALLOW TO STAND FOR 2 HOURS AND THEN SIPHON SAMPLE 3 TIMES.
- 7. DRY AND REWEIGH.
- 8. PERCENTAGE CaCO $_3$ = WEIGHT LOSS/ORIGINAL WEIGHT X 100.

APPENDIX 2: RESULTS FROM METHODS OF STUDY

- 2.A. Selected descriptions of submarine vibrocores
- 2.B. Grain-size analysis
- 2.C. CaCO₃ analysis
- 2.D. Sand and organism count results

APPENDIX 2.A.

SELECTED DESCRIPTIONS OF SUBMARINE VIBROCORES

ABBREVIATIONS USED

FORAMS.	FORAMINIFERA
Τ.	GENUS <u>TRILOCULINA</u>
E .	" <u>ELPHIDIUM</u>
D.	" <u>DISCORBIS</u>
Q.	" QUINQUELOCULINA
S.	" SPIROLOCULINA
P.	" PENEROPLIS
N.	" <u>NUBECULARIA</u>
BI.	BIVALVE
T.G.	TURRETED GASTROPOD
N.T.G.	NON-TURRETED GASTROPOD
ECH.SPINE	ECHINOID SPINES
COR.ALG.	CORALLINE ALGAE
BRY.	BRYOZOA
A.NEEDLES	ARAGONITE NEEDLES
COMM.SHELL	COMMINUTED SHELL
O.M.	OMARMA
QTZ.	QUARTZ
P.S.	POORLY SORTED
M.S.	MODERATELY SORTED
W.S.	WELL SORTED
R.	ROUNDED
S.R.	SUB_ROUNDED
S.A.	SUB-ANGULAR
Α,	ANGULAR
POL.	POLISHED
FeOx	IRON OXIDE
CALC.	CALCAREOUS
>	GREATER
<	LESS
THRU	THROUGH
%	PERCENTAGE
Į	WITH DEPTH, OR, TOWARDS BASE
HOL.	HOLOCENE
DIETO	DI DI CIDADID

PLEISTOCENE

PLEIS.

APPENDIX 2.A.

SELECTED DESCRIPTIONS OF SUBMARINE VIBROCORES

LINE II WHYALLA CENTRAL

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
S G11 <i>5</i>			
0-15	FORAMST., Q.AND S., COARSE		HOL.
	BI.HASH & T.G.HASH AND ECH.		
	SPINES, A.NEEDLES, COR.ALG.,		
	AND COMM.SHELL. SMALL % GREY	PRESENT BED	
	STAINING. PALE GREY/BROWN	LOAD	
•	SILT. SHELLY SILT. ORGANIC		
	MATTER PRESENT.		
15-105	FORAMSFEW-Q.AND S. AND BI.	SUBTIDAL	
	HASH. COARSE SHELL.MAINLY		
	GREY SILT. ORGANIC MATTER		
	PRESENT.		
105-165	FORAMST., Q.AND S. MAINLY,		
	WITH FEW E. COARSE BI.HASH	INTERTIDAL	PLEIS.
	AND T.G.s, ECH.SPINES AND		
	COMM.SHELL 60% SHELL, 5%		
	QTZ. GREY/GREEN IN COLOUR.		
	NODULES THRU.OUT. $>$ SIIT \checkmark		
	NO WHOLE SHELLS.		
165-240	MAINLY COARSE SHELL. BI.		
•	HASH, T.G.'s, N.T.G.'s, ECH.	SUBTIDAL	
	SPINES AND COMM.SHELL. WHOLE		
	BI.s. SILT CHANGES COLOUR		
	FROM GREEN/BROWN AT TOP TO		
	PALE BROWN/WHITE AT BASE.		
	NODULOSE.		
240-400	QTZ.CONTENT VARIES FROM 10%		
	AT TOP TO 40% AT BASE. M.S.,	SUPRATIDAL	
	S.RS.A., POL., FeOx STAIN.		
	PLUS SILT. GREEN/WHITE AT		
	TOP TO YELLOW/BROWN AT BASE.		
	COMM.SHELL-5% MAXIMUM.		
	MAINLY SILT. OXIDIZED.		

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG117			
0-170	FORAMSQ.AND S. BI.HASH, T.G.s	SUBTIDAL	
	AND A.NEEDLES PLUS COMM.SHELL.		HOL.
	GREY STAINING ON SHELL.		
	AVERAGE OF 70% SILT-GREY.		
	SHELLY SILT.		
170-247	FORAMSQ.AND S. VERY COARSE		
	BI.HASH AND T.G.s AND COMM.		
	SHELL. QTZ-P.S., S.AS.R.,	SUBTIDAL	PLEIS.
	POL AND FROSTED, FeOx STAIN		
	SILT-OXIDIZED, YELLOW/ORANGE		
	COLOUR. MEAN PERCENTAGES OF		
	35% SHELL, 25% QTZ., 40% SIIT		
	WHOLE BIVALVES.		
247-260	COARSE BI.HASH & WHOLE VALVES	BEACH	
	AND COMM.SHELL. FINE, OFFWHILE		
	SILT.		
260-315		SUPRATIDAL	
	SLIGHTLY OXIDIZED.		
31 <i>5</i> -330	5% SHELL, 55% QTZP.S., S.A	INTERTIDAL	
	S.R., POL.AND FROSTED, FeOx		
	STAINED RED. WHITE SILT.		
330-345		SUPRATIDAL	
345-365	IDENTICAL TO 315-330	INTERTIDAL	

NOTE: 347 to 365 ONE SEQUENCE.
SUPRATIDAL WITH INTERTIDAL
PROBABLY SPRING OR STORM-TIDE
DEPOSITS ON SUPRATIDAL FLAT.

DEPTH (CM.) SG119	DESCRIPTION	ENVIRONMENT	<u>AGE</u>
0-100	FORAMSQ.AND S. MAINLY BI. HASH & FEW T.G.s. 10% SHELL 90% FINE GREY SILT.	PRESENT BED	HOL.
100-130	FORAMST.,Q.,S.,E., AND P. ECH.SPINES, T.G.s, AND BI. HASH. SEA GRASS PRESENT. 50% SHELL, 50% SILT.	SUBTIDAL	
130-160	FORAMSFEW-E., Q., AND S. BI.AND T.G.HASH, COR.ALG., AND BRY.		
	QTZ P.S., S.RS.A., POL. FeOX STAINING ON QTZ AND SILT. SILT-VERY FINE, GREY/WHITE. 45% SHELL, 15% QTZ, AND 40% SILT.	INTERTIDAL	PLEIS.
160-185	VERY COARSE BI.HASH AND T.G.s. VERY PALE, ORANGE/BROWN, NODULAR, SILT ADHERES TO THE SHELL. 50% SHELL, 50% SILT.	INTERTIDAL	
	FORAMSMANY-T.,Q., AND S. WITH BI.HASH AND T.G.s. 30% SHELL.70% VERY FINE, OFF WHITE SILT.	SUPRATIDAL/ INTERTIDAL	
215-295	FORAMST.,Q.AND S. COARSE BI. AND T.G.HASH AND COMM.SHELL. QTZP.S., S.RS.A., POL.AND	SUBTIDAL/	
	•	INTERTIDAL	
	SIMILAR TO 215-295 BUT > SILT AND NO COARSE SHELL. > QTZ. AND > FINE SHELL. OXIDIZED PINKY/BROWN COLOUR.	INTERTIDAL	

DEPTH (CM	DESCRIPTION	ENVIRONMENT	AGE
SG121			
0- 20	BI.HASH AND FEW T.G.s.AND COMM.	AFTER	PLEIS.
	SHELL. OFFWHITE, PALE BROWN SILT. MAINLY SHELL. SHELLY SILT.	DREDGING	
20-100	SPINES AND COR.ALG.	•	HOL.
	70% GREY SILT. ORGANIC MATERIAL PRESENT. COARSE NODULES AT BASE.	ı	
100-200	FORAMST.,S.AND Q. ECH.SPINES, COARSE BI.AND T.G.HASH AND COMM.SHELL. OFFWHITE, PALE BROWN SILT. MAINLY SHELL, COARSE SHELL. NODULOSE.	SUBTIDAL	PLEIS.
200-345	MARGINOPORA, T.G.s, N.T.G.s., ECH.SPINES, BI.HASH AND AN ANADARA TRAPEZIA (BROKEN) AND COMM.SHELL. 20% QTZ., P.S.ALTHOUGH MAINLY FINE GRAIN-SIZE, S.RS.A., POL.AND FeOx STAINED. > COARSE SHELL ✓, < SILT ✓	INTERTIDAL	

LINE 33 WHYALLA CENTRAL

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG83			
0-125	FORAMSQ.AND S. BI.HASH,		
	T.G.s AND A.NEEDLES.		
	30% VERY FINE SHELL HASH.	SUBTIDAL	
	70% GREY/BROWN SILT, PLUS OR-		
	GANIC MATERIAL.		
125-145	PROLIFIC FORAMSE.,P.,Q.AND		
	S. BI, T.G.AND N.T.G.HASH.		
	BLUE/GREY STAIN ON SHELL. 70%	SUBTIDAL	HOL.
	SHELL AND 30% ORGANIC RICH SILT.		•
	GREY COLOUR.		
145-375	FORAMST., Q.AND S. VERY COARSE		PLEIS
	BI.HASH AND MINOR T.G.s AND ECH.	. SUBTIDAL	
	SPINES. MEAN % OF 70% SHELL.		
	SMALL % QTZ.		
	CREAMY/WHITE SILT. MEAN % OF		
	30% SIIT.		
375-395	SMALL % SHELL. 50% QTZ. REST	INTERTIDAL	
	WHITE SILT.		

DEPTH	(CM.)	DESCRIPTION	ENVIRONMENT	<u>AGE</u>
SG8	35			
0-	35	FORAMST., E., Q.AND S. BI.HASH,		HOL.
		T.G.s, N.T.G.s, ECH.SPINES,		
		AND COR.ALG. COR.ALG.VERY	SUBTIDAL	
	•	PROMINENT. GREY STAINING ON		
		SHELL. SEAGRASS. 5% GREY SILT.		
		SHELLY SILT.		
35-1	20	25% BI.HASH, N.T.G.s AND ECH.	INTERTIDAL	PIEIS.
		SPINES.		
		50% QTZ., P.S., S.AS.R., POL.	(BEACH)	
		25% ORANGE/BROWN SILT. SMAIL		
		% ORGANIC MATTER.		
120-2	240	APPROX.5% QTZ., REST FINE	SUPRATIDAL	
		SHELL AND SILT. NODULOUS.		
		ORANGE/BROWN V RED/BROWN SILT.		
240-3	310	BI.HASH, T.G.s, N.T.G.s, COR.		
		ALG., AND ECH.SPINES.		
		APPROX.50% SHELL AT TOP TO	INTERTIDAL?	
		80% SHELL AT BASE. COARSE SHELL.	SUBTIDAL	
		REST A PALE OFFWHITE COLOURED		
		SILT.		

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
sg86			
0-158	FORAMSE.,T.,S.AND Q. BI.HASH,		HOL.
	T.G.s, ECH.SPINES, AND COR.ALG.	*********	
	PLUS LIGHT GREY DARK GREY.	SUBTIDAL	
7.50.000	>SILT AND > ORGANIC MATTER ♥ .	ØYDD A MITD A T	DESTA
158-200	NODULAR PALE BROWN SILT. FEW COARSE FRAGMENTS.	SUPRATIDAL	PLEIS.
200-225	E.,Q,AND S. BI, T.G.AND N.T.G.	INTERTIDAL/	
	HASH. PLUS ECH.SPINES AND COR.	SUBTIDAL?	
	ALG. REST KHAKI/GREEN SIIT.		
	40% SHELL. SILT ADHERES TO		
	SHELL.		
225-260	FORAMSE.,S.AND Q. BI., T.G.s,	INTERTIDAL	
•	ECH.SPINES AND COR.ALG.		
	100% SHELL. COARSE SHELL.		
	NODULES.		
260-283	ORANGE/BROWN NODULAR SILT.	SUPRATIDAL	
SG126			
0- 15	FORAMST., Q.AND S. BI.HASH PLUS	DUMPED	PLEIS.
	COR.ALG., ECH.SPINES AND A.NEE-	AFTER	
	DLES. BASICALLY BI.HASH AND	DREDGING	
	FINE, LIGHT BROWN SILT. SHELLY		
	SILT.		
15- 90	FORAMST. BI.HASH AND FEW T.G.s	SUBTIDAL	HOL.
	30% SHELL 70% GREY SILT.		
	BURROWS PRESENT.		
90-105	CALC.FINE, PALE, YELLOW SILT.	SUPRATIDAL	PLEIS.
105-204	FORAMST.,Q.AND S. BI.HASH,		
	T.G.s, N.T.G.s, COR.ALG., ECH.		
	SPINES AND BRY. 70% SHELL.		
	10% QTZ. M.S., S.R., POL.AND	INTERTIDAL	
	FeOx STAINED. 20% ORANGE/BROWN		
	SILT. CALCAREOUS CRUST AT BASE.		
	NODULES THRU.OUT.		
	WHITE CALC.CLAY.	SUPRATIDAL	
280-360	99% QTZ. W.S., S.R., POL.AND	AEOLIAN	
	FROSTED. SLIGHTLY OXIDIZED		

QTZ.-2 POPULATIONS. FINE, S.R. SUPRATIDAL

S.A., POL., AND COARSE S.R.-R. POL. 10% QTZ.AT TOP INCREA-

ORANGE TO RED/BROWN SILT.

SES TO 30% AT BASE.

NODULOSE. OXIDIZED.

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
\$G132	TODAMO TO MOANDS FEW		
0-130	TUBES. SEA GRASS PRESENT. SMALL % PALE-GREY SILT. SILT AND ORGANIC CONTENT INCREASE TO AN EXTENT THAT AT BASE 5% SHELL, 10% ORGANIC MATTER,	SUBTIDAL	HOL.
130-155	85% GREEN/GREY SILT. 10% SHELL HASH-FORAMS., A. NEEDLES AND WORM TUBES ARE		PLEIS.
	PROMINENT. 10% QTZP.S., S.RS.A., POL. ORGANIC MAT- TER PRESENT. 80% ORANGE/BROWN SILT.	:	
155-202	> 40% SHELL HASH-SIMILAR TO ABOVE BUT COR.ALG. PROLIFIC. SILT ADHERES TO SHELL AND QTZ. LARGE T.G. BROWN SILT.	SUBTIDAL	
202-300	FORAMST.,E.,S.AND Q. SIMI- IAR TO 155-202 BUT > SHELL. A.NEEDLES-PROLIFIC. IARGE BI.S. AND T.G.S. COR.ALG. AISO PROMINENT. SMALL T.G.S & N.T.G.S ALSO. NOLULES. QTZP.S., S.AS.R., POL 10%. 10% SILT. 80% SHELL. OFFWHITE COLOUR.	INTERTIDAL	
300-360	BROKEN ANADARA TRAPEZIA PRESENT. 30% SILT AT TOP INCREASING TO 70% AT BASE. QTZ.VERY FINE. SHELL AND QTZ. %5 DECREASE & DEPTH. BOTH SILT AND QTZ.OXIDIZED. DEPO- SIT GREEN AT TOP, BROWN/GREEN	SUPRATIDA <u></u> L	

SG132 CONTD.

340-415

AT BASE. SMALL % OF BLACK-STAINED SHELL AT TOP. NO COARSE SHELL.

SG237 0- 25 FORAMS.-T., BI., N.T.G. AND T.G. HOL. HASH AND A.NEEDLES. QTZ.-P.S. SUBTIDAL S.A.-S.R., POL., SOME FROSTED. 10% MAXIMUM. 10% GREY SILT. 80% SHELL. FORAMS.-T.-FEW. BI & T.G.HASH. 25-100 ALSO A.NEEDLES ARE PROMINENT. NODULAR. 70% SHELL AT TOP DOWN INTERTIDAL PLEIS. TO 30% SHELL AT BASE. 0% QTZ. AT TOP INCREASING TO 40% QTZ. AT BASE, PALE ORANGE AT TOP, THRU TO PALE GREY AT BASE. 10% SHELL HASH. 35% QTZ., P.S., SUPRATIDAL/ 100-140 S.R.-S.A., POL. AND REST CALC. INTERTIDAL GREEN CLAY. PINKY-BROWN BOUNDARY COLOUR AT BASE. NODULES. 140-255 10% SHELL HASH MAXIMUM. ECH. SPINES AND A.NEEDLES. 40-50% QTZ.-M.S., S.A.-R., POL.AND FeOx STAINED. REST WHITE, CALCAREOUS, NODULAR SILT. 255-340 FORAMS.-T., A. NEEDLES AND BI. HASH. 30% SHELL. 60% QTZ.-P.S., S.R., POL., FeOx STAINED. INTERTIDAL -10% PALE-ORANGE SILT. FINE

ROCK FRAGMENTS ALSO PRESENT.

NOTE: ALLUVIAL FAN OR DECAL-

100-140: OXIDIZED AND THEN RE-DUCED SECTION OF 100-255 SUPRA-

CIFIED SUPRATIDAL DEPOSIT.

TIDAL/INTERTIDAL DEPOSIT.

NO FAUNA NOR QTZ.

BROWN/RED NON-CALCAREOUS CLAY. ALLUVIAL

FAN

LINE 30 WHYALLA NORTH

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG225			
0-105	FORAMS., T.G.s, ECH.SPINES, A.NEEDLES AND BI.HASH. GREY SILT. > SILT \(\sqrt{.} \)	SUBTIDAL	HOL.
105-180	FORAMS., T.G.s, A.NEEDLES AND BI.HASH AND 15% P.S., S.R., POL.AND FROSTED, FeOx-STAINED QTZ. OFFWHITE AT TOP TO YELLOW/BROWN AT BASE. GREATER	INTERTIDAL	PLEIS.
	OXIDATION WITH DEPTH.		
180-225	GREATER THAN 80% QTZW.S., S.R., FeOx STAINED. RED SILT. NO SHELL.	AEOLIAN	
225-240	FEW FORAMS. AND 25% QTZM.S., R., POL. 75% WHITE SILT.	SUPRATIDAL	
240-270	SMALL % SHELLY MATERIAL. < 5% QTZ. LARGE S.A. ROCK FRAGMENTS. REST NODULOSE. PALE-GREEN SILT.	SABKHA	
270-310	HIGHLY CALC., WHITE CLAY.	SABKHA	
310-342	COARSE BROKEN SHELL AND 50% P.S., S.RS.A., FROSTED AND		
	POL. AND FeOX-STAINED QTZ. SLIGHTLY OXIDIZED. LESS CLAY ↓.	BEACH.	
342-375	80% W.S., S.R. FeOx-STAINED QTZ. 20% SILT AND ECH.SPINES. PINK IN COLOUR.	AEOLIAN	
	COARSE ROCK FRAGMENTS AT BOUN- DARY. WHITE CALC.CLAY. SMALL % QTZ.	SABKHA	
425-430	IDENTICAL TO 342-375.	AEOLIAN	

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG227			
0-100	FORAMST.,E.,Q.AND S. WHOLE T.G.S AND BI.S AS WELL AS BI. AND T.G.HASH. GREY STAINING ON SHELL. 10% M.S., S.RR., POL.QTZ. 80% SHELL. 10% GREY	SUBTIDAL	HOL.
100-140	SILT. > SILT, < QTZ. V . < 5% SHELLY MATERIAL AND OR-		PLEIS.
140	GANIC MATERIAL. 40% P.S., S.RR., POL.QTZ. FeOx STAINED. REST GREY/BROWN SILT. AVERAGE OF 40% SHELL INCREA-	PALEOSOL	
140	SING AT BASE. A.NEEDLES, BI.s	INTERTIDAL	
	AND T.G.s, WHOLE AND HASH.		
	MEAN OF 20% M.S., S.RR., POL., QTZ. REST NODULAR, WHITE,		
	CAIC.CLAY.		
0.00-5			
SG231 0-100	FORAMS., BI.s, T.G.s WHOLE AND	SHRTTDAT.	t HOL.
02100	HASH, AND GREY SILT. GREY SHELLY SILT.		
100-135	SMALL % SHELL HASH-BI.s AND	AEOLIAN	PLEIS.
	T.G.s. REST QTZ. W.S., S.R S.A., FROSTED QTZ. BROWN IN		
	COLOUR.		
135-160	5% QTZ. REST WHITE, CAIC., NO- DULOSE SILT.	SABKHA	
160-295	QTZ. AND SHELL. SHELL-FORAMS.,		
	ECH.SPINES, T.G.s AND BI.s.		
	QTZM.S., S.AS.R., POL.AND	BEACH	
	FeOx STAINED. PALE YELLOW		

COLOUR AT TOP TO PINK AT BASE.

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
S G233			
* *	COARSE SHELL HASH-FORAMS., BI.s, A.NEEDLES, T.G.s. HIGH % OF SHELL STAINED BLACK, AND SILT. GREY SHELLY SILT.	S UBTIDA L	HOL.
40-120	IARGE NODULES PROMINENTQTZ., SHELL AND CAIC.CEMENT. 50% QTZ. SHELL-BI.s, T.G.s, AND FORAMS. MAINLY, AND YELLOW/BROWN SILT. BROWN IN COLOUR.	INTERTIDAI	PLEIS.
SG284			
0-110	COARSE SHELL HASH-BI.s, FORAMS., T.G.s, WHOLE AND HASH. SEA GRASS, AND SILT. > SILT ↓. GREY SHELLY SILT.	SUBTIDAL	HOL.
110-380	COMM.SHELL AND FINE-GRAINED QTZ. MOSTLY CALC.CLAY. 110-275: WHITE IN COLOUR. 275-380: YELLOW/BROWN WITH RED PATCHES.	SABKHA	PLEIS.
\$G285			
	COARSE SHELL-FORAMS., T.G.s, N.T.G.s, BI.s, AND A.NEEDLES. BLACK STAINING ON SHELL. SEA GRASS PRESENT, AND SILT. SHELLY GREY SILT. NOTE: ZONES WHERE BROWN IN	SUBTIDAL	HOL.
150-320	COLOUR AND NODULES PRESENT. 2% SHELL HASH-BI.s, T.G.s, AND FORAMS. 30% M.S., S.AS.R., POL., FINE QTZ. BURROWS. REST SILT-GREEN IN COLOUR. 150-255. GREATER OXIDATION TO A RED COLOUR 255-320. NOTE: ORIGINALLY OXIDIZED CLAY (RED)	SUPRATIDAL	PLEIS.

REDUCED AT TOP TO GREEN COLOUR.

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
S G294			
0- 13 6	0% SHELL. FORAMST.,E.,S.	SUBTIDAL	HOL.
A	ND Q. BI.HASH AND T.G.s.		
40	0% GREY SILT.		
13- 40 2	5% SHELL HASH-BI.s, T.G.s		PLEIS.
A:	ND FEW FORAMS.		
2	5% P.S., S.RR., POL.QTZ.	INTERTIDAL	
50	0% RED SILT. BEDS AS DESCRI-		
Bi	ED ABOVE WITH BEDS OF JUST		

RED SILT IN BETWEEN.

4:

LINES 56, 57, 58, 59, 60 PORT AUGUSTA

DEPTH	(CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG	54			
0-	20	10% BI., T.G. AND N.T.G. HASH.	PRESENT	
		80% QTZW.S., RS.R. AND	BED LOAD	
		POL. AND 10% SEA GRASS.		
20-1	L 50	20% BI., T.G. AND N.T.G. HASH.	ESTUARINE	HOL.
		60% FINE QTZW.S., RS.R.		
		POL. 15% BROWN SILT & 5% SEA		
		GRASS.		
150-2	250	10% BI.HASH. > SHELL AT BASE,		PLEIS.
		INCLUDING COARSE BI.s, N.T.G.s,		
		T.G.s, AND ANADARA TRAPEZIA.	INTERTIDAL	
		AT TOP 50% QTZ., 40% RED SILT.		
		THESE %s DECREASE WITH DEPTH.		
		QTZP.S., RS.R., POL.AND		
		FROSTED AND FeOx STAINED.		
250-1	310	FORAMSE. 10% SHELL INCLUDING		
		ECH.SPINES, BI.AND T.G.HASH.		
		45% QTZFeOx STAINED AND 45%	TMTEKTIDAL	
		GREEN SILT. HEAVILY BURROWED		
210	220	IN SECTIONS.		
310-1	3 30	FORAMSQ.AND S. BI.HASH, COR. ALG., ECH.SPINES AND ANADARA		
		TRAPEZIA AND MARGENOPORA.	INTERTIDAL	
		50% SHELL. 25% QTZ. AND 25%	11/10111110111	
		CREAM-COLOURED SILT.		
330-	380	PALE-GREEN AND YELLOW SILT	SUPRATIDAL	
))U=.		WITH COMM.SHELL. ROOTIETS		
		PRESENT.		

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG 55			
0- 60	FORAMSE.,P. AND D. BI.HASH,		
	ECH.SPINES AND T.G.s. 40% SHELL,		
	· · · · · · · · · · · · · · · · · · ·	SUBTIDAL	HOL.
	GRASS. 40% QTZW.S., S.R.AND		
	POL.		D#-T0
60-110	FORAMSP. BI.HASH. 5% SHELL.		PLEIS.
	50% QTZP.S., RS.A., POL.	INTERTIDAL	
•	AND FROSTED AND FeOx STAINED,		
	AND 50% PALE-ORANGE/BROWN SILT.		
110-165	SIMILAR TO ABOVE BUT > SHELL.	SUBTIDAL	
	LARGE BI.s, T.G.s AND N.T.G.s.		
	WHOLE FORMS FOUND.		
165-200	FORAMSQ.AND S. 10% FINE SHELL	L SUPRATIDAL	
	HASH OF BI.s, N.T.G.s AND T.G.s.	or	
	45% QTZP.S., RS.A., POL.	ESTUARINE	
	AND FROSTED AND FeOx STAINED.	or	
	45% YELLOW SILT.	INTERTIDAL	
200-360	5% COARSE SHELL-BI.HASH AND	INTERTIDAL	
	WHOLE T.G.s. 40% FINE QTZ.AND		
	60% FINE YELLOW SILT AND COMM.		

SHELL.

DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE
SG 56			
0- 15	BROWN, QTZRICH, SILTY,		
	SHELLY DEPOSIT. SHELL GREY-		
	STAINED. ROCK FRAGMENTS		
	PRESENT.	WASH	
15- 23	SLIGHTLY CALCAREOUS, QTZRICH,		
	AQUA-GREEN CLAY.	FROM	
23- 40	70% ROCK FRAGMENTS AND QTZ		
	P.S., R., POL.AND FROSTED AND	CLIFFS	
	FeOx STAINED. PALE SILT AND		
	5% SHELL MAKE UP REST OF DEPO-		
	SIT. ORANGE AT TOP TO PALE		
	BROWN AT BASE.		
40-195	FORAMST.,Q.AND S. WHOLE BI.s		
	AND BI.AND T.G. HASH. 55%		
	SHELL. SEA-GRASS FIBRE PRESENT	•	
	SOME SHELL HIGHLY STAINED. 25%		
	FINE QTZ. AND 20% FINE BROWN		
	SILT. GREY/BROWN COLOUR.		HOL.
	>QTZ. AND SILT .		D. F. T. C.
195-223	5% SHELL, 70% QTZ. AND ROCK	WASH FROM	PLEIS.
	FRAGMENTS. QTZP.S., R., POL.	####### OD	
	AND FROSTED AND FeOx STAINED.	CLIFFS OR	
	SILT YELLOW AT TOP TO PALE		
	OFFWHITE AT BASE.	BEACH	
223-235	5% SILT. NO SHELL. ONLY QTZ.		
	AND ROCK.	CLIFFS OR	
	DARK-ORANGE COLOUR.	BEACH	
235–360	80% PEBBLES AND 20% QTZ.	WASH FROM	
		CLIFFS OR	
	•	BEACH	

DEPTH (CM.)	DESCRIPTION	<u>ENVIRONMENT</u>	AGE
SG156			
0- 10	60% BI.HASH AND WHOLE BI.s	PRESENT	
	AND N.T.G.s. 20% BROWN SILT.	BED LOAD	
	20% DARK ORANGE/RED FeOx-		
	STAINED QTZ. QTZW.S., R.		
	AND POL.		
10- 50	SIMILAR TO O-10 BUT < SHELL		
	AND > SILT. GREY/BROWN SILT.	SUBTIDAL	HOL.
	GREY STAINING ON SHELL.		
50- 83	60% W.S., S.R., POL., QTZ.	AEOLIAN	PLEIS.
	(SIIT SIZE) AND 40% BROWN SIIT.		
83-100	SIMILAR TO 50-83 PLUS AQUA-		
	GREEN CLAY NODULES WHICH BREAK		
	UP TO CAUSE GREY COLOUR. 2% BI.		
100-110	SIMILAR TO 50-83 BUT > SILT.		
110-125	AQUA-GREEN IN COLOUR OVERALL,		
	OTHERWISE IDENTICAL TO 50-83.		
125-330	LIGHT-GREEN AND ORANGE SILT	ESTUARINE	
	WITH UP TO 5% COMM.SHELL.		
	POCKETS OF COARSE SHELL -		
	ANADARA AND OTHER BI.s.		
SG159			
0- 15	40% FINE COMM.SHELL, GREY	PRESENT	
	STAINED. 20% FINE QTZ. 40%	BED LOAD	
	PALE-BROWN/GREY SILT.		
15-145	FORAMSS.AND Q. 20% SHELL-	SUBTIDAL	
	5% BI.HASH, 15% FORAMS., 2%		
	SEA GRASS. 20% FINE QTZ. REST		
	GREY SILT.		
145-175	> SEA GRASS, > SILT. NO QTZ.	SUBTIDAL	
175-200	BLUE/GREY AND LIGHT-BROWN SILT		
	WITH ORGANIC MATERIAL.		HOL.
200-220	VERY FINE QTZ RICH CLAY.	DUNE	PLEIS.
220-310	90% W.S., S.R., POL., FeOx-	INTERTIDAL	
	STAINED QTZ. 10% PALE-BROWN		
	SILTADHERES TO QTZ.		
310-353	50% P.S., RA., POL.AND FROS-		
	TED QTZ. AND 50% RED SILT.	\mathtt{FLAT}	

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DEPTH (CM.)	DESCRIPTION	ENVIRONMENT	AGE		
SG160 0- 70	FORAMSQ.,S.,E.AND P. BI.HASH				
	AND T.G.S. 40% SHELL AND 60% GREY SILT AND ORGANIC MATERIAL. WITH DEPTH THE SHELL % DE-CREASES, SILT AND QTZ. %.s IN-CREASE.	SUBTIDAL	HOL.		
70-100	FORAMSE.,P.,D.,S.AND Q. AND BI.HASH. 15% SHELL. 70% W.S., R., POL., QTZ. AND 15% PALE-BROWN SILT. COARSE BI.HASH.	INTERTIDAL	PLEIS.		
100-240	10% SHELL AT MOST. COMM.SHELL MAINLY WITH FEW WHOLE BI.S. 30% QTZMAXIMUM. REST FINE	TOWAY DINT			
	CALCAREOUS SILT. COLOUR CHAN- GES FROM PALE GREEN AT TOP TO YELLOW-GREEN TO KHAKI AT THE BASE.	ESTUARINE			
SG162					
0-130	20% SHELL-ECH.STEMS, COR.ALG., T.G.s, AND BI.HASH. 70% P.S., S.AS.R., POL.AND FROSTED AND FeOx-STAINED QTZ. 10% SILT. ORANGE/BROWN SAND WITH ROCK FRAGMENTS. NON-CALC., QTZRICH, RED SILT AT BASE-ALLUVIAL.	SUBTIDAL			
	POSSIBILITY THAT HOL.HERE FORMED PARTLY BY REWORKED ALLUVIAL FLAT MATERIAL.		HOL.		
130-220	GREY, CALCAREOUS, FINE QTZ RICH CLAY. ONE ANADARA TRAPE- ZIA FOUND. ONLY SHELL NOTED.	ESTUARINE	PLEIS.		
220-233	10% SHELL-BI.HASH. 80% P.S., RS.A., POL.AND FROSTED. FeOx-STAINED QTZ. AND ROCK FRAG-MENTS. 10% YELLOW SILT. POCKET OF NON-CAIC.AQUA-GREEN CLAY.				

APPENDIX 2
B. GRAIN
GRAIN SIZE ANALYSIS.
• WHYALLA
NORTH.

0.82	2.20	3.62	1.42	0.70	1.78	1.08	2.17	3.26	1.09	5.90	Top 25.79	=	19.88	78
0.29	0.87	2.48	1.61	0.38	1.49	- <u>-</u>	4.12	5.19	1.07	5.68	130 25.51	=	19.83	17
0.97	1.12	2.52	1.40	0.28	1.34	1.06	4.71	5.82	 	7.08	140 27.35	=	20.26	16
0.46	0.73	2.33	1.60	0.46	1.55	1.09	44.9	7.54	1.10	8.11	Bot 28.83	SG222	20.72	15
0.72	2.05	3.62	1.57	0.37	1.45	1.08	4.17	5.25	1.08	7.23	Top 28.00	Ξ	20.68	14
0.12	0.52	1.97	1.45	0.12	1.21	1.09	5.13	6.24	-1 -1 -1	5.90	130 26.09	=	20.18	13
0.62	0.58	2.19	1.61	0.19	1.29	1.10	4.26	5.36	1.10	5.67	180 25.88	=	20.21	12
0.73	0.88	2.29	1.41	0.18	1.27	1.09	4.59	5.69	1.10	6.40	222 28.01	=	21.61	7
0.67	0.43	2.00	1.57	0.20	1.30	1.10	5.09	6.16	1.07	6.40	270 26.99	S G223	20.58	10
0.67	1.30	2.70	1.40	0.37	1.47	1.10	6.40	7.48	1.08	8.76	Top 28.93	=	20.16	9
1.03	0.67	2.33	1.66	0.15	1.21	1.06	5.66	6.75	1.09	7.52	150 29.12	=	21.60	œ
0.42	0.68	2.11	1.43	0.21	1.30	1.09	3.50	4.59	1.09	4.82	195 25.10	=	20.28	7
0.56	0.25	1.83	1.58	0.06	1.15	1.09	2.49	3.56	1.07	3.37	240 23.95	=	20.58	- 1
0.65	0.26	1.68	1.42	0.11	1.22		3.68	4.76	1.08	4.71	300 24.95	7	20.23	VI
2.86	2.08	3.66	1.58	0.27	1.37	1.10	4.16	5.23	1.07	9.38	315 29.27	=	19.89	7
0.56	0.31	1.72	1.41	0.083	1.18	1.10	5.74	6.83	1.09	6.70	330 28.32	=	21,61	w
0.26	0.29	1.93	1.64	0.051	1.17	1.12	5.46	6.56	1.10	6.07	360 25.94	=	19.87	2
0.08	0.20	1.64	1.44	0.044	1.16	1.12	6.38	7.50	1.12	6.71	· 400 28.31	SG224	21.59	ل بيب
ms.	wt in grams.	F.Silt			C.Silt			Sand		Wt.	• &Samp Wt	No	Wt.	MO.
Clay	F.Silt	ች ₽ +	다	C.Silt	н Р +	Ħ P	Sand	+ +	ب ب ب	Sample		. Sample	Beaker	Beaker

/ Cont.

19	18	17	16	7	14	Ω 1	7	1	10	9	Φ	7	0	J	F	ىب	N	=1	20	19
19.87	19.88	19.83	20.26	20.72	20.68	20.18	20.21	21.61	20.25	20.16	21.60	20.28	21.56	20.23	19,89	21.61	19.87	21.59	21.62	19.87
=======================================	8G230 T	= ယ	= 2	= 2	SG228 T	= B	= -1	SG227 T	= ω	==	SG226 T	÷ :	= ω	:	= 20	: N	=	SG225 T	:: H	SG221 B
170 2	Top 2	390 2	250 2	210 2	Top 2	Bot 2	155	Top 2	370 ;	190 ;	Top :	415	360	300	250	210	160	Top	Top	Bot
26.08	25.88	27.09	26.16	27.40	26.56	26.07	26.55	28.68	27.80	27.26	29.39	29.16	29.00	27.06	26.92	28.26	26.54	27.73	26.70	26.36
6.20	5.99	7.25	5.89	6.67	5 · 88	5.88	6.33	7.06	7.55	7.09	7.78	8.88	7.43	6.82	7.03	6.65	6.67	6.13	5.07	6.49
1.06	1.08	1.06	1.08	1.07	1.07	1.07	1.08	1.05	1.08	1.07	0.97	0.97	0.90	0.89	0.90	0.89	0.95	0.97	1.08	→ • •
4.07	5,50	7.50	6.30	7.31	5.55	6.15	5.84	7.31	6.93	5.81	6.13	3.24	8.00	2.60	6.25	6.16	5.94	5.61	4.32	5.01
3.01	4.42	6.44	5.22	6.24	84.4	5.08	4.76	6.26	5.85	4.74	5.16	2.27	7.00	1.71	. 5.35	5.27	4.99	4,64	3.24	3.90
1.09	1.09	1.06	1.09	1.08	1.06	1.07	1.06	1.10	1.06	0.88	0.90	0.88	0.97	0.96	0.96	0.97	0.90	0.89	1.07	1.09
1.18	1.40	1.13	1.22	1.17	1.31	1.22	1.36	1.35	1.35	1.12	1.47	1.09	1.02	1.24	1.19	1.37	1.04	1.15	1.36	1.38
0.09	0.31	0.07	0.13	0.09	0.25	0.15	0.30	0.25	0.29	0.24	0.57	0.21	0.05	0.28	0.23	0.40	0.14	0.26	0.29	0.29
1.49	1.43	1.42	1.49	1,40	1.43	1.48	1.39	1.41	1.37	1.47	1.45	1.37	1.46	1.44	1.48	1.47	1.39	1.46	1.42	1.62
3.67	2.53	2.05	1.96	1.65	2.46	1.97	2,40	1.92	2.24	3. 11	3.19	5.51	1.74	4.43	2.46	2.00	2.14	2.56	2.61	3.28
2.18	1.10	0.63	0.47	0.25	1.03	0.49	1.01	0.51	0.87	1.64	1.74	4.14	0.28	2.99	0.98	0.53	0.75	1.10	1.19	1.66
0.90	0.14	0.10	0.06	0.07	0.11	0.14	0.24	0.02	0.53	0.45	0.30	2.24	0.08	1.83	0.45	0.43	0.78	0.11	0.35	0.62

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20	19	1 8	17	16	3	1,	<u>1</u>	12	<u> </u>	10	9	∞	7	6	Λ	£	ω	N	→	<u>2</u> 0
21.62	19.87	19.88	19.83	20.26	20.72	20.68	20.18	20.21	21.61	20.58	20.16	21.60	20.28	21.56	20.23	19.89	21.61	19.87	21.59	21.62
" 320	m 230	SG284 Top	" 303	II 200	SG285 Top	" 310	" 220	SG291 Top	" 320	n 205	SG282 Top	¹¹ 160	ıı 120	SG281 Top	" 380	" 250	SG232 Top	155	SG231 Top	SG230 220
0 28.40	0 25.30	p 23.38	3 26.28	0 25.94	p 28.25	0 26.84	0 25.26	p 25.89	0 25.25	5 24.19	p 27.79	0 28.08	0 28.79	p 29.27	0 28.30	0 27.17	p 28.79	5 27.76	p 30,29	0 28.14
6.77	5.42	3.49	44.9	5.67	7.52	6.15	5.07	5.67	3.64	3.60	7.62	6.48	8.50	7.70	8.46	7.28	7.18	7.89	8.69	6.51
0.91	0.91	0.91	0.99	0.98	0.97	0.91	0.91	0.97	0.94	0.97	0.91	0.90	0.97	0.92	0.93	0.97	0.97	0.95	0.97	1 08
5.31	3.60	2.82	2.77	3.91	7.50	4.87	1.73	5.97	1.93	2.91	7.74	5.52	7.77	7.39	3.04	2.50	6.68	5.88	8.15	6.24
04.4	2.69	1.91	1.76	2.93	6.53	3.96	0.82	5.00	0.99	1.94	6.83	4.62	6.80	6.47	N -1 -1	1.53	5.71	4.93	7.18	5.16
0.96	0.98	0.99	0.91	0.92	0.97	0.94	0.97	0.93	0.90	0.96	0.92	0.96	0.95	0.91	0.97	0.92	0.95	0.93	0.92	1.05
1.38	1.68	1.49	1.20	-1 -1 2	1.17	1.17	1.47	1.19	1.44	1.17	1.08	-1 -18	1.13	1.13	1.30	1.61	1. 13	1.06	N -1	1 12
0.42	0.70	0.50	0.29	0.20	0.20	0.23	0.50	0.26	0.544	0.21	0.16	0.22	0.18	0.22	0.33	0.69	0.18	0.13	0.20	0.07
1.44	1.42	1.39	1.45	1.39	1.38	1.43	1.38	1.37	1.46	1.38		1.47	1.46	1.40	1.48	1.46	1.38	1.49	1.47	
2.94	2.87	2.34	4.16	2.98	2.03	2.70	3.88	1.64	2.96	2.51	1.98	2.65	2.62	2.25	84.5	4.68	2.51	3.52	2.61	2.43
1.50	1.45	0.95	2.71	1.59	0.65	1.27	2.50	0.27	1.50	1.13	0.57	-1 -1 -0	1.16	0.85	4,00	3.22	1.13	2.03	1.14	1.02
144°O	0.57	0.12	1.67	0.94	0.12	0.68	1.24	0.13	0.59	0.03	0.05	0.45	0,35	0.14	1.61	1.83	0.14	0.78	0.15	0.25

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Λ	F	ļμ	Ŋ	=
20.23	19.89	21.61	19.87	21.59
3	=	=	=	SG293
Bot	150	100	80	Top
Bot 26.83	150 28.66	100 28.13	" 80 27 . 11	28.86
6.59	8.76	6.25	7.24	7.26
0.90	0.95		1.14 2.61	1. 13
0.90 4.71	0.95 7.87	2.88	2.61	1.13 7.41
3.81	6.92	1.74	1.47	6.28
0.96	1.10	1.07	1.15	• -1 -1
1.93	1.52	1.07 1.38	1.42	1 31
0.97	0.42	0.31	0.27	0.20
1.42	1.43	1.69	1.51	1.47
2.98	2.37	1.69 4.93	5.79	2,01
1.56	0.94	3.24	4.28	0.54
0.23		1,21		

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<u> </u>	18	17	16	15	14	13	12	1	10	9	œ	7	δ	J	ţ	ہب	Ŋ	-	20	19
19.87	19.88	19.83	20.27	20.73	20.63	20.18	20.22	21.61	20,25	20.16	21.55	20.28	21.57	20.24	19.89	21.61	19.87	21.60	21.63	19.87
=	=	SG238 I	:	=	***	SG237 1	==	acti ang	SG:133 1	=	=	=	SG131 1	=	=	112	SG127 1	=	=	SG128
4	20	Top	Bot	190	0	Top	Bot	90	Top	Bot	150	100	Top	10E	120	50	Top	Bot	250	150
25.53	25.73	26.55	26.42	27.00	26.24	26.74	26.57	27.20	27.22	26.84	28.19	27.23	27.42	26.54	25.60	25.98	25.79	27.58	27.27	24.20
5.66	5.85	6.72	6.15	6.27	5.61	6.56	6.35	5.59	6.97	6.68	6.64	6.95	5.85	6.30	5.71	4.37	5.92	5.98	5.64	4.33
	1.10	1.12	1.10	1.09	1.12	1.08	1.08	1.08	1.10	1.12	1 • 1 1	1.08	1.09	1.10	1.09	1.07	1.10	1.09	1.09	1.08
3.62	4.78	6.83	1.19	5.06	4.90	6.83	2.97	6.03	6.92	6.01	5.71	5.87	6.38	6.60	4.93	1.79	46.6	46.4	4.69	4.27
2.51	3.68	5.71	0.09	3.97	3.78	5.75	1.89	4.95	5.82	4.89	4.60	4.79	5.29	5.49	3.84	0.72	4.54	3.85	3.60	3.19
1. 13	1.15	1.14	1. 13	<u>-1</u> -1	1.08	-1 -1 -1	1.10	1.10	/ / /	1.12	 	1.12	1 1 1	1.12	1.10	1.10	1.09	1.12	- <u>'</u>	1.10
1.42	1. ⁴ 3	1.35	1.34	1.41	1.53	1.28	1.37	1.21	1.28	1.28	1.36	1.38	1.27	1.27	1.43	1,29	1 • +1	1.56	1.70	1,42
0.29	0.28	0.21	0.21	0.30	44.0	0.17	0.27	0.11	0.17	0.16	0.25	0.26	0.16	0.15	0.33	0.19	0.32	0.44	0.59	0.32
1.67	1.66	1.52	1.49	1.46	1.46	1.47	1.51	1.47	1.52	1.46	1.42	1.53	1.50	1.40	1.51	1.41	1.45	1.51	1.50	1.45
3.51	3.03	2.21	÷.83	2.67	2,40	2.03	4.38	1.93	2.39	1,88	2.25	2,63	1.89	1.92	2.68	3.45	2.40	2.65	2,47	2.12
1.84	1.37	0.69	3.34	1.21	0.94	0.55	2.87	0.46	0.87	0.42	0.83	1.10	0.39	0.52	1.17	2.04	0.95	1.14	0.97	0.67
1.02	0.51	0.11	2.51	0.79	0.44	0.08	1.32	0.07	0.11	1.21	0.96	0.80	0.00	0.13	0.37	1.42	0.11	0.55	0.48	0.15

20	19	18	17	16	15	14	13	12	1	10	9	∞	7	σ	પ	+	ىبا	Ŋ	ⅎ	20
21.63	19.87	19.88	19.83	20.27	20.73	20.63	20.18	20.22	21.61	20.25	20.16	21.55	20.28	21.57	20,24	19.89	21.61	19.87	21,60	21.63
SG125	2	=	=	SG292	=	=	=	SG240	=	=	s c 239.	=	=	vice) iong	SG 241	7	=	11	=	SG238
Top	310	250	200	Top	Bot	100	70	Top	370	250	Top	Bot	200	150	Top	190	130	195	90	70
28.35	26.97	25.39	25.68	25.58	25.60	26.48	25.80	27.24	27.22	29.95	25.58	27.44	26.50	26.57	25.60	26.18	27.41	27.13	28.85	27.85
6.72	7.10	5.51	5. 85	5.31	4.87	5.85	5.62	7.02	5.61	9.70	5.42	5.89	6.22	5.00	5.36	6.29	5.80	7.26	7.25	6.22
1.14	1 1 1 +	1.16	1.15	1.16	1.17	1.14	1.18	1.15	1.15	1.10	1.14	1.16	1.17	1.17	1.19	1.15	1.15	1.15	1.15	1.10
6.14	2.04	4.81	4.39	5.39	5.06	5.67	4.05	7,19	5.81	5.79	5.61	+0.4	5.85	3.55	6.04	3.15	4.03	7.35	6.97	5.23
5.00	0.90	3.65	3.24	4.23	3.89	4.53	2.87	6.04	4.66	4.69	4.47	2.88	4.68	2.38	4.85	2,00	2.88	6.20	5.82	÷ 13
1.15	1.15	1.10	1.14	1.13	1.15	1.1	1.16	1.10	1.18	1.18	1.15	1.16	1.18	1.17	1.17	1.10	1.16	1.09	1.14	-7 • -7
1 + 1	1.33	1.61	1.52	1.53	1.27	1.35	1.84	1.35	1.29	1.64	1.35	1.50	1.63	1.47	1.33	1.24	1.54	1.23	1.35	1.44
0.26	0.18	0.51	0.38	0.40	0.12	0.21	0.68	0.25	0.11	0.46	0,20	0.34	0.45	0.30	0.16	0.14	0.38	0.14	0.21	0.33
1 • ¥ 1	1.62	1.54	1.54	1.60	1.47	1.61	1.49	1.65	1.53	1.66	1.49	1.65	1.54	1.64	1.55	1.61	1.50	1.63	1.68	1.66
2 28	5.09	2.27	2.59	1.96	1.87	2.29	2.67	2.02	1,40	4.83	1.94	3.12	2.12	3.19	1.79	3.48	2.91	2.17	2.48	2.75
0.87	3.47	0.73	1.04	0.36	0.40	0.68	1.18	0.37	0.37	3.17	0.45	1.47	0.58	1.55	0.24	1.87	1.41	0.54	0.80	1.09
0.59	2.55	0.62	1.19	0.32	0.46	0.43	0.88	0.36	0.47	1.38	0.30	1.20	0.50	0.77	0.11	2.27	1.13	0.38	0.42	0.67

20	19	18	17	16	15	7	<u>.</u>	12	<u>¬</u>	10	9	<u></u>	7	6	VI	ţ	W	N	-7	
19.94	19.87	19.88	19.83	20.27	20.73	20.54	20.18	20.22	21.61	20.19	19.76	21.08	20.28	21.03	20.24	1989	21.61	19.87	21.55	
Ξ	SG 85	 		Ξ	SG 86	-st -m2	==	=		SG 87	geria meng		7	11	SG 124		:::	7	SG 125	
90	αoΤ	2,45	220	190	Top	tog.	350	250	150	Top	ti ot	210	140	90	Top	Bot	320	230	170	
26.59	26,90	25,80	26.07	25,16	28,61	26.22	27.06	25.09	26.05	26.34	27.71	26.72	28.73	27.11	27.19	25.05	27.80	26.39	27.39	
65	7.03	5,92	6.24	4.89	7.88	5.68	6.88	4.87	4.44	6.15	7.95	5.64	£,	6.08	6.95	5.16	6.19	6.52	5.8+	
<u>-</u> -	1 • 14	1,14	1.12	1.14	1.16	1.17	1.15	1.15		1.13	1 1 1	1. 13	1.16	1.13	1.15	1.15	1.15	1.17	1 • 14	
6 . 88	7.10	5.78	4.27	4.47	40.8	6.18	6.27	4.78	4.07	6.39	4.98	4.69	7.95	5.28	7.18	1.43	2.02	++++	4.95	
5.75	5.96	49.4	3.15	3. 3. 3.	6 88	5.01	5.12	3.63	2.96	5.26	3.84	3.56	6.79	÷.15	6.03	0.28	0.87	3.27	3.81	
1 1 0	√ 1 -1	 	1 1 3	1.12	1.13	1.13	1.10	1.16	1.13	1.19	1.16	1.16	1 1 W	1.13	1.15	1.15	1.16	1.15	1.12	
ω ω	1.37	1.36	1.49	1.46	1.51	1,40	1.42	1.37	1.40	1.47	1.62	1.55 55	1.38	1.54	1.37	1.58	1.71	1.95	1.37	
0.20	0,25	0.25	0.36	0.34	0.38	0.27	0.32	0.21	0.27	0.28	0.46	0.39	0.25	0.41	0.22	0.43	0.55	0.80	0.25	
1 5 2	1.58	1.48	1.58	1.49	1.57	1.48	1.59	1.48	1.59	1.50	1.56	1.47	1.59	1.48	1.53	1.53	1.60	1.53	1.63	
N 10	2.20	2.26	3.36	2,34	205	1.69	2.52	2.51	2.52	1.87	ω • 83	2.83	2.59	2.57	2.13	4.09	5.19	ω 11	2.84	
0.58	0.62	0.78	1.78	0.85	0.48	0.21	0.93	1.03	0.93	0.47	2.27	1.36	1.00	1.09	0.60	2.56	3.59	1.58	1.21	
0.12	0.20	0.25	0.94	0.37		0.19	0.51	0.00	0.28	0.14	1.38	0.33	0+0	0.43	0.10	1.89	1.18	0.87	0.57	4

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20	19	18	17	16	75	14	<u>J</u>	7	<u> </u>	10	9	CΩ	7	σ	J	F	W	N	_7	
20,61	19.87	19.88	19.83	20,20	20.73	21.61	20.18	20.22	21.61	19.85	20.17	20.59	20.28	19.84	20.19	19.89	21.61	19.87	21.55	
= 201	15	" 101	SG123 Top		SG122 To	п п	28	n 16	SG121 To	p.	= 27	= 2	SG120 To	п В	16	SG 82 To	n B0	·· 25	SG 85 18	
25.31	0 27.06	01 25.35	op 25.65	00 24.78	op 28,21	ot 25.90	80 25.73	60 25.58	op 28.45	ot 26.43	270 28.85	10 27.41	op 25.99	ot 28.07	60 27.21	op 23.93	ot 26.64	0 25.82	30 27.31	
4.70	7.19	5.47	5.82	4.58	7.48	4.29	5.55	5.36	48.0	6.57	8.68	6.82	5.71	8.23	7.02	+0.0+	5.03	5.95	5.76	
_7 _7 ▶	1.14	1 1	1.06	1.02	1.10	1.02	1.12	1 1 1	1.07	1.06	1.01	1.09	1.09	1.10	1.07	1.09	1.09	1.09	1 • 1 1	
4.50	5.92	1+.1+1	5.50	4.50	8.11	4.76	5.61	5.13	5,12	4.65	6.61	7.49	2.89	7.60	5.99	2.46	+.83	٠٠ - 10	2.72	
3.39	4.78	3.30	44.4	3.48	7.01	3.74	64.4	4.02	4.35	3.59	5.60	6.40	1.80	6.50	4.92	1.37	3.74	2.95	1 0 1	
1.09	1 0	1.09	1.00	_7 _7 •	1.05	-7 -7 -7	1.06	1.07	1.01	1.05	1.09	1.07	1.08	1.06	1.06	1.07	1.07	1.06	1.08	
1.35	1.57	1.52	1.34	1.36	1.23	1 20	1.22	1.37	1.35	1.45	1.45	1 1 8	1.93	1.21	1.30	1.53	1.31	1.56	1.38	
0.26	0,47	0.43	0.34	0.25	0.18	0.09	0.16	0.30	0.34	0+40	0.36	0.11	0.85	0.15	0.24	0.46	0.24	0.50	0.30	
1.14	1 55	1.43	1 • 5 8	1.42	1.58	1.45	1.57	1.45	1.51	1.49	1.59	1.44	1.56	1.46	1.56	1.4.1	1.58	1.45	1.63	
2 10	2,89	2.95	2.37	2.09	1.85	1.75	2.09	2.17	3.22	3.01	3.65	1.75	3.85	2.44	3.04	ω • 11	2.33	3.15	4.32	
0,66	1.31+	1.52	0.79	0.67	0.27	0.30	0.52	0.72	1.71	1.52	2.06	0.31	2.29	0.98	1.48	1.70	0.75	1.70	2.69	
0.39	0.60	0.22	0.25	0.18	0.02	0.15	0.38	0.32	0.43	1.06	0.66	0.00	0.77	0.60	0.38	0.51	0.30	0.80	1.16	1

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20	19	<u>~</u>	17	0\ 1	<u>\7</u>	1,	(J)	7	>	10	9	∞	7	σ	J	‡	W	N	
20,61	19.87	19.88	19.83	20.20	20.73	21.61	20.18	20.22	21.61	19,85	20.17	20.59	20.28	19.84	20,19	19.89	21.61	19.87	21.55
-24 -29	=	==	SG 84	=	===	=	SG 83	=	=	=	=	17	SG 81	=	=	11	=	=	SG126
Bot	120	95	Top	70H	280	210	qoT	Bot	350	310	260	190	Top	340	290	250	190	95	Top
28,97	25.72	26.29	26.44	25.45	25.23	28.09	24.73	26,68	28.42	25.89	25.61	25.14	25.54	26,28	26,19	24.76	28,20	25.81	27.89
8.36	5. 85	0,+1	6.61	5,25	4.50	84.0	1.55	6.46	6.81	6.03	5,44	7.55	5.26	6.44	6.00	4.87	6.59	5.94	6.34
1,17	1.15	1. 15	1.14	1.16	1 1 3	1.13	1.15	1.13	1 1 12	1.12	1.09	1 1 +	1.12	1.12	-7 -7 •	1.1	1.13	1.17	1.16
+1+	5.74	5 28	6.13	5.25	4.62	5.77	3.46	5.46	6.19	5.48	4.68	4.30	4.03	5.53	6.81	1.21	7.27	1.84	6.34
7.27	4.59	4.13	4.99	4.09	3.49	49.4	2.31	¥•33	5.07	4.36	3.59	3.16	2.91	+ + 1	5.70	0.07	6.14	0.67	5.18
1,12	1.12	1.12	1.14		1.13	1.15	1.12	1.14	1.13	1.13	1 • 1 }	-1 -1 -1	1.12	1.12	1.10	1.12	1 • 1 +	1 - 1 +	-7 • -7
1.39	1,43	1.40	1.36	1.33	1.30	1.27	1.76	1 (U 1	1.43	1.35	1.30	1.23	1.54	1.29	-7 -7 -7	1•13	1.16	1.44	1.30
0.27	0.31	0.28	0 N 0	0.22	0.17	0,12	0.64	0.17	0.30	0.22	0.16	0.12	0.42	0.17	0.01	0.01	0.02	0.30	0.19
1 + 5	1,62	1.50	1.65	1.47	1.64	1.46	1.65	1.47	1.66	1.47	1.66	1.44	1.69	1.43	1.69	1.43	1.70	1.48	1.61
1.92	2.15	2.74	2.53	1 88	2.01	1.97	2.76	2.75	2.49	1.99	2.30	2 • 18	3.00	2.51	1.71	3.15	1.92	7.47	2.23
0.47	0,53	1.24	0.88	0.41	0.37	0.51	-7 •	1.28	0.83	0,52	0.64	0.74	1.31	1.08	0.02	1.72	0.22	2.99	0.62
0	0,42	0.76	0.52	0.53	74.0	-1 (N) -1	0.49	0.68	0.61	0.93	1.05	0.53	0.62	0.78	0.27	3.07	0.21	1.97	0.35

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	20	19	<u>~</u>	17	<u>o</u> ,	15	7	13	7	<u> </u>	10	Ω,	Φ	7	0	VΊ	ť	س)	N	_7
	20.61	19.87	19.88	19.83	20,20	20.73	21.61	20,18	20.22	21,61	19.85	20.17	20.59	20.28	19.01	20.19	19.89	21.61	19.87	21 • 55 5
	:: ti ot	SG114 Top	" Bot	180	" 150	SG115 Top	SG116 Top	" Bot	335	" 325	SG117 Top	n 290	u top	SG118 260	tot:	n 200	n 170	·· 150	8G119 Top	SG123 Bot
	27.14	25.09	25.34	25.30	25.53	26.99	26,26	27.41	28.74	30.55	24.79	29.64	27.46	28.39	27.41	27.90	26.00	29,28	25.22	27.10
	6.53	5.22	2.16	5.47	5.33	6.26	4.65	7.23	00 57 22	46.8	4.93	9.47	6.87	∞ • 1	7.57	7.71	6.11	7.67	5.35	5.55
	-1 -1 -2		1 13	1.16	1.13	1.15	1.15	1.16	1.16	0, -7	1.1+	1.13	1.12	1.13	1.16	-7 • -7	1 1 1	1.12	1 • 1 +	1 1 2
	57 9 W	J. 33	2,44	5.97	4.34	44.9	2.67	7.29	5.79	8 20	3.99	9.06	3,29	7.43	6.31	4.77	5.94	7.75	3.24	6.24
	F. 81	4.22	1 W 1	÷.81	ω 21	5.29	1.52	6.13	4.63	7.04	2.85	7.93	2.17	6.30	5.15	3.66	4.80	6.63	2.10	5.12
	1 -1 +	1 • 1 3	1.16	1.16	-1 -1 -1	1.12	1 • 12	1.09	¬ ¬ •	1,13	1 W	1.17	1.14	1 • 12	1 • 1 +	1.16	1.15	1.17	1.13	
	 N 1	1.29	1 12 12	1.39	1	1.29	-1 0 -1	1.16	1.36	1 21	-1 +-1	1 U U	1.86	1.32	1.54	1.58	1.43	1.29	1.80	1 • 19
	0.07	0.16	0,06	0.23	0.20	0.17	0.49	0.07	0.25	0.08	0.28	0.18	0.72	0.20	0.40	0.42	0 28	0.12	0.67	0,06
	1.59	1.64	1.43	1,67	1.44	1.62	1.45	~ 0 ~	1.63	1.49	1.46	1.64	1.50	1,67	1.46	1.63	1.43	1.64	1.49	1.67
	2.03	2.14	2.58	1.86	2	2.00	3.24	2.25	3.63	2.76	2.67	2.51	¥•19	2.62	2.30	3.33	2.09	2.32	3.29	1.84
` `	0.44	0.50	1.15	0.19	0.87	0.38	1.79	0.64	2.00	1.27	1.21	0.87	2.69	0.95	0.84	1.70	0.66	0.68	1.80	0.17
) S +	- 12 1	0.34	2.94	0.24	1.05	0. 1.	0.85	0.39	1.0+	0.55	0.59	0.49	1.29	0.66	1 1 8	1.92	0.37	0.24	0.78	0.20

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つ , , , , ,	0, 75	<u>,</u> 20 20	1 . л 0	0.06	1.20	1.14	フ, , , , , , , ,	8.60	1.08	ω Σ	20 20 20 20 20 20 20 20 20 20 20 20 20 2	SG162 Ton	νο ₋ νο	7
0,40	0,71	2.17	1.46	0.23	1.35	1.12	5.67	6.73	1.06	7.01	27.74	SG161 Top	20.73	15
0.61	1.84	ω • ω 1	1.47	0.42	1.57	1.15	3.75	÷.85	1 1 0	6.62	28.23	SG160 Top	21.61	14
1.51:	1.97	3.46	1.49	0.11	1.27	-1 -16	5.97	7.05	1.08	9.59	29.77	" Bot	20.18	-1 ω
0.91	2.19	3.70	1.51	0.61	1.75	1 1	1.86	2.92	1.06	5.57	25.79	SG159 Top	20.22	12
0.25	0.43	1.93	1.50	0.22	1.36	1.1+	6.37	7.46	1.09	7.27	28	SG158 Top	21.61	_7 _7
0.14	0.17	1.66	1.49	0.03	1 18	1 . 15	84.8	9.56	1.08	8 8 8 8	28.68	·· 150	19.85	10
0.59	0.84	2.34	1.50	0.54	1.69	1.15	5.30	6.39	1.09	7.27	27.44	8G157 Top	20.17	9
3.36	1.17	2.69	1.52	0.22	1.32	1.10	1.17	2.28		5.92	26.51	112	20.59	œ
0.23	0.35	1 • 88	1.53	0.08	1.16	1.08	6.80	7.94	1 1 1 +	7.46	27.74	SG156 Top	20.28	7
1.20	0.25	1.89	1.63	0.12	1 .21	1.09	5.27	6.39	1.12	6.85	26.69	56155 Top	19.84	6
0.87	0.34	1.84	1.50	0.22	1.32	1.10	7.83	8.93	1.10	9.26	29.45	" Bot	20.19	<i>ত</i> া
0.21	0.41	1.91	1.50	0.03	1.15	1.12	6.60	7.71		7.25	27.14	SG154 Top	19.89	+
0.19	0.27	1.83	1.56	0.14	1.22	1.08	6.72	7.82	1.10	7.32	28.93	SG152 Top	21.61	W
1.06	2.32	3.83	1.51	0.24	1.29	1.05	1.93	3.09	1.16	5.55	25.42	" Bot	19.87	2
0.25	0.43	1.97	1.54	0.44	1.56	1 1 2	5.23	6.32	1.09	6.35	27.90	SG151 Top	21.55	_7
		F.Silt			C.Silt			Sand		Wt.	&Samp Wt	No	W Wt.	No.

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17	16	15	14	13	12	<u> </u>	10	9	œ	7	6	J	t	ω	N	_7	20	19	
19.83	20,20	20.73	21.61	20,18	20.22	21.61	19.85	20.17	20.59	20.28	19.84	20.19	19.89	21.61	19.87	21.55	20.61	19.87	
SG 66 Top	SG 65 Top	SG 63 Top	SG 62 Top	SG 61 Top	SG 58 Top	" Bot	SG 57 Top	SG 56 Top	" Bot	SG 55 Top	™ Bot	n 180	SG 54 Top	SG165 Top	" Bot	SG164 Top	" Bot	SG163 80	
26.48	27.61	25.81	28,08	26.15	25.76	27.62	25.47	25.67	28.51	25.57	25.26	26.48	25.65	27.26	24.80	28.07	29.15	28.00	
6.65	7.41	5.08	6.47	5.97	5.5.4	6.01	5.62	5.50	7.92	5.29	5.42	6.29	5.76	5.65	4.93	6.52	8,57	8,13	
1 • 1 +	1 1 1	1.12	1.14	1.12	1.16	1 • 14	1.15	1.15	1.10	1.13	1.13	1.16	1.17	1.16	1 • 1 +	1 14	1. 13	1.16	
5.77	7.76	1.96	3.50	94.4	5.67	2,13	4.92	4.97	+.81	46.4	3.02	1.1	5.81	5.92	5.13	5.92	9.27	5.69	
4.63	6.65	0.84	2.36	3•34	÷ 51	0.99	3.77	3.82	3.71	3.81	1.89	3.25	49.4	4.76	3.99	4.78	8.14	4.53	
1.12		1 1 3	1,12	1.10	- - - -	1.13	1.12	1.10	1.13	1.10	1.10	1.10			1.08	1.08	1.08	1.06	
1.27	1,23	1.99	1.72	1.47	1.24	1.39	1.42	1.34.	1.87	1.38	1.39	1.47	1 • + 1	1.27	1.17	1.25	1.07	1.37	
0.15	0 1 2	0.86	0.60	0.37	0.13	0.26	0.30	0.24	0.74	0.28	0.29	0.37	0.30	0.16	0.09	0.17	0.01	0.31	
1.64	1.42	1.62	1.49	1.65	1.44	1•41	1.64	1.42	1.64	1.43	1.62	-1 - -	1.68	1.44	1.43	1.67	1.53	1.51	
2.38	1.70	3.45	3.70	2.87	1.78	2.95	2.47	2.19	2.82	2.17	2.96	2.39	2.08	1.75	1.52	2.49	1.66	2.30	
0.74	0.28	1.83	2 2 1	1.22	0.34	1.54	0.83	0.77	1.18	0.74	1.34	0.95	0.40	0.31	0.09	0.82	0.13	0.79	
1.13	0.36	1.55	1.30	1.04	0.56	ω. 2	0.72	0.67	2.28	0.45	1.90	1.72	0.42	0.42	0.76	0.74	0.26	2.50	2

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APPENDIX 2. C. CaCo PERCENTAGES. WHYALLA NORTH.

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Beaker No.	Beaker Wt.	Sam	nple No.		Beaker & Samp	Sample Wt.	Wt. After Digestion	Wt.Loss	CaCo ₃ % Wt.Loss /Wt x100
	20 0).	- CC	001	Top	วว กรี	1.90	32•755	n GRAMS 1 . 195	62.89%
21	32.04	υα			33.95			0.845	43.22%
22	32.45		11	Bot -	34.405	1.955	33.56	•	
23	32.01	SG		Top	33.76	1.75	32.29	1.47	84.00%
24	32.25		tt	130	34.09	1.84	32.61	1.48	80.43%
25	31.48		11	140	33.29	1.81	32.36	0.93	51.38%
26	32.76		11	Bot	34.72	1.96	33.33	1.39	70.91%
27	32.11	SG	223	Top	33.875	1.765	32.60	1.275	72.23%
28	32.63		11	130	34.67	2.04	34.29	0.38	18.62%
29	32.02		11	180	34.00	1.98	33.16	0.84	42.42%
30	32.35		11	222	34.30	1.95	34.23	0.07	03.58%
31	32.26		tt	270	34.21	1.95	34.17	0.04	02.05%
32	32.615	SG	224	Top	34.625	2.01	33.07	1.555	77.36%
33	30.90		11	150	32.97	2.07	31.93	1.04	50.24%
34	31.935		11	195	33.94	2.005	32.81	1.13	56.35%
35	31.90		tt	240	33.86	1.96	32.83	1.03	52.55%
36	32.365		11	300	34.33	1.965	34.21	0.12	06.10%
37	32.26		11	315	34.115	1.855	33.10	1.015	54.71%
38	32.59		11	330	34.69	2.10	34.40	0.29	13.30%
39	31.64		11	360	33.52	1.88	33 • 355	0.165	08.77%
1 ₊ O	31.89		11	400	33.905	2.015	33.67	0.235	11.66%
69	32.17	SG	225	Top	34.11	1.94	32.69	1.42	73.19%
70	30.84		11	160	32.905	2.065	32.20	0.705	34.14%
71	32.39		11	210	34.66	2.27	34.48	0.18	07.92%
72	31.865		11	250	33.96	2.095	32.79	1.17	55.84%
73	31.91		11	300	33.72	1.81	32.46	1.26	69.61%
73 74	32.39		11	360	34.73	2.34	34.41	0.32	13.67%
7 - 75	30.68		11	415	32.76	2.08	31.65	1.11	53.36%
()	JU•00			' ' ' /	JL • / O	_, 00	J.••/	. •	/ J ♥ J → I*

76	31.81	SG	226	Top	33.93	2.12	32.50	1.43	67.45%
77	31.95		11	140	34.34	2.35	33.265	1.075	42.55%
78	31.355		11	370	33.46	2.105	33.07	0.39	18.52%
79	31.72	SG	227	Top	33.90	2.18	32.32	1.58	72.47%
80	31.23		11	155	33.25	2.02	32.28	0.97	48.01%
l+W	29.66		11	Bot	32.07	2.41	30.78	1.29	53.52%
107	31.83	SG	228	T_{op}	34.24	2.41	32.44	1.80	74.68%
108	29.75		11	210	32.41	2.66	32.16	0.25	09.39%
3 T	29.875		11	250	32.40	2.525	31.855	0.545	21.58%
103	32.58		11	390	34.92	2.34	34.73	0.19	08.11%
(6M) 5M	29.79	SG	230	Top	31.88	2.09	30.155	1.725	82.53%
3B	30.28		11	170	32.51	2.23	31.34	1.17	52.46%
5B	30.63		tt	220	32.85	2.22	32.50	0.35	15.76%
11	31.07	SG	231	Top	33.00	1.93	31.77	1.23	63.73%
12	32.30		11	155	34.38	2.08	33.05	1.33	63.94%
13	32.24	SG	232	Top	34.24	2,00	32.72	1.52	76.00%
14	32.32		11	250	34.32	2.00	32.575	1.745	87.25%
15	32.015		11	380	33.94	1.925	32.62	1.32	68.57%
16	32.17	SG	281	Top	34.05	1.88	32.61	1.44	76.59%
17	31.60		Ħ	120	33.54	1.94	32.54	1.00	51.54%
18.	31.16		17	160	32.95	1.79	31.58	1.37	76.53%
19	32.02	SG	282	Top	33.95	1.93	32.495	1.455	75.38%
20	32.06		11	205	33.88	1.82	32.60	1.28	70.32%
41	30.04		11	320	31.84	1.80	30.555	1.285	71.38%
42	30.45	SG	291	Top	32.38	1.93	30.62	1.76	91.19%
43	29.94		11	220	31.85	1.91	31.05	0.80	41.88%
7+7+	30.02		11	310	31.755	1.735	31.515	0.24	13.83%
45	30.85	SG	284	Top	32.76	1.91	31.12	1.64	85.86%
46	29.50		11	230	31.435	1.935	29.59	1.845	95.34%
47	30.07		11	320	31.99	1.92	31.03	0.96	50.00%

	48	30.08	Sg	285	Top	31.995	1.915	30.505	1.49	77.80%
	49	30.43		11	200	32.525	2.095	32.30	0.225	10.73%
	50	29.65		11	305	31.66	2.01	31.45	0.21	10.44%
	6T	30.715	SG	293	Top	32.88	2.165	30.93	1.95	90.06%
	5 T	29.10	•	11	80	31.09	1.99	30.56	0.53	26.63%
	7B	30.14		11	100	32.25	2.11	31.39	0.86	40.75%
(6B 102)	31.74		11	150	34.12	2.38	32.545	1.575	66.17%
	7T 101)	28.77		11	Bot	31.06	2.29	29.46	1.60	69.86%

Beaker No.	Beaker Wt.	Sar	mple No	(m)	Beaker & Samp	Sample Wt.	Wt. After Digestion	Wt. Loss	CaCo ₃ % Wt Loss/ Wt x100
1B(104)	30.15	SG	132	Top	31.85	1,70	30.27	1.58	92.94%
1M(110)	32.04		11	140	33.80	1.76	32.96	0.84	47.72%
7T(101)	28.77		11	170	30.82	2,05	29.51	1.31	63.90%
4T(7)	30.74		11	210	33.42	2.68	31.49	1.93	72.01%
4B(9)	29.912		11	280	32.07	2.158	30.49	1.58	73.21%
1 T(105)	31.88		11	310	34.04	2.16	33.02	1.02	47.22%
6B(102)	31.74		11	350	33.75	2.01	32.88	0.87	43.28%
7M(106)	29.58	SG	130	Top	31.38	1.80	29.80	1.58	87.77%
3M(106)	30.624		11	50	32.97	2.346	31.31	1.66	70.75%
(6) 7B(109)	30.14		11	133	31.75	1.61	30.74	1.01	62.73%
4M(8)	29.66		n	175	31.81	2.15	30.45	1.36	63.25%
108	29.75	SG	129	Top	31.79	2.04	30.03	1.76	86.27%
6M(103)	32.58		11	190	34.45	1.87	33.09	1.36	72.72%
3B (2)	30.28		11	250	32.31	2.03	31.04	1.27	62.56%
6T (5)	30.715		11	310	34.13	3.415	32.60	1.53	44.80%
107	31.83		tt	350	35.43	3.60	33.89	1.54	44.77%
3T (1)	29.875		11	390	33.06	3.185	32.14	0.92	28.88%
5M (4)	29.79	SG	128	Top	32.28	2.49	30.45	1.83	73.49%
5B(10)	30.63		tt	150	32.54	1.91	30.91	1.63	85.34%
5T(3)	29.10		11	250	31.63	2.53	29.55	2.08	82.21%
77	31.95		11	Bot	34.63	2.68	33.44	1.19	44.40%
75	30.68	SG	127	Top	33.19	2.51	31.32	1.87	74.50%
78	31.355		11	50	33.09	1.735	32.08	1.01	58.21%
36	32.365		11	120	33.93	1.565	32.53	1.40	89.45%
76	31.81		11	Bot	33.72	1.91	32.51	1.21	63.35%
80	31.23	SG	131	Top	33.82	2.59	31.97	1.85	71.42%
38	32.59		11	100	34.82	2.23	34.12	0.70	31.39%
37	32.26		n	150	35.91	3.65	35.03	0.88	24.10%
31	32.26		11	Bot	34.71	2.45	34.59	0.12	04.89% · / 5

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73	31.91	SG	133	Тор	33.64	1.73	32.17	1.47	84.97%
50	29.65		11	90	33.23	3.58	32.67	0.56	15.64%
49	30.43		11	Bot	33•33	2.90	31.88	1.45	50.00%
48	30.08	SG	237	Top	32.29	2.21	30.68	1.61	72.85%
47	30.07		11	40	32.37	2.30	30.72	1.65	71.73%
46	29.50		11	190	31.6 ¹ 4	2.14	30.70	0.94	43.92%
45	30.85		11	Bot	33.77	2.92	33.38	0.39	13.35%
1414.	30.02	SG	238	Top	32.16	2.14	30.86	1.30	60.74%
43	29.94		tt	20	32.03	2.09	30.83	1.20	57.41%
42	30.45		tt	54	32.36	1.91	31.35	1.01	52.87%
41	30.04		11	70	32.40	2.36	31.12	1.28	54.23%
11	31.07		tt	90	34.06	2.99	32.49	1.57	52.50%
12	32.30		11	105	3 ¹ +•92	2.62	33.81	1.11	42.36%
13	32.24		11	130	35.32	3.08	34.03	1.29	41.88%
14	32.32		11	190	34.60	2.28	34.23	0.37	16.22%
15	32.015	SG	241	Top	33.83	1.815	32.31	1.52	83.74%
16	32.17		11	150	34.18	2.01	32.99	1.19	59.20%
17	31.60		11	200	33.61	2.01	32.24	1.37	68.15%
18	31.16		11	Bot	33.75	2.59	33.11	0.64	24.71%
19	32.02	SG	239	Top	33.66	1.64	32.35	1.31	79.87%
20	32.06		11	250	34.67	2.61	33.49	1.18	45.21%
21	32.04		11	370	34.60	2.56	33.04	1.56	60.93%
22	32.45	SG	240	Top	34.48	2.03	32.96	1.52	74.87%
23	32.01		tt	70	34.67	2.66	32.80	1.87	70.30%
24	32.25	SG	240	100	35.17	2.92	33.34	1.83	62.67%
25	31.48		11	Bot	34.77	3.29	32.90	1.87	56.83%
26	32.76	SG	292	Top	34.42	1.66	32.98	1.44	86.74%
27	32.11		11	200	34.48	2.37	33.29	1.19	50.21%
28	32.63		11	250	34.56	1.93	33.15	1.41	73.05%
29	32.02		11	310	34.45	2.43	34.05	0.40	16.46%
30	32.35	SG	125	Top	34.54	2.19	33.03	1.51	68.95%
1B	30.15		11	170	32.45	2.30	31.07	1.38	60.00%

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1 T	30.74	SG	125	230	32.66	1.92	31.38	1.28	66.66%
6B	31.74		11	320	33.66	1.92	32.12	1.54	80.20%
1T	31.88		11	Bot	33.40	1.52	32.375	1.025	67.43%
7B	30.14	SG	124	Top	32.52	2.38	30.58	1.94	81.51%
7M	29.58		11	90	31.76	2.18	30.215	1.545	70.87%
1M	32.04		11	140	34.26	2.22	33.07	1.19	53.60%
7T	28.77		11	210	31.07	2.30	30.43	0.64	27.82%
4В	29.91		11	Bot	32.51	2.60	32.03	0.48	18.46%
3M	30.62	SG	87	Top	32.82	2.20	30.81	2.01	91.36%
21	32.04		11	150	34.68	2.64	32.68	2.00	75 • 75%
22	32.45		11	250	34.38	1.93	32.98	1.40	72.53%
23	32.01		11	350	34.57	2.56	33.29	1.28	50.00%
24	32.25		tt	Bot	35.23	2.98	33.86	1.37	45.97%
25	31.48	SG	86	Top	34.42	2.94	31.83	2.59	88.09%
26	32.76		11	190	34.555	1.795	33.26	1.295	72.14%
27	32.11		n	220	34.18	2.07	32.81	1.37	66.18%
28	32.63		H	245	36.14	3.51	33.79	2.35	66.95%
29	32.02	SG	85	Top	33.88	1.86	32.24	1.64	88.17%
30	32.35		11	90	35 •63	3.28	35.06	0.57	17.37%
41	30.04		11	180	32.32	2.28	31.215	1.105	48.46%
42	30.45		H	250	32.34	1.89	30.865	1.475	78.04%
43	29.94		11	Bot	31.79	1.85	30.45	1.34	72.43%
7+7+	30.02	SG	82	Top	32.535	2.515	30.49	2.045	81.31%
45	30.85		11	160	33.44	2.59	32.19	1.25	48.26%
46	29.50		11	Bot	32.11	2.61	31.72	0.39	14.94%
47	30.07	SG	120	Top	33.37	3.30	30.765	2.605	78.94%
48	30.08		11	210	32.82	2.74	31.88	0.94	34.30%
49	30.43		11	270	33.32	2.89	31.735	1.585	54.84%
50	29.65		11	Bot	32.58	2.93	30.44	2.14	73.03%
11	31.07	SG	121	Top	33.92	2.85	31.45	2.47	86.66%
12	32.30		11	160	35 • 305	3.005	33.23	2.075	69.05%
13	32.24		11	280	34.38	2.14	33.29	1.09	50.93%
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14	32.32	SG	121	Bot	34.57	2.25	33.45	1.12	49.77%
15	32.015	SG	122	Top	34.67	2.655	32.47	2.20	82.86%
16	32.17		11	100	34.74	2.57	32.93	1.81	70.43%
17	31.60	SG	123	Top	33.60	2.00	31.98	1.62	81.00%
18	31.16		11	101	33.92	2.76	32.06	1.86	67.39%
19	32.02		11	150	33.98	1.96	32.82	1.16	59.18%
20	32.06		II.	201	34.28	2.22	33.27	1.01	45.49%
3 T	29.875	SG	126	Top	32.74	2.865	30.23	2.51	87.60%
3B	30,28		11	95	32.82	2.54	30.77	2.05	80.70%
5 T	29.10		11	190	32.28	3.18	30.15	2.13	66.98%
5M	29.79		11	250	33.105	3.315	30.70	2.405	72.55%
6T	30.715		11	290	33.715	3.00	32,675	1.04	34.66%
108	29.75		17	340	32.11	2.36	31.73	0.38	16.10%
6M	32.58	SG	81	Top	34.46	1.88	32.91	1.55	82.44%
ЧM	29.66		11	190	34.57	4.91	32.80	1.77	36.05%
107	31.83		11	260	34.56	2.73	32.935	1.625	59.52%
5B	30.63		11	310	34.82	4.19	33.49	1.33	31.74%
31	32.26		11	350	33.98	1.72	32.64	1.34	77.90%
32	32.615		11	Bot	36.02	3.405	34.05	1.97	57.85%
33	30.90	SG	83	Top	32.92	2,02	31.30	1.62	80.20%
34	31.935		11	210	34.10	2.165	32.82	1.28	59.12%
35	31.90		11	280	33.93	2.03	32.24	1.69	83.25%
36	32.365		11	Bot	34.87	2.505	33.535	1.335	53.29%
37	32.26	SG	84	Top	34.88	2.62	32.68	2.20	83.97%
38	32.59		11	95	34.69	2.10	33.32	1.37	65.24%
39	31.64		II	120	34.67	3.03	32.28	2.39	33.00%
40	31.89		11	Bot	35 • 35	3.46	33.43	1.92	55.49%
11	31.07	SG	123	Bot	34.53	3.46	33.07	1.46	42.19%
12	32.30	SG	119	Top	34.29	1.99	32.71	1.58	79 • 39%
13	32.24		11	150	35.915	3.675	33.595	2.32	63.13%
14	32.32		11	170	35 • 39	3.07	32.72	2.67	86.97%
15	32.015		!!	200	35.05	3.035	32.67	2.38	78.42%

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16	32.17	SG	119	Bot	34.79	2.62	34.07	0.72	27.48%
17	31.60	SG	118	Top	35.095	3.495	32.54	2.555	73.10%
18	31.16		11	260	33.69	2.53	31.45	2.24	88.54%
19	32.02		11	290	34.77	2.75	32.75	2.02	73.45%
20	32.06	SG	117	Top	35.295	3.235	32.72	2.575 ·	79.60%
21	32.04		11	325	34.39	2.35	33.43	0.96	40.85%
22	32.45		11	335	34.70	2.25	32.915	1.785	79•33%
23	32.01		11	Bot	34.69	2.68	33.27	1.42	52.98%
24	32.25	SG	116	Top	36.175	3.925	33.00	3.175	80.89%
25	31.48	SG	115	Top	34.29	2.81	32.40	1.89	67.26%
26	32.76		11	150	34.72	1.96	33.70	1.02	52.04%
27	32.11		11	180	35.41	3.30	32.76	2.65	80.30%
28	32.63		11	Bot	35.53	2.90	34.72	0.81	27.93%
29	32.02	SG	114	Top	34.405	2.385	32.73	1.675	70.23%
30	32.35		H .	Bot	34.93	2.58	34.84	0.09	03.49%
31	32.26	SG	134	Top	34.105	1.845	32.70	1.405	76.15%
32	32.615		11	110	34.89	2.275	33.31	1.58	69.45%
33	30.90		11	170	34.65	3.75	33.55	1.10	29.33%

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Beaker No.	Beaker Wt.	Sa	mple No	(m)	Beaker & Samp	Sample Wt.	Wt. After Digestion	Wt. Loss	CaCo ₃ % Wt.Loss, Wt.x100
11	31.07	SG	151	Top	32.94	1.87	31.70	1.24	66.31%
12	32,30		11	Bot	34.38	2.08	33.61	0.77	37.01%
13	32.24	SG	152	Top	34.44	2.20	34.175	0.265	12.04%
14	32.32	SG	154	Top	34.09	1.77	33.64	0.45	25.42%
15	32.015		11	Bot	34.73	2.715	34.69	0.04	1.47%
16	32.17	SG	155	Top	34.35	2.18	33.54	0.81	37.15%
17	31.60	SG	156	Top	33•39	1.79	32.355	1.035	57.82%
18	31.16		11	112	33.85	2.69	33.825	0.025	0.92%
19	32.02	SG	157	Top	35.265	3.245	34.88	0.385	11.86%
20	32.06		11	150	36.54	4.48	36.43	0.11	2.45%
21	32.04	SG	158	Тор	34.18	2.14	33•345	0.835	39.01%
22	32,45	SG	159	Top	34.29	1.84	33.62	0.67	36.41%
23	32.01		11	Bot	36.925	4.915	35•93	0.995	20.24%
24	32.25	SG	160	Top	34.47	2.22	33.47	1.00	45.04%
25	3 1. 48	SG	161	Top	33.655	2.175	32.76	0.895	41.15%
26	32.76	SG	162	Top	35.95	3.19	35.30	0.65	20.37%
27	32.11		ti	150	35.03	2.92	33.85	1.18	40.41%
28	32.63	SG	163	Top	34.64	2.01	33.96	0.68	33.83%
29	32.02		tt	80	34.65	2.63	34.62	0.03	1.14%
30	32.35		11	Bot	37.38	5.03	37.37	0.01	0.20%
31	32.26	SG	16½	Top	34.935	2 , 675	33.56	1.375	51.40%
32	32.615		11	Bot	34.58	1.965	34.55	0.03	1.52%
33	30.90	SG	165	Top	33.29	2.39	32.59	0.70	29.29%
34	31.935	SG	54	Top	34.01	2.075	33.44	0.57	27.47%
35	31.90		11	180	34.37	2.47	33.79	0.58	23.48%
36	32.365		11	Bot	36.075	3.71	34.47	1.605	43.26%
37	32.26	SG	55	Top	34.575	2.315	33.29	1.285	55.50%
38	32.59		11	Bot	35.53	2.94	34.755	0.775	26.36%
39	31.64	SG	56	Top	34.74	3.10	34.285	0.455	14.68%

(50) 40	29.65	SG 57 Top	32.05	2.40	30.94	1.11	46.25%
41	30,04	" Bot	32.38	2.34	32.19	0.19	8.12%
42	30.45	SG 58 Top	32.53	2.08	32.105	0.425	20.43%
43	29.94	SG 61 Top	32.05	2.11	31.16	0.89	42.18%
1+1+	30.02	SG 62 Top	32.53	2.51	31.525	1.005	40.04%
45	30.85	SG 63 Top	32.48	1.63	31.675	0.805	49.38%
46	29.50	SG 65 Top	31.17	1.67	30.75	0.42	25.15%
47	30.07	SG 66 Top	33.83	3.76	32.12	1.71	45.48%

Core	Core	BYT.	Q.	BI			FORAMS.			!	QTZ.	4	ECH.	MISC.	TOT.
No.	Alg.		N.H.G		Ç+0	H	Ħ	Ъ.	ė	N.		needles	Spines		Forams
221	ر د	t	0,	17.8	ر ش	2.0	1.6	0.2	ı	,	35.2	0•9	0.4	25.6	7.3
222	ı	ı	0.0	34.8	2.0	0.2	2.0	ı	1	ı	14,0	15.0	1.0	30.0	7, 4
223	1.2	0.2	2.4	29.2	9.	0,2	1.0	1	ı	ı	14.40	17.0	1.6	30.8	3.2
224	0,0	6.0	13.44	23.63	6.0	2.54	3.27	0.18	0.2	t	21.81	11.63	2,18	18,8	7.45
225	0.19	0.39	44.5	23.52	0.19	0.58	0.39	ı	1	1	31.76	13.52	2.94	20.39	1.74
226	7.	0.2	7.5	30.2	1.2	ᡮ•0	1.6	ı	ı	ı	23.2	12,2	0.9	19.6	3.2
227	9.0	9.0	6. ب	33.4	ᡮ•0	7.	8.0	ı	1	ı	23.0	9.6	7.0	21.6	2.6
228	0.18	0.36	2.72	26.9	0.18	06.0	1.09	ı	i	1	06.04	6.36	1.09	19.27	2.17
230	1.2	0.8	ν. 8	55.2	0,8	1.00	7,	1	ı	1,2	12.0	3.0	1.0	16.6	† •†
231	1.2	1.2	9.0	29.6	†•0	0.8	ᡮ•0	ı	ı	2.	37.2	5.0	0.	21.2	3.0
232	1.04	0.34	2.08	32.34	69.0	0.86	0.34	ı	ŧ	1	23.65	8.52	3.13	26.78	2.06
281	0.72	· 1	1.26	27.45	ı	0.54	2.54	1	ı	1	21.45	11.63	4.72	29.45	3.26
282	1.0	ī	2.6	33.2	0.2	7.	9.0	1	1	ı	20.8	5.8	5.0	29.6	2.0
284	1	t	1.6	50.0	0.8	7, 1	2.0	0.2	٠,0	1	10.2	14.6	ተ•0	17.6	5.6
285	1.92	0.19	3.06	32.69	ı	0.19	96.0	ı	1	t	19.23	8.65	3.26	29.61	1.34
291	2.36	i	1.99	21,81	ı	0.54	3.09	ı	8,364,72	4.72	5.45	1.63	14,00	35.45	17.25
293	1.0	ı	4,6	746.60	9.0	0.7	2.0	٥.	ı	ı	7.0	3.40	5.0	22.60	8.6

RESULTS FROM ORGANIC AND NON ORGANIC PARTICLE COUNTS.

86	87	124	125	126	127	128	129	130	132	131	133	237	238	239	240	241	292	Core
ı	1.1	0.7	0.3	28.0	2.4	1	2.6	20	5°3	1 • 0	0,8	0.6	0.6	1.6	2.6	3.6	0	Core
ı	0.2	ı	ı	ο ω	0.2	0.4	ŝ	i	ı	0.6	ı	4.0	ł	4.0	4.0	0,8	0.2	BRY.
4.0	1.2	0.9	0.9	ν ω	2.2	2.4	57 • 8	0.6	3•4	0.4	1.0		I	0.6	1.6	∞ •	<u>~</u> ∞	H.C.+
37.2	54.0	41.6	27.6	1.2.5	37.8	8.04	36.0	32.4	29.4	36.0	43.6	25.6	15.4	40.8	26.4	30.6	20.4	BI.
ı	0.2	ı	ı	i	1.0	0.2	0.6	0.2	0.6	٠ 4.	1	4,0	4.0	0.2	i	ı	÷+	S+Q.
ı	4.0	о ;т	ı	ł	1.0	1.0	1.6	1 • }	1	3 . 4	J. 6	∞ • -7	0 • 8	0.6	N. 0	0.8	0	H
2.6	1.8	1 •2	ı	1 • +	• 6	1.0	3.4	2.0	л °0	0.4	0.8	٠, ١	· 4	0.8	0.6	2.2	ω œ	FORAMS
ı	1	0.1	ŧ	1	0.2	0.2	0.6	0.2	ı	ı	1	ı	1	I,	ı	ı	0.2	PO .
1	0.4 -	1	ŧ	1	t	1	2.6 5.3	3.0 4.0	10,010,6	1	1	1	1	0.8 5.3	0.8 5.3	2.8 7.3	5.28.0	D. N.
14.0	9.8	18.7	30.9	6.5	22.4	18.4	7.0	12.0	1 • 8	28.4	21.0	34.0	50.0	18.8	22.0	4.9	10.8	QTZ.
0.8	12.6	14.3	8.1	6.3	7.2	4.2	3.4	15.0	2 8	3 • 2	3.0	2.4	57 +	8.6	2.0	2.2	1 • 0	A. Needles
15.0	1.8	5 8	2.0	- + +	3.2	11.4	+ • +	1 •}	0.4	3 2	1.8	N • •	3.0	·1 •2	6.0	16.8	12.6	ECH. Spines
30.0	16.2	15.8	30.0	7.8	20.2	18.6	26.6	25.8	27.0	23.0	22.4	31.0	24.0	22,2	30.2	24.6	30.8	ECH. MISC.
2.6	N •	2.0	ı	1.8	4.1	4.2	14.1	10.8	26.2	4. 2	4.9	2.6	1.6	5.7	8.7	13.1	18.4	TOT.

Allas

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57 6	55	45	165	164	163	162	161	160	159	158	157	156	155	151	152	151	Core No.
ı	1.0	0.2	1	-1 • 0	0.9	1	0.2	ı	ı	ů	4.0	N ,0	0.9	1.	ŧ	~ • ∞	Core
ı	ı	ı	ı	4.0	4.0	1	i	1	ı	ı	0.2	0,18	0.2	0.2	ı	4.0	BRY.
1.	N 8	7.2	14.2	ω •	7.0	0	1 • 0	<u>-1</u>	ı	2.2	0.2	3.58	0.4		0.2	ω. 	T.G.+
7.6	10.6	8.6	16.4	20.6	18.0	0.8	22.4	16.8	12.0	13.8	1 • 0	24,36	4.92	13.4	0.6	19 2	BI.
•	ı	0.2	ı	0.8	1.6	ı	ı	2,4	2.6	ı	ı	0.36	i	2.0	0.2	1	© + \
- • 0	0.6	ı	1	-∠	3.6	ı	1,2	1.0	, L	ł	2.0	0.73	•	2.6	ı	0,6	H
0 • 2	4,0	4.0	1	0,2	0.2	ı	0,2	0.2	4.0	4.0	0.2	0.9	N	0,6	ı	<u>,</u>	FORAMS
0	0.2	0.2	0.4	3°+	• 0	0.4		0.2	4.0	-1 - -	0.2	0.36	0.8	1	1	0	P .
1	4.0	0.2	ı	. 4.0	0.6	ı	1.0	1	ı	0.8	î	0.36	ı	0.6	ı	2	
1	1	i	1	• 0	0.9	f	ı	ł	ŀ	ώ	f	l	0,9	ł	ł	-1 •⊗	N.
85.2	74.0	82.2	69.0	58.0	53.4	97.8	63.2	54.6	66.2	65.2	95.6	50.3	53.4	68.8	98.0	34.0	QTZ.
-1 • O	3.0	0.6	ı	N	0.8	1	2.6	8.0	5.0	1.6	ı	6.0	1.6	ν ω		\$ 8	A. Needles
4.0	2.0	0 • N	t	1 •}	3.4	0.2	-\ - +	4.0	4.0	4.8	0.2	1 8	₀ 8	2.0	1	18.6	ECH Spines
3.0	5.0	ı	1	5.6	5.2	•	5.0	15.4	1 1 0	3.6	1	0	5. 8	5.6	ı	10.0	MISC.
1.4	1.6	1 • 0	4.0	7.0	10.9	4.0	•	ω • 8	9.	3.9	2.4	2.7	٠.	5.8	0,2	6.8	TOT.

/Cont.

Cont.

66	50	6μ	62	67	58	57
0.8	ı	ı	ı	ŧ	0.6	1
1	0.2	t	+	t	4.0	ı
ဝ တ	-1 -6	0.4	1 • 0	1.0	† -1	6.4
23.6	16.4	16.4	12.4	15.0	12.6	23.4
ı	ı	0.2	1 • ‡	0,8	0.2	0.2
0.6	0.2	1.0	N	-7	‡. 	4.0
i	ŧ	.0	1 , 2	1.0	0.2	0 &
0.	ı	0.2	4.0	0.2	0.2	7.2
ŧ	0.2	0.2	ŧ	1	i	1.+
ı	ŧ	ì	ı	ı	ı	ı
71.0	75.2	50.4	59.4	57.8	79.4	53.0
-1 -6	1.6	11.4	10.0	8.0	ı	÷,
0.2	0.1	1 6	0.6	1		
			7	1	N • 8	S
•	ν. ω	17.2	2	Ņ	œ	, O

ABBREVIATIONS USED.

TOT. Forams	ECH. Spines	QTZ.	D.	.	₽•	Forams	T.C.+ N.T.C. 2	Cor. Alg.
1	ı	ı	ı	1	1	ı	₩	ı
Total Foraminifera	Echinoid Spines and Plates	Quartz	Discorbis and Rosalina	Elphidium	Quinqueloculina	Foraminifera	Turreted and Non Turreted Castropods	Coralline Algae
	MISC.	A. Needles	N.	^ነ ህ •	T•	• •	BI.	BRY.
	ı	ı	ı	i	ı	ı	1	i
	Miscellaneous	Aragonite Needles.	Nubecularia	Peneroplis and Spirolina	Triloculina	Spiroloculina	Bivalves	Bryozoa

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